

1 **An analytical verification test for numerically simulated convective flow**
2 **above a thermally heterogeneous surface**

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4
5 **Abstract.** An analytical solution of the Boussinesq equations for the motion of a
6 viscous stably stratified fluid driven by a surface thermal forcing with large horizontal
7 gradients (step changes) is obtained. This analytical solution is one of the few available
8 for wall-bounded buoyancy-driven flows. The solution can be used to verify that
9 computer codes for Boussinesq fluid system simulations are free of errors in formulation
10 of wall boundary conditions and to evaluate the relative performances of competing
11 numerical algorithms. Because the solution pertains to flows driven by a surface thermal
12 forcing, one of its main applications may be for testing the no-slip, impermeable wall
13 boundary conditions for the pressure Poisson equation. Examples of such tests are
14 presented.

15 1 Introduction

16 Thermal disturbances associated with variations in underlying surface properties can
17 drive local circulations in the atmospheric boundary layer (Atkinson, 1981; Briggs, 1988;
18 Hadfield et al., 1991; Segal and Arritt, 1992; Simpson, 1994; Mahrt et al., 1994; Pielke,
19 2001; McPherson, 2007; Kang et al., 2012) and affect the development of the convective
20 boundary layer (Patton et al., 2005; van Heerwaarden et al., 2014). Computational fluid
21 dynamics (CFD) codes for modeling such flows commonly solve the Boussinesq
22 equations of motion and thermal energy for a viscous/diffusive stably stratified fluid. In
23 this paper we present an analytical solution of the Boussinesq equations for flows driven
24 by a surface thermal forcing with large gradients (step changes) in the horizontal. The
25 solution can be used to verify that CFD codes for Boussinesq fluid system simulations
26 are free of errors, and to evaluate the relative performances of competing numerical
27 algorithms. Such verification procedures are important in the development of CFD
28 models designed for research, operational, and classroom applications.

29 We solve the linearized Navier-Stokes and thermal energy equations analytically
30 for the case where the surface buoyancy varies laterally as a square wave (Fig. 1).
31 Attention is restricted to the steady state. No boundary-layer approximations are made;
32 the solution is non-hydrostatic, and both horizontal and vertical derivatives are included
33 in the viscous stress and thermal diffusion terms. The solution is similar to that of
34 Axelsen et al. (2010) for katabatic flow above a cold strip, but is easier to evaluate (no

35 slope present) and applies to the more general scenario where the viscosity and
36 diffusivity coefficients can differ. The flow is also similar to a special case (no slope)
37 considered by Egger (1981), although a final analytical solution was not provided in
38 that study. Strictly speaking, the linearized Navier-Stokes equations apply to a class of
39 very low Reynolds number motions known as creeping flows. Such flows appear in
40 studies of lubrication, locomotion of microorganisms, lava flow, and flow in porous
41 media. Of course, for the task at hand, if our linear solution is to serve as a benchmark
42 for a nonlinear numerical model solution, it is essential that the parameter space be
43 restricted to values for which the model's nonlinear terms are negligible.

44 Because the solution pertains to flows driven by a surface thermal forcing, one of
45 its main applications may be as a test for surface boundary conditions in the pressure
46 Poisson equation. In models of atmospheric boundary-layer flows, the buoyancy is a
47 major contributor to the forcing term in the Poisson equation and also appears in the
48 associated surface boundary condition. The pressure boundary condition on a solid
49 boundary in incompressible (Boussinesq) fluid flows is an important and complex issue
50 that has long been fraught with technical difficulties and controversies (Strikwerda,
51 1984; Orszag et al., 1986; Gresho and Sani, 1987; Gresho, 1990; Temam, 1991; Henshaw,
52 1994; Petersson, 2001; Sani et al., 2006; Rempfer, 2006; Guermond et al., 2006;
53 Nordström et al., 2007; Shirokoff and Rosales, 2011; Hosseini and Feng, 2011; Vreman,
54 2014). Typical fractional-step solution methodologies and associated pressure (or

55 pseudo-pressure) boundary-condition implementations are often verified using various
56 prototypic flows such as Poiseuille flows, lid-driven cavity flows, flows over cylinders or
57 bluff bodies, viscously decaying vortices, and dam-break flows. We are unaware of
58 verification tests in which flows were driven by a heterogeneous surface buoyancy
59 forcing. Our solution is designed to fill this gap.

60 The analytical solution is derived in Sect. 2. In Sect. 3, this solution is compared
61 to numerically simulated fields in a steady state. Two versions of a numerical code are
62 run: a version in which the correct surface pressure boundary condition is applied, and a
63 version in which the pressure condition is mis-specified. A summary follows in Sect. 4.

64

65 **2 Analytical solution**

66 We derive the solution for steady flow over an underlying surface along which the
67 buoyancy varies laterally as a single harmonic function. This single-harmonic solution is
68 then used as a building block in a Fourier representation of the square-wave solution.

69

70 **2.1 Governing equations**

71 Consider the flow of a viscous stably stratified fluid that fills the semi-infinite domain
72 above a solid horizontal surface (placed at $z = 0$). This surface undergoes a steady
73 thermal forcing that varies periodically in the right-hand Cartesian x direction, but is
74 independent of the y direction. The two-dimensional (x, z) flow is periodic in x , and

75 satisfies the linearized (assuming the disturbance is of small amplitude) governing
 76 equations under the Boussinesq approximation,

$$77 \quad 0 = -\frac{\partial \Pi}{\partial x} + \nu \nabla^2 u, \quad (2.1)$$

$$78 \quad 0 = -\frac{\partial \Pi}{\partial z} + b + \nu \nabla^2 w, \quad (2.2)$$

$$79 \quad 0 = -N^2 w + \alpha \nabla^2 b, \quad (2.3)$$

$$80 \quad \frac{\partial u}{\partial x} + \frac{\partial w}{\partial z} = 0. \quad (2.4)$$

81 Apart from notational differences, (2.1)–(2.4) are the two-dimensional steady state
 82 versions of (55)–(57) of Sect. II of Chandrasekhar (1961). Equations (2.1) and (2.2) are
 83 the horizontal (x) and vertical (z) equations of motion, respectively, (2.3) is the thermal
 84 energy equation (differential form of the first law of thermodynamics) expressed in
 85 terms of the buoyancy variable (defined below), and (2.4) is the incompressibility
 86 condition. Here u and w are the horizontal and vertical velocity components, $\Pi \equiv$
 87 $[p - p_e(z)]/\rho_w$ is the kinematic pressure perturbation [p is pressure, $p_e(z)$ is pressure in a
 88 hydrostatic environmental state in which the density profile is $\rho_e(z)$, ρ_w is a constant
 89 reference density, say, $\rho_e(0)$], and $b \equiv -g[\rho - \rho_e(z)]/\rho_w$ is the buoyancy, where ρ is the
 90 actual density, and g is the acceleration due to gravity. The Brunt-Väisälä frequency
 91 $N \equiv \sqrt{-(g/\rho_w)d\rho_e/dz}$ of the ambient fluid (Kundu 1990), kinematic viscosity ν , and
 92 thermal diffusivity α are taken constant.

