

Interactive comment on “Mass-conserving subglacial hydrology in the Parallel Ice Sheet Model” by E. Bueller and W. Van Pelt

Anonymous Referee #3

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Summary of the manuscript

The manuscript (MS) describes a novel subglacial hydrology model implemented as part of the PISM ice sheet model. To my knowledge, this model together with the model of de Fleurian et al. (2014) (Elmer/Ice) are the currently most complex hydrology models included in large scale ice sheet models. The hydrology model consists of a cavity-like layer which can conduct the water horizontally, and two storage components: a till layer and an englacial aquifer. The coupling to ice flow would be through the yield strength of the till, which in terms depends on the amount of water stored (although no two-way coupled runs are demonstrated). The model performs well on test cases with analytic solutions and on an application to the Greenland ice sheet.

The MS is very detailed and describes the mathematical model, some analytic solu-

C2006

tions, the numerical implementation and some test applications. The MS is suitable for publication in GMD after the comments below are addressed.

General comments

Mathematical model

My main comments are that water in englacial storage is not accounted for, that the statements $0 \leq P \leq P_o$ and $W = Y$ are inconsistent, and that boundary conditions are omitted. Further, in the mathematical sections it is never explained how in detail the bounds on P and also W_{till} are enforced, although it can be deciphered from the later sections on numerical implementation. Also the authors mention that their pressure regularisation is necessary to allow enforcing $0 \leq P \leq P_o$ (by projection). Why is this so? Why could this not be done using the elliptic pressure equation?

Mass conservation (Eq. 1, 34a) should also take into account W_{eng} , the equivalent layer of water stored englacially:

$$\frac{\partial W}{\partial t} + \frac{\partial W_{till}}{\partial t} + \frac{\partial W_{eng}}{\partial t} + \nabla \cdot \mathbf{q} = m/\rho_w. \quad (1)$$

In particular, for the void ratios ($\phi_0 = 0.01$) considered in this MS the W_{eff} term is important. For instance, a relatively small pressure difference of 10 m water head leads to a change in W_{eff} of 0.1 m which is on the order of W . In fact, having $\phi_0 = 0.01$ is probably beyond what may be considered a regularisation (i.e. having negligible effect on the solution), and the MS should be updated accordingly.

If possible, it would be nice to state the bounds on the various state equations more explicitly, e.g.:

$$\frac{\partial W_{till}}{\partial t} = \begin{cases} m/\rho_w - C_d & \text{if ...} \\ 0 & \text{otherwise.} \end{cases} \quad (2)$$

Or if that is not possible, state the bounds next to the equations.

C2007

For the pressure, according to the numerics outlined in section 7.6, the authors solve Eq.33 on the whole domain for P and then project/update P such that $0 \leq P \leq P_o$ (except where $W = 0$ also $P = P_o$). Therefore $Y = P$ is only true in the so-called “normal-pressure” regions, which should be stated. In the overpressure or underpressure regions the authors instead use the mathematical closures $P = P_o$ and $P = 0$, which should also be stated. Also, it seems that the pressure equation is solved for the whole domain using boundary conditions at the edge of the domain, which is in contrast to Schoof et al. (2012). This difference needs to be discussed in a section about boundary conditions.

Even apart from the storage term (which the authors acknowledge), the presented scheme is not quite equivalent to the one in Schoof et al. (2012): To determine the regions where pressure equation needs to be solved (Eq.34c in this MS) Schoof et al. (2012) uses constraints on W and not on P (see their equations 4.1, 4.7 and 4.11). In the region where the pressure equation is solved, Schoof et al. (2012) uses appropriate boundary condition to link to the adjacent regions. Also in underpressure regions Schoof et al. (2012) solve both for Y and W (their h and h_w).

To illustrate the impact of the different models, here a pathological case which (I think) the mathematical model of Schoof et al. (2012) handles fine but the one in this MS less so:

Starting with an initial, steady state with a region where $W > W_r$ and $P = P_o$. Decrease input into that region until $P < P_o$, i.e. something like a draining subglacial lake. Now (as far as I understand the equations in the MS) W in that region would evolve according to Eq.13, i.e. shrink by viscous creep (unless $P < 0$ at which point it would again evolve according to Eq.34a). This contrasts to Schoof et al. (2012) which keeps $P = P_o$ until $W \leq W_r$.

Not getting this and other corner cases right is not bad and still results in a great subglacial hydrology model, in particular for the application intended here. However,

C2008

Schoof et al. (2012) gets them right(er) (as far as I understand) and thus the authors' claims that they successfully solve that problem should be a bit more qualified (see line-comments below).

Other comments

The manuscript is quite lengthy and could do with some streamlining. Among others, Section 4.1 and 5.2 should be merged, Section 9.2.1 should be shortened and Fig. 6 removed.

It would help if the authors would state the unknown variables at the beginning of the mathematical description.