93 We obtain our solution using a standard vorticity/streamfunction formulation.

94 Cross-differentiating (2.1) and (2.2) yields the vorticity equation,

$$95 \quad 0 = -\frac{\partial b}{\partial x} + \nu \nabla^2 \eta, \quad (2.5)$$

96 where $\eta \equiv \partial u / \partial z - \partial w / \partial x$ is the vorticity. Eliminating b from (2.3) and (2.5) yields

$$97 \quad \nabla^4 \eta = \frac{N^2}{\nu \alpha} \frac{\partial w}{\partial x}. \quad (2.6)$$

98 Introducing a streamfunction ψ defined through

$$99 \quad u = \partial \psi / \partial z, \quad w = -\partial \psi / \partial x, \quad (2.7)$$

100 guarantees that (2.4) is satisfied, and transforms (2.6) into a single equation for ψ ,

$$101 \quad \nabla^6 \psi + \frac{N^2}{\nu \alpha} \frac{\partial^2 \psi}{\partial x^2} = 0. \quad (2.8)$$

102 The dependent variables are assumed to vanish far above the surface ($z \rightarrow \infty$). On the

103 surface we apply no-slip ($u = 0$) and impermeability ($w = 0$) conditions, and specify a

104 periodic (in x) buoyancy distribution. As we will now see, restricting the dependent

105 variables to steady periodic forms that vanish as $z \rightarrow \infty$ also restricts acceptable

106 distributions of the surface buoyancy. The restriction was first noted by Egger (1981,

107 Sect. 3c), though without details. Averaging (2.3) over one period (using $w = -\partial \psi / \partial x$)

108 yields $d^2 \bar{b} / dz^2 = 0$, which integrates to $\bar{b} = A + Bz$ (\bar{b} is the average of b ; A and B are

109 constants). Taking $b \rightarrow 0$ as $z \rightarrow \infty$, implies that $\bar{b} \rightarrow 0$ as $z \rightarrow \infty$, in which case $A =$

110 $B = 0$, and $\bar{b}(z) = 0$. In particular, at the surface, $\bar{b}(0) = 0$. If a surface distribution

111 $b(x,0)$ violates this condition, the ground acts as a net heat source/sink. In an unsteady
 112 model, such a source/sink would force a continually upward-developing disturbance, and
 113 a steady state could never be attained.

114

115 **2.2 Single-harmonic forcing**

116 For a surface buoyancy of the form $b(x,0) \propto \sin kx$, (2.3) indicates that ψ is of the form

$$117 \quad \psi = A(z) \cos kx. \quad (2.9)$$

118 Application of (2.9) in (2.8) yields

$$119 \quad \left(\frac{d^2}{dz^2} - k^2 \right)^3 A - \frac{N^2 k^2}{\nu \alpha} A = 0, \quad (2.10)$$

120 which has solutions of the form $A \propto e^{Mz}$ for M satisfying

$$121 \quad (M^2 - k^2)^3 = \frac{N^2 k^2}{\nu \alpha}. \quad (2.11)$$

122 Taking the one-third power of (2.11) yields a useful intermediate result:

$$123 \quad M^2 - k^2 = \frac{N^{2/3} k^{2/3}}{\nu^{1/3} \alpha^{1/3}} e^{2n\pi i/3}, \quad (2.12)$$

124 where n is an integer. Rearranging (2.12) and taking the square root yields

$$125 \quad M = \pm \sqrt{k^2 + \frac{N^{2/3} k^{2/3}}{\nu^{1/3} \alpha^{1/3}} e^{2n\pi i/3}}. \quad (2.13)$$

126 Equation (2.13) furnishes six roots, two for each of $n = 0, 1, 2$. To ensure that $A(z) \rightarrow 0$

127 as $z \rightarrow \infty$, we reject the roots with a positive real part. With the radicand of (2.13)

128 expressed in polar form, the physically acceptable roots are

$$129 \quad M_0 = -\sqrt{k^2 + \frac{N^{2/3}k^{2/3}}{\nu^{1/3}\alpha^{1/3}}}, \quad (n = 0), \quad (2.14a)$$

$$130 \quad M_1 = -r^{1/2}e^{i\phi/2}, \quad (n = 1), \quad (2.14b)$$

$$131 \quad M_2 = -r^{1/2}e^{-i\phi/2}, \quad (n = 2), \quad (2.14c)$$

132 where the subscript on M denotes the associated value of n , and r and ϕ are defined by

$$133 \quad r \equiv \sqrt{\left[k^2 + \frac{N^{2/3}k^{2/3}}{\nu^{1/3}\alpha^{1/3}} \cos\left(\frac{2\pi}{3}\right) \right]^2 + \left[\frac{N^{2/3}k^{2/3}}{\nu^{1/3}\alpha^{1/3}} \sin\left(\frac{2\pi}{3}\right) \right]^2}, \quad (2.15)$$

$$134 \quad \cos \phi = \frac{1}{r} \left[k^2 + \frac{N^{2/3}k^{2/3}}{\nu^{1/3}\alpha^{1/3}} \cos\left(\frac{2\pi}{3}\right) \right], \quad \sin \phi = \frac{1}{r} \left(\frac{N^{2/3}k^{2/3}}{\nu^{1/3}\alpha^{1/3}} \right) \sin\left(\frac{2\pi}{3}\right) > 0. \quad (2.16)$$

135 While solving (2.16) for ϕ , care must be taken when evaluating arcsin or arccos

136 functions that ϕ appears in the correct quadrant (ϕ should be in quadrant I or II so

137 $\phi/2$ should always be in quadrant I). Also note from (2.14b) and (2.14c) that M_2 is the

138 complex conjugate of M_1 ($M_2 = M_1^*$), a fact that will often be used below.

139 With the general solution for ψ written as

$$140 \quad \psi = (B e^{M_0 z} + C e^{M_1 z} + D e^{M_2 z}) \cos kx, \quad (2.17)$$

141 where B , C , and D are constants, the vorticity becomes,

$$142 \quad \eta = \left[B(M_0^2 - k^2) e^{M_0 z} + C(M_1^2 - k^2) e^{M_1 z} + D(M_2^2 - k^2) e^{M_2 z} \right] \cos kx, \quad (2.18)$$

143 and the buoyancy follows from (2.3) as

144
$$b = \frac{kN^2}{\alpha} \left(\frac{B}{M_0^2 - k^2} e^{M_0 z} + \frac{C}{M_1^2 - k^2} e^{M_1 z} + \frac{D}{M_2^2 - k^2} e^{M_2 z} \right) \sin kx + b_h, \quad (2.19)$$

145 where $\nabla^2 b_h = 0$. In view of (2.12), equation (2.19) becomes

146
$$b = \frac{k^{1/3} \nu^{1/3} N^{4/3}}{\alpha^{2/3}} (B e^{M_0 z} + e^{-2\pi i/3} C e^{M_1 z} + e^{-4\pi i/3} D e^{M_2 z}) \sin kx + b_h. \quad (2.20)$$

147 Applying (2.18) and (2.20) in (2.5) yields an equation for $\partial b_h / \partial x$, which upon use of

148 (2.12) and $M_2 = M_1^*$ reduces to $\partial b_h / \partial x = 0$. So b_h is, at most, a function of z . Since

149 $\nabla^2 b_h = 0$, b_h is, at most, a linear function of z , and since b should vanish as $z \rightarrow \infty$,

150 that linear function must be 0. Thus, $b_h = 0$.

151 The pressure follows from (2.1) and (2.12) as

152
$$\Pi = \frac{\nu^{2/3} N^{2/3}}{k^{1/3} \alpha^{1/3}} (B M_0 e^{M_0 z} + C M_1 e^{2\pi i/3} e^{M_1 z} + D M_2 e^{4\pi i/3} e^{M_2 z}) \sin kx + G(z), \quad (2.21)$$

153 where $G(z)$ is a function of integration. Applying (2.21) in (2.2), and using (2.11) yields

154 $dG/dz = 0$, so G is constant. For Π to vanish as $z \rightarrow \infty$, this constant must be zero.

155 The surface conditions determine B , C , and D . The surface buoyancy is

156
$$b(x, 0) = b_0 \sin kx, \quad (2.22)$$

157 where b_0 is a constant forcing amplitude. Application of (2.20) in (2.22) yields

158
$$B + e^{-2\pi i/3} C + e^{-4\pi i/3} D = \frac{b_0 \alpha^{2/3}}{k^{1/3} \nu^{1/3} N^{4/3}}. \quad (2.23)$$

159 In view of (2.7) and (2.17), the impermeability condition $w(x, 0) = 0$ and no-slip

160 condition $u(x,0) = 0$ yield

$$161 \quad B + C + D = 0, \quad (2.24)$$

$$162 \quad BM_0 + CM_1 + DM_2 = 0. \quad (2.25)$$

163 Straightforward but lengthy manipulations yield the solution of (2.23)–(2.25):

$$164 \quad B = - \left(\frac{b_0 \alpha^{2/3}}{\sqrt{3} k^{1/3} \nu^{1/3} N^{4/3}} \right) \frac{2r^{1/2} \sin(\phi/2)}{M_0 + 2r^{1/2} \cos(\pi/3 + \phi/2)}, \quad (2.26)$$

$$165 \quad C = -i \left(\frac{b_0 \alpha^{2/3}}{\sqrt{3} k^{1/3} \nu^{1/3} N^{4/3}} \right) \frac{M_2 - M_0}{M_0 + 2r^{1/2} \cos(\pi/3 + \phi/2)}, \quad (2.27)$$

$$166 \quad D = i \left(\frac{b_0 \alpha^{2/3}}{\sqrt{3} k^{1/3} \nu^{1/3} N^{4/3}} \right) \frac{M_1 - M_0}{M_0 + 2r^{1/2} \cos(\pi/3 + \phi/2)}. \quad (2.28)$$

167 Applying (2.26)–(2.28) in (2.17), (2.20), and (2.18), with (2.12) used in the latter
 168 equation, and noting that B is real, while $D = C^*$ (since $M_2 = M_1^*$), we obtain

$$169 \quad b = \frac{2b_0}{\sqrt{3}} \frac{e^{-Z_c} [\mu \cos(Z_s + \pi/6) + \cos(Z_s + \pi/6 + \phi/2)] - e^{M_0 z} \sin(\phi/2)}{\mu + 2 \cos(\pi/3 + \phi/2)} \sin kx, \quad (2.29)$$

$$170 \quad \psi = \frac{2b_0 \alpha^{2/3}}{\sqrt{3} k^{1/3} \nu^{1/3} N^{4/3}} \frac{e^{-Z_c} [\mu \sin Z_s + \sin(Z_s + \phi/2)] - e^{M_0 z} \sin(\phi/2)}{\mu + 2 \cos(\pi/3 + \phi/2)} \cos kx, \quad (2.30)$$

171 where

$$172 \quad Z_s \equiv z r^{1/2} \sin(\phi/2), \quad Z_c \equiv z r^{1/2} \cos(\phi/2), \quad \mu \equiv M_0 / r^{1/2}. \quad (2.31)$$

173 Application of (2.30) in (2.7) yields the velocity components as

174
$$u = \frac{2b_0 \alpha^{2/3} r^{1/2}}{\sqrt{3} k^{1/3} \nu^{1/3} N^{4/3}} \frac{e^{-Z_c} [\mu \sin(\phi/2 - Z_s) - \sin Z_s] - \mu e^{M_0 z} \sin(\phi/2)}{\mu + 2 \cos(\pi/3 + \phi/2)} \cos kx \quad (2.32)$$

175
$$w = \frac{2b_0 \alpha^{2/3} k^{2/3}}{\sqrt{3} \nu^{1/3} N^{4/3}} \frac{e^{-Z_c} [\mu \sin Z_s + \sin(Z_s + \phi/2)] - e^{M_0 z} \sin(\phi/2)}{\mu + 2 \cos(\pi/3 + \phi/2)} \sin kx. \quad (2.33)$$

176

177 **2.3 Piecewise constant (square wave) forcing**

178 Next, consider the case where the surface buoyancy varies horizontally as a square

179 wave, with a distribution over one period L given by

180
$$b(x, 0) = \begin{cases} b_{\max}, & 0 < x < L/2, \\ -b_{\max}, & L/2 < x < L. \end{cases} \quad (2.34)$$

181 Such a distribution can be expressed as the Fourier series:

182
$$b(x, 0) = \sum_{n=1}^{\infty} b_n \sin\left(\frac{n\pi x}{L}\right), \quad (2.35)$$

183
$$b_n = \frac{2}{L} \int_0^L b(x, 0) \sin\left(\frac{n\pi x}{L}\right) dx. \quad (2.36)$$

184 Application of (2.34) in (2.36) yields

185
$$b_n = \frac{2b_{\max}}{n\pi} [1 - 2 \cos(n\pi/2) + \cos(n\pi)]. \quad (2.37)$$

186 The solutions for b , ψ , u , and w can then be written as summations over the single-

187 harmonic solutions (2.29), (2.30), (2.32), and (2.33), with k related to n by

188
$$k = \frac{n\pi}{L}, \quad (2.38)$$

189 and with b_0 replaced by b_n :

$$190 \quad b = \frac{2}{\sqrt{3}} \sum_{n=1}^{\infty} b_n \frac{e^{-Z_c} [\mu \cos(Z_s + \pi/6) + \cos(Z_s + \pi/6 + \phi/2)] - e^{M_0 z} \sin(\phi/2)}{\mu + 2 \cos(\pi/3 + \phi/2)} \sin\left(\frac{n\pi x}{L}\right), \quad (2.39)$$

$$191 \quad \psi = \frac{2\alpha^{2/3}}{\sqrt{3} \nu^{1/3} N^{4/3}} \sum_{n=1}^{\infty} \frac{b_n}{k^{1/3}} \frac{e^{-Z_c} [\mu \sin Z_s + \sin(Z_s + \phi/2)] - e^{M_0 z} \sin(\phi/2)}{\mu + 2 \cos(\pi/3 + \phi/2)} \cos\left(\frac{n\pi x}{L}\right), \quad (2.40)$$

$$192 \quad u = \frac{2\alpha^{2/3}}{\sqrt{3} \nu^{1/3} N^{4/3}} \sum_{n=1}^{\infty} b_n \frac{r^{1/2}}{k^{1/3}} \frac{e^{-Z_c} [\mu \sin(\phi/2 - Z_s) - \sin Z_s] - \mu e^{M_0 z} \sin(\phi/2)}{\mu + 2 \cos(\pi/3 + \phi/2)} \cos\left(\frac{n\pi x}{L}\right), \quad (2.41)$$

$$193 \quad w = \frac{2\alpha^{2/3}}{\sqrt{3} \nu^{1/3} N^{4/3}} \sum_{n=1}^{\infty} b_n k^{2/3} \frac{e^{-Z_c} [\mu \sin Z_s + \sin(Z_s + \phi/2)] - e^{M_0 z} \sin(\phi/2)}{\mu + 2 \cos(\pi/3 + \phi/2)} \sin\left(\frac{n\pi x}{L}\right). \quad (2.42)$$

194

195 **3 Verification tests**

196 A solution of the linearized equations may be used to verify a nonlinear code if the
 197 nonlinear terms are sufficiently small. Unfortunately, *a priori* estimates of such terms
 198 expressed, for example, through a Reynolds number, are not straightforward since the
 199 relevant velocity and length scales in our problem are only evident after a solution has
 200 been obtained. We thus seek an appropriate set of test parameters through trial and
 201 error, guided by *a posteriori* linear solution estimates of the terms $\mathbf{u} \cdot \nabla b$ and $\mathbf{u} \cdot \nabla \eta$
 202 [$\mathbf{u} = (u, w)$] present in nonlinear versions of (2.3) and (2.5), respectively. Specifically, for
 203 any computed candidate solution, we formed the ratios of the largest values of those
 204 nonlinear terms to the largest values of the corresponding linear terms, that is, the

205 terms actually present in (2.3) and (2.5). We need only consider one such linear term
 206 per ratio since (2.3) and (2.5) are comprised of two terms of equal magnitude. A
 207 solution was deemed to be sufficiently linear if

$$208 \quad R_\eta \equiv \frac{\max|\mathbf{u} \cdot \nabla \eta|}{\max|\partial b / \partial x|} < \varepsilon, \quad \text{and} \quad R_b \equiv \frac{\max|\mathbf{u} \cdot \nabla b|}{\max|\alpha \nabla^2 b|} < \varepsilon, \quad (3.1)$$

209 where ε ($\ll 1$) is a prescribed threshold. The suitability of this approach was
 210 confirmed by the very close agreement between the analytical solutions and the
 211 numerical solutions obtained with the correct surface pressure condition.

212 The numerical model employed in our tests is a variant of a direct numerical
 213 simulation (DNS) code used in the boundary-layer and slope-flow studies of Fedorovich
 214 et al. (2001), Fedorovich and Shapiro (2009a,b), and Shapiro and Fedorovich (2013,
 215 2014). The model solves the Boussinesq governing equations on a staggered (Arakawa
 216 C) grid. Although designed for three-dimensional simulations, the model was run in a
 217 two-dimensional (x, z) mode. The overall solution procedure is patterned on a fractional
 218 step method proposed by Chorin (1968). In our version, the prognostic equations are
 219 integrated using a filtered leapfrog scheme with explicit treatment of the viscous term.
 220 The pressure is diagnosed from a Poisson equation (equation (A3b), discussed in the
 221 Appendix), which is solved using a fast Fourier transform technique in horizontal
 222 planes, and a tridiagonal matrix inversion in the vertical. The surface condition on
 223 pressure is the inhomogeneous Neumann condition (INC) that arises from projecting the

224 vertical equation of motion into the vertical, and imposing the impermeability condition
 225 (Vreman, 2014; also see the Appendix). We also run a version of the code in which the
 226 surface pressure condition is mis-specified as a homogeneous Neumann condition (HNC).
 227 We hasten to add, however, that our implementation of the HNC may be quite different
 228 from implementations described in the literature. We elaborate on these technical
 229 differences and review general aspects of the problem of surface pressure specification in
 230 the Appendix.

231 The analytical solution was evaluated on an un-staggered (x, z) grid extending
 232 over one period of the square wave ($x = 0$ to $x = L$). The series were truncated at
 233 50000 terms. The governing parameters were adjusted so that the linearity criteria
 234 were satisfied in comparisons with $\varepsilon = 5 \times 10^{-3}$.

235 In the first test, we set $\nu = \alpha = 0.001 \text{ m}^2 \text{ s}^{-1}$, $N = 0.02 \text{ s}^{-1}$, $L = 5.12 \text{ m}$, and b_{\max}
 236 $= 1 \times 10^{-5} \text{ m s}^{-2}$. For the analytical solution A-1, the (x, z) grid consisted of 513 points
 237 in the x direction and 1025 points in the z direction, with grid spacings
 238 $\Delta x = \Delta z = 0.01 \text{ m}$. The linearity criteria (3.1) were satisfied with $R_\eta \cong 8.2 \times 10^{-5}$ and
 239 $R_b \cong 2.8 \times 10^{-3}$. The analytical b and w fields shown in Fig. 2 depict a broad zone of
 240 ascent above the warm surface and a compensating zone of descent over the cold
 241 surface, roughly for $z < 1.8 \text{ m}$. In the upper part of these zones (at roughly
 242 $0.9 \text{ m} < z < 1.8 \text{ m}$), adiabatic expansion/compression has reversed the senses of the

243 buoyancy fields. Surprisingly, the numerical fields in the inhomogeneous INC-1 and
 244 homogeneous HNC-1 cases are very similar to each other and to the A-1 fields. The u
 245 fields from A-1, INC-1, and HNC-1 shown in Fig. 3 are visually indistinguishable from
 246 one another.

247 To understand why the INC-1 and HNC-1 simulations are so similar, and to
 248 identify simulation parameters that might evince more substantial differences, we
 249 consider the idealized problem in which a specified buoyancy $b = b_0 e^{-\gamma z} \sin kx$ ($\gamma = h^{-1}$,
 250 where h is the e-folding depth scale) is the only forcing term in the Poisson equation
 251 $\nabla^2 \Pi = \partial b / \partial z$, with Neumann surface condition $\partial \Pi / \partial z|_0 = b(x, 0)$. This idealized
 252 problem is solved as

$$253 \quad \Pi_{\text{INC}}^* = \frac{b_0}{\gamma^2 - k^2} \left(k e^{-kz} - \gamma e^{-\gamma z} \right) \sin kx. \quad (3.2)$$

254 The corresponding solution obtained with the homogeneous Neumann condition,
 255 $\partial \Pi / \partial z|_0 = 0$, is

$$256 \quad \Pi_{\text{HNC}}^* = \frac{b_0}{\gamma^2 - k^2} \left(\frac{\gamma^2}{k} e^{-kz} - \gamma e^{-\gamma z} \right) \sin kx. \quad (3.3)$$

257 The relative error (RE) in the vertical pressure gradient force associated with (3.2) and
 258 (3.3), defined as the local absolute error in that force divided by the local buoyancy, is
 259 calculated as

260
$$RE \equiv \left| \frac{\partial \Pi_{\text{INC}}^* / \partial z - \partial \Pi_{\text{HNC}}^* / \partial z}{b} \right| = e^{(a-1)kz}, \quad (3.4)$$

261 where $a \equiv \gamma/k$. Written in terms of the depth scale h and wavelength $\lambda = 2\pi/k$, a can
 262 be interpreted as an aspect ratio characterizing the width to depth scales of the
 263 disturbance, $a = \lambda/(2\pi h) \propto \lambda\gamma$. From (3.4) we see that RE decreases exponentially with
 264 z for disturbances characterized by small aspect ratios, $a < 1$ (which we refer to as deep
 265 disturbances) and increases exponentially with z for disturbances characterized by large
 266 aspect ratios, $a > 1$ (which we refer to as shallow disturbances). The buoyancy in Fig. 2
 267 is suggestive of $a < 1$, which indicates that the first test could be classified as a deep
 268 (error-forgiving) simulation.

269 The preceding analysis suggests that simulations with shallow thermal
 270 disturbances ($a > 1$) might yield large differences between cases with inhomogeneous
 271 and homogeneous Neumann conditions. There did not appear to be a straightforward
 272 way to increase the effective a by systematically varying the parameters (e.g., increasing
 273 L tended to increase the effective h), but a set of suitable parameters were identified
 274 through trial and error and were used as the basis for the second test case.

275 In the second test, we set $\nu = \alpha = 0.0001 \text{ m}^2 \text{ s}^{-1}$, $N = 0.2 \text{ s}^{-1}$, $L = 10.24 \text{ m}$, and
 276 $b_{\text{max}} = 5 \times 10^{-6} \text{ ms}^{-2}$. The analytical solution A-2 was generated with 2049 points in the
 277 x direction and 513 points in the z direction, with grid spacings of $\Delta x = \Delta z = 0.005 \text{ m}$.
 278 The linearity criteria were satisfied with $R_\eta \cong 4.8 \times 10^{-5}$ and $R_b \cong 3.8 \times 10^{-3}$. In

279 contrast to the counter-rotating convection rolls seen in the first test, the analytical b
280 and w fields shown in Fig. 4 depict narrow updraft/downdraft pairs straddling the
281 buoyancy discontinuities. Between the narrow updrafts is a broad region of relatively
282 weak ascent. The w and b fields above the cold surface are mirror images of the fields
283 above the warm surface. Note the change in the scales of the x and (especially) the z
284 axes between Figs. 4 and 2: the low-level thermal disturbance in the second test is much
285 shallower than the disturbance in the first test (and is suggestive of $a > 1$). In this
286 second test case we find dramatic differences between the inhomogeneous INC-2 and
287 homogeneous HNC-2 cases. Specifically, while the INC-2 and A-2 fields are in excellent
288 agreement, the HNC-2 fields showed no signs of even approaching a steady state. Long
289 after the INC-2 simulation had reached a steady state, the HNC-2 fields continued to
290 amplify and develop asymmetric structures associated with flow nonlinearities. The very
291 close agreement between the A-2 solution and the steady state in the INC-2 simulation
292 is shown for the u field in Fig. 5. The u field in the disastrous HNC-2 simulation, at a
293 time when a steady state had already been attained in the INC-2 simulation, is shown
294 in Fig. 6.

295

296 4 Summary

297 The linearized Boussinesq equations for the motion of a viscous stably stratified fluid
298 are solved analytically for a surface buoyancy that varies laterally as a square wave.

299 The solution describes two-dimensional laminar convective structures such as thermal
300 convective rolls and updraft/downdraft pairs. The main applications of the solution may
301 be in code verification and the evaluation of different implementations of the surface
302 pressure condition for the pressure Poisson equation. Tests have been conducted for
303 cases where the aspect ratios of the thermal disturbance have been large and small.
304 With attention restricted to disturbances of sufficiently small amplitude, the linear
305 solution and numerically simulated fields with the inhomogeneous Neumann condition
306 for pressure (which is appropriate in the context of the particular fractional step
307 procedure adopted in our DNS code) have been found to be in excellent agreement for
308 both tests. However, in tests with a mis-specified Neumann condition, an excellent
309 agreement with the analytical solution has been found only for the deep (small aspect
310 ratio) disturbance case; errors in the shallow (large aspect ratio) disturbance case have
311 been catastrophic.

312 **Appendix A: Comment on the pressure condition at a lower solid surface**

313 Consider a three-dimensional Boussinesq system with equation of motion,

314
$$\frac{\partial \mathbf{u}}{\partial t} = -\nabla \Pi + \nu \nabla^2 \mathbf{u} + \mathbf{F}. \quad (\text{A1})$$

315 Here $\mathbf{u} = (u, v, w)$ is the three-dimensional velocity vector, Π is a kinematic pressure
316 perturbation, ν is the kinematic viscosity coefficient, and \mathbf{F} is the sum of nonlinear
317 acceleration and buoyancy terms. Applying the incompressibility condition,

318
$$\nabla \cdot \mathbf{u} = 0, \quad (\text{A2})$$

319 in the equation that results from taking the divergence of (A1) (e.g., Orszag et al., 1986)
320 yields the Poisson equation,

321
$$\nabla^2 \Pi = \nabla \cdot \mathbf{F}. \quad (\text{A3a})$$

322 Although (A1) and (A2) imply (A3a), the reverse statement is not generally true.
323 Indeed, eliminating Π from between (A3a) and the equation arising from taking the
324 divergence of (A1) yields the diffusion equation $\partial \delta / \partial t = \nu \nabla^2 \delta$ for the velocity
325 divergence $\delta \equiv \nabla \cdot \mathbf{u}$, whose solution is (A2) only if δ is zero initially and on all
326 boundaries (Orszag et al., 1986; Gresho and Sani, 1987, Vreman 2014).

327 The same steps leading to (A3a) also lead to an alternative Poisson equation,

328
$$\nabla^2 \Pi = \nabla \cdot (\nu \nabla^2 \mathbf{u} + \mathbf{F}). \quad (\text{A3b})$$

329 Although $\nabla \cdot \nu \nabla^2 \mathbf{u}$ was omitted in (A3a) [this term is zero if (A2) is satisfied], without
330 further constraints on δ (described above), (A2) may not be satisfied. Gresho and Sani

331 (1987) showed that the retention of $\nabla \cdot \nu \nabla^2 \mathbf{u}$ in (A3b) assures that (A2) is satisfied,
 332 and thus leads to the paradox: "If you include it, you don't need it; if you don't include
 333 it, you need it." Vreman (2014) revisited this paradox, and showed that for a standard
 334 staggered method, the discretized form of (A3b) is equivalent to that of (A3a)
 335 supplemented with the constraint that $\nabla \cdot \nabla^2 \mathbf{u} = 0$ ($\nabla^2 \delta = 0$) on points adjacent to the
 336 solid boundary [with the same inhomogeneous Neumann boundary condition for Π
 337 implied for (A3a) and (A3b)]. When supplemented with this $\nabla^2 \delta = 0$ near-wall
 338 condition, the diffusion equation for δ led to $\delta = 0$ for all time. We note that (A3b) is
 339 the form adopted in our numerical code.

340 Evaluating the vertical component of (A1) on the surface, where the
 341 impermeability condition applies, yields the inhomogeneous Neumann condition,

$$342 \quad \left. \frac{\partial \Pi}{\partial z} \right|_0 = \nu \left. \frac{\partial^2 w}{\partial z^2} \right|_0 + F_z \Big|_0, \tag{A4}$$

343 where $w \equiv \mathbf{k} \cdot \mathbf{u}$, $F_z \equiv \mathbf{k} \cdot \mathbf{F}$, \mathbf{k} is the upward unit vector, and $(\) \Big|_0$ is a surface value. It
 344 has been argued that (A4), by itself, is not a proper boundary condition because it does
 345 not provide new information (it is not independent of the governing equations) and does
 346 not enforce the incompressibility condition (A2) at the boundary (Strikwerda, 1984;
 347 Henshaw, 1994; Sani et al., 2006). However, as pointed out by Henshaw (1994), many
 348 studies that impose (A4) (or a variant of it) also apply (A2) on the boundary.

349 In our numerical model, (A1) is integrated using a fractional step procedure with

350 explicit treatment of the viscous term. First, a provisional velocity field $\tilde{\mathbf{u}}$ that does not
351 satisfy (A2) is obtained by integrating a discretized form of (A1) in which the pressure
352 gradient is omitted. The provisional velocity is equal to the velocity at the end of the
353 previous time step plus the sum of the forcing terms (nonlinear acceleration, buoyancy,
354 and viscous stress) multiplied by the time step Δt . With the forcing terms explicitly
355 evaluated, $\tilde{\mathbf{u}}$ is readily computed throughout the flow domain, including on the surface,
356 where, in surface-forced flows, the buoyancy will make a substantial contribution. In
357 terms of $\tilde{\mathbf{u}}$ and its vertical component \tilde{w} , (A3b) and (A4) become,

$$358 \quad \nabla^2 \Pi = \frac{\nabla \cdot \tilde{\mathbf{u}}}{\Delta t}, \quad (\text{A5})$$

$$359 \quad \left. \frac{\partial \Pi}{\partial z} \right|_0 - \frac{1}{\Delta t} \tilde{w} \Big|_0 = 0. \quad (\text{A6})$$

360 In the second step, a velocity field that does satisfy (A2) is obtained by solving (A5) for
361 Π and then adding the pressure gradient force associated with Π (multiplied by Δt) to
362 $\tilde{\mathbf{u}}$.

363 In some explicit fractional step procedures (including the DNS code used in our
364 study), the problem of solving (A5) subject to (A6) with $\tilde{\mathbf{u}}|_0$ evaluated from model data
365 is replaced by what appears to be an entirely different (but is actually equivalent)
366 problem: solving (A5) subject to the homogeneous Neumann condition,

$$367 \quad \left. \frac{\partial \Pi}{\partial z} \right|_0 = 0, \quad (\text{A7})$$

368 in concert with $\tilde{\mathbf{u}}|_0$ being set to 0, obviating the need to calculate $\tilde{\mathbf{u}}|_0$ from model data.
 369 It can be shown that $\tilde{w}|_0$ and the discretized form of $\partial\Pi/\partial z|_0$ appear in the discretized
 370 form of (A5) valid half a grid point above the physical surface as $\partial\Pi/\partial z|_0 - \tilde{w}|_0/\Delta t$, that
 371 is, in the same combination as they appear in (A6). Thus, setting $\tilde{w}|_0$ and $\partial\Pi/\partial z|_0$ to 0,
 372 is equivalent to implementing (A6) with the model-computed values of $\tilde{w}|_0$: the
 373 discretized form of (A5) near the surface is the same in either case. Moreover, on the C
 374 grid, setting the tangential components $\tilde{u}|_0$ and $\tilde{v}|_0$ to 0 only affects the values of \tilde{u} and
 375 \tilde{v} half a grid point beneath the physical boundary. These values do not appear in the
 376 discretized form of (A5) at any z -level, and thus have no bearing on the solution. In
 377 essence, the errors associated with the conflation of the two physically unjustifiable
 378 specifications (homogeneous Neumann condition for pressure, and $\tilde{\mathbf{u}}|_0 = 0$) cancel out.

379 The homogeneous Neumann condition for pressure can be the source of confusion
 380 if the context in which the condition is applied is not made clear: it would be a correct
 381 condition if $\tilde{\mathbf{u}}|_0$ is set to zero (per the equivalence described above), but it would be an
 382 incorrect condition if the explicit model-computed values of $\tilde{\mathbf{u}}|_0$ are used. In the
 383 experiments with the mis-specified condition described in Sect. 3, the homogeneous
 384 condition is imposed in the latter context. Unfortunately, in many numerical model
 385 descriptions, the nature of the surface pressure condition is left vague, for example, by

386 not indicating whether a Neumann condition is homogeneous or inhomogeneous, or, if a
387 homogeneous Neumann condition is indicated, not mentioning how $\tilde{\mathbf{u}}|_0$ is treated.

388 Finally, we note that in fractional step procedures that treat the viscous term
389 implicitly (e.g., Kim and Moin, 1985; Gresho, 1990; Armfield and Street, 2002;
390 Guermond et al., 2006, and many others), the homogeneous Neumann condition is often
391 applied as a surface condition for a Poisson equation, but it is again different from our
392 implementation described in Sect. 3. In the implicit treatments, the provisional velocity
393 is obtained as the solution of a boundary value problem ($\tilde{\mathbf{u}}|_0$ should be specified; often
394 it is set to 0) in which the relevant Poisson equation resembles (A5) but applies to a
395 scalar function (sometimes called a pseudo-pressure) that is not the real pressure.
396 Temam (1991) refers to this scalar as, "... a technical quantity, a mathematical
397 auxiliary..." and advocates that it should not even be considered as an approximation of
398 the pressure. Interestingly, in the context of implicit treatments, the homogeneous
399 Neumann condition on the pseudo-pressure has sometimes been implicated as corrupting
400 solution accuracy through the development of spurious numerical boundary layers
401 adjacent to solid boundaries (Gresho, 1990; Guermond et al., 2006; Hosseini and Feng,
402 2011).

403

404 **Code availability**

405 The Fortran program used to generate output data files from the analytical solution is

406 available as a supplement to this article. That program (square.f) is configured for test
407 A-1, but can be easily adjusted to run test A-2 or other tests. Running square.f
408 automatically generates an output file for each dependent variable (e.g., u.dat) as well
409 as an output file (square.out) that summarizes the test parameters and gives the
410 computed values of the linearity ratios R_η and R_b defined in (3.1).

411

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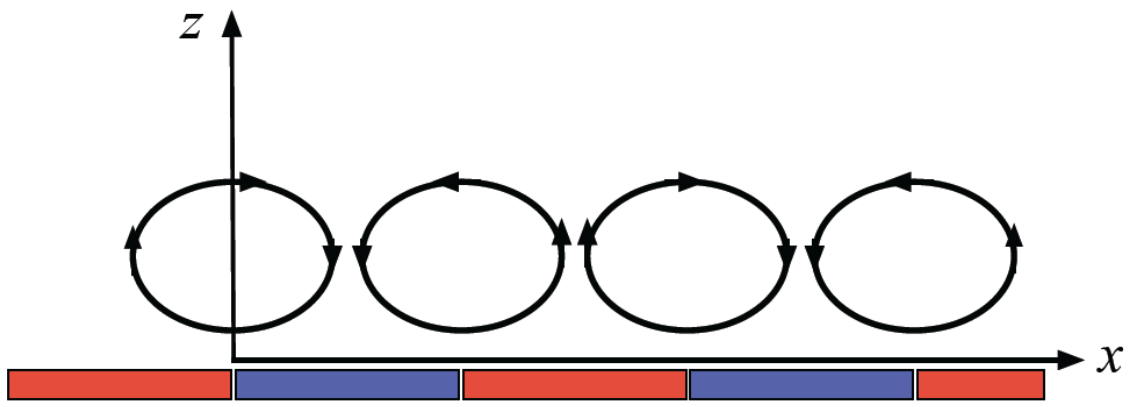
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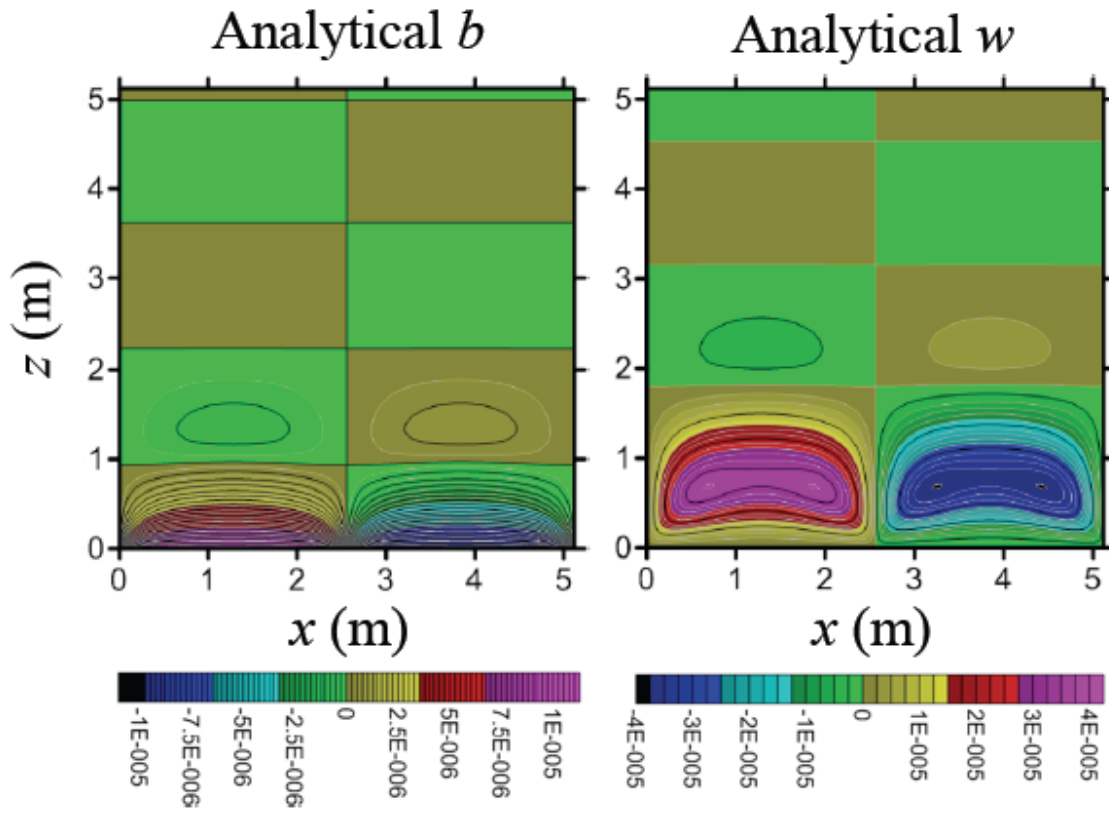
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511 **Figure 1.** Schematic of two-dimensional (x, z) thermal convection induced by a surface

512 buoyancy that varies horizontally (x) as a square wave. Red denotes positive surface

513 buoyancy, blue denotes negative surface buoyancy.

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520 **Figure 2.** Vertical cross section of the analytical (A-1) buoyancy b and vertical velocity

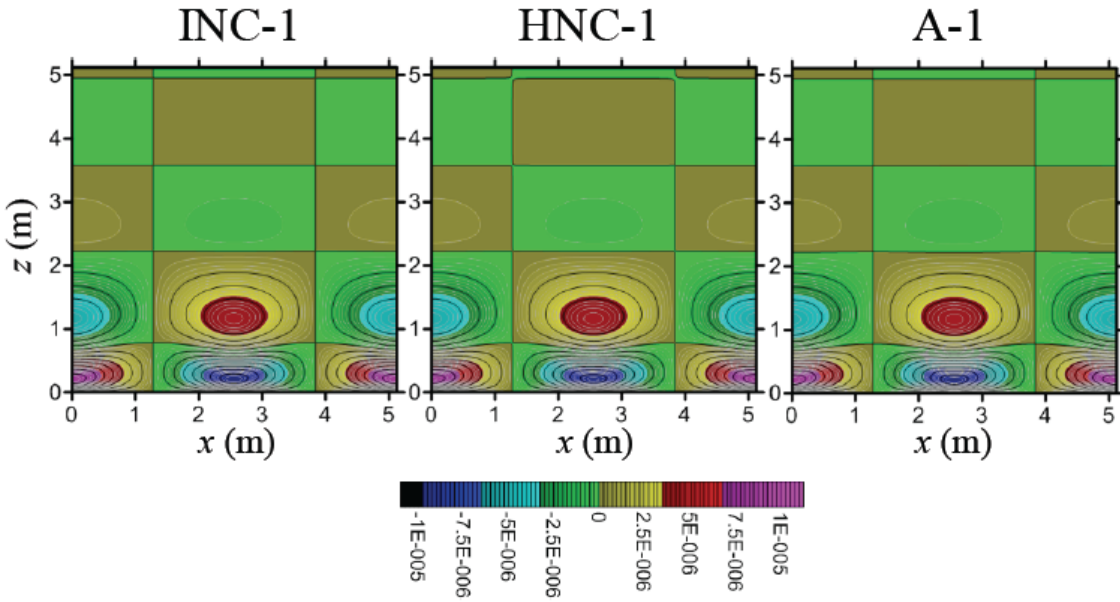
521 w fields from the first test case. Color bar units are m s^{-2} for b , and m s^{-1} for w .

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531 **Figure 3.** Vertical cross section of u from the first test case. A-1 is the analytical

532 solution. INC-1 is the numerical simulation with inhomogeneous Neumann condition for

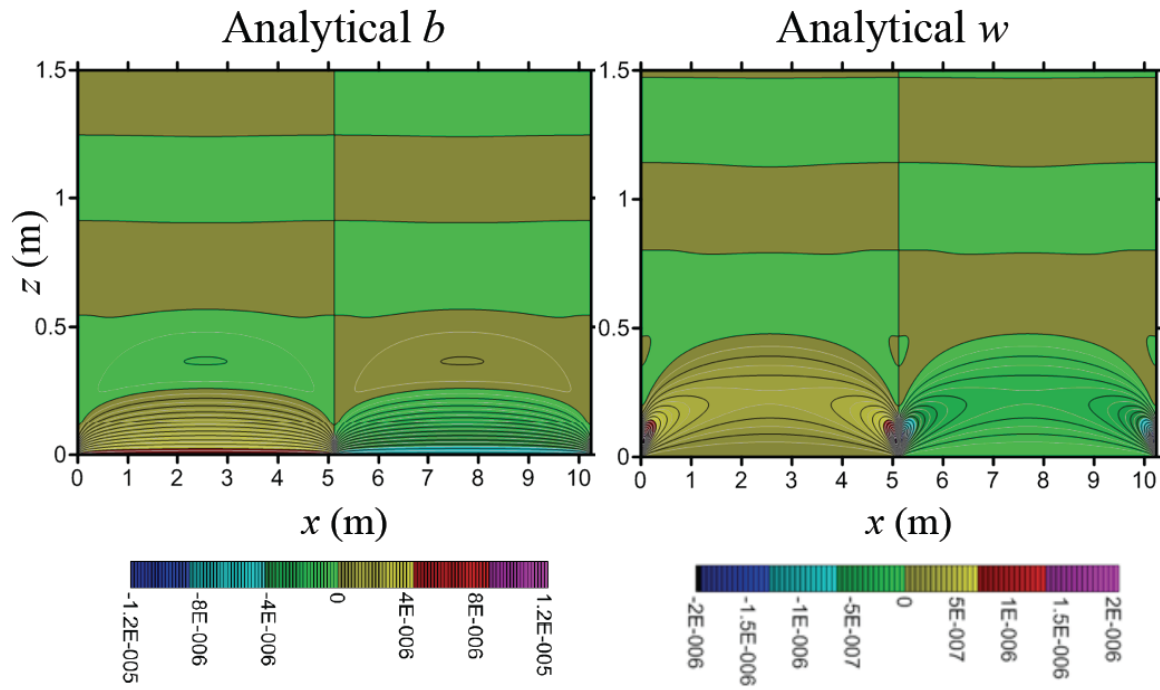
533 pressure. HNC-1 is the numerical simulation with the homogeneous Neumann condition

534 for pressure. Color bar units are m s^{-1} .

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542 **Figure 4.** Vertical cross section of the analytical (A-2) buoyancy b and vertical velocity

543 w fields from the second test case. Color bar units are m s^{-2} for b , and m s^{-1} for w .

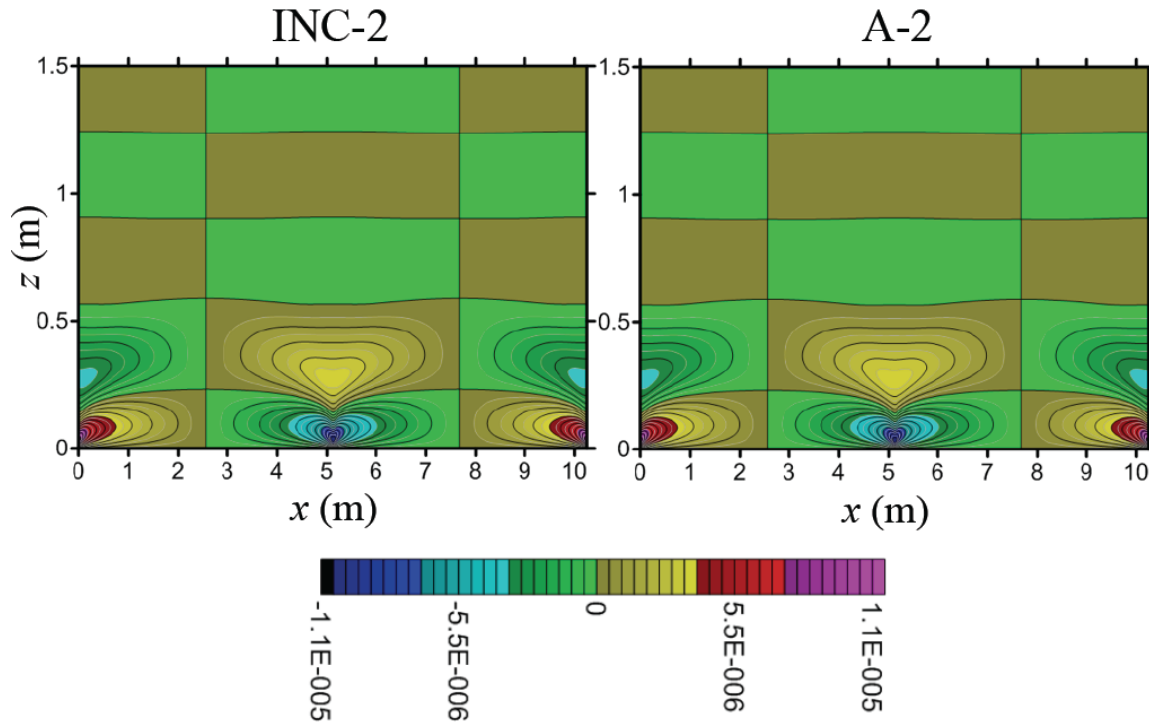
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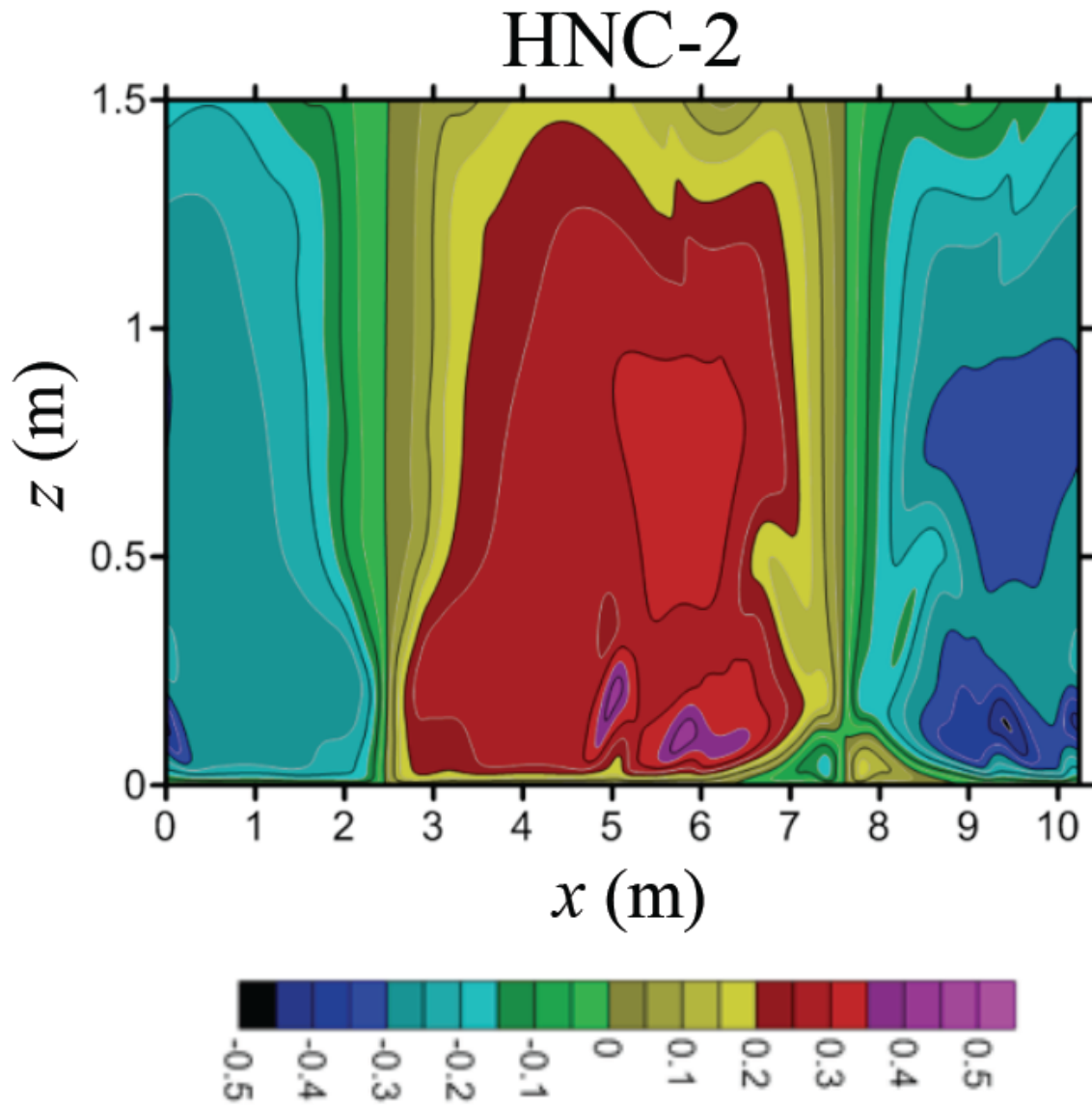
551

552 **Figure 5.** Vertical cross section of u from the second test case. A-2 is the analytical

553 solution. INC-2 is the numerical simulation with inhomogeneous Neumann condition for

554 pressure. Color bar units are m s^{-1} .

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560 **Figure 6.** Vertical cross section of u from HNC-2, the numerical simulation with

561 homogeneous Neumann condition for pressure in the second test case. Color bar units

562 are m s^{-1} .