The authors mention frozen conditions but never go into details about them. What happens to the cavity sheet and till layer when input is negative? What does the water pressure do? What do the cavities and thus W do? In fact, the evolution equation for Y does not contain a melt/freeze term so $Y > 0$ even when frozen. How does this link to setting $P = P_o$ when $W = 0$ (p.4742 l.4). This should warrant at least a paragraph.

Comments by line

Comments by page and line number (add 4700 to the page number):

p.6 l.8 State how many parameters are used

p.6 l.8 Instead of “We use englacial porosity as a regularization, and we preserve physical bounds on the pressure.” write “We use englacial porosity as a regularization to impose physical bounds on the pressure.” But in fact, I am not sure this statement is right, as bounds on the pressure are enforced by projecting it onto $0 \leq P \leq P_o$.

p.6 l.21 reword “reasonable”

p.8 l.4-6 This is not quite right, see my Section above.

C2009

- p.8.1.19-24** Maybe this paragraph should be moved to start at line 7.
- p.8.1.29** Whilst no mathematical proven of convergence of grid-based models is available, they do seem to converge under grid refinement in a statistical sense (see appendix of Werder et al. (2013)). Also, their parameters are independent of the grid. Thus automatic grid-resolution determination should be possible.
- p.9.1.8** “closures” here and elsewhere can be confused with “creep closure”, reword.
- p.9.1.19** It would help to briefly introduce which processes will be described and in particular which are the unknown variables (or major variables as the authors call them later).
- p.10 Eq.1** add a term $\partial W_{eng}/\partial t$
- p.10.1.9** it is not quite clear what “the two-dimensional subglacial layer” is. Presumably it is the layer which has thickness W .
- p.10.1.18** Specify that the pressure P is at the top of the water layer too.
- p.16.1.8** write “and $N_{til} = P_o - P_{til}$ is the effective pressure of the overlying ice on the saturated till ...”.
- p.16.1.10** Should be “previous section” but specify section number instead.
- p.16.1.19** I find N_0 confusing. The very similar looking subscript “o” in P_o refers overburden but the “0” is something else. Maybe N_r or N_{ref} ?
- p.17.1.8-16** What follows in this part is unclear. Reformulate of this introductory sentence “On the other hand we will describe the maximum capacity of the till by specifying...” to prepare the reader that instead of working with δ you change to W_{til}^{max} .

C2010

- p.17.1.10** Should this not just be $W_{til} < W_{til}^{max}$? The lower bound is never used, or is it?
- p.19.1.10** For this section the Y equation is not needed/decoupled. That should be mentioned.
- p.20.1.22** comma after “consider”.
- p.22.1.10** Expand here (or maybe elsewhere) on how $P \leq P_o$ is enforced.
- p.23.1.15-22** This paragraph is a bit misplaced in this section. Maybe the enforcement of the various constraints, including $0 \leq P \leq P_o$, warrants its own section. Which is where this paragraph would belong.
- p.23.1.19-22** These two sentences suggest that the authors have solved the “prohibitively expensive” problem of Werder et al. (2013). But as discussed above, they only solve a simplified version of Schoof et al. (2012) without channels. Reword.
- p.24 Sec.5.1** I like this summary. One suggestion: write the equations in Eq.34 all as “time derivative of unknown = something”. Add the boundary conditions.
- p.25.1.1-8** Either be specific about which functions are what type or leave the paragraph away.
- 25 Sec.5.2** This section should be merged with section 4.1, probably at this location in the MS.
- p.25.1.11-19** as stated above, I don’t think this is quite the Schoof et al. (2012) model.
- p.27.1.23** write “layer thickness” instead of “amount”
- p.28.1.1** write “layer thickness” instead of “amount” (and other places in the MS)

C2011

p.37 I.15 What happens when $W < 0$ should probably be discussed in the mathematical section too.

p.39 I.14 For a mountain glacier porosity seems to be around 0.01 (Bartholomäus et al., 2011). Porosity for an ice sheet may be more on the order of 10^{-4} .

p.40 I.13 What is the “active subglacial layer”?

p.42 I.17 is this connected to the statement on p. 37, I.15? How?

p.45 I.20 write “The spin-up grid sequence...”

45 Sec.9.2.1 This section is too long and detailed considering this is not about ice flow modelling. Is this spin-up different from others used before? Also in a similar vein, Fig. 6 could be removed.

p.48 I.5-7 The till is either completely full or empty. If I understand the dependence of sliding on the till hydrology correctly, this means either fully slippery or not at all. So, is there no dependence of sliding on hydrology? Maybe this point could be briefly discussed.

Comments for tables and figures

Tab. 3 Why is W_r so much higher here?

Fig. 2 Label R_1 and L .

Fig 2&3 they could be combined.

Fig. 6 could be left away

Fig. 8 & 11 mention what model run this is for

C2012

Fig. 11 Add a label to the colour-scale. Also, I think there is a inconsistency between the caption and the text (p.48, I.18), one says ice thickness one says sliding speed.

References

- Bartholomäus, T. C., Anderson, R. S., and Anderson, S. P. (2011). Growth and collapse of the distributed subglacial hydrologic system of Kennicott glacier, Alaska, USA, and its effects on basal motion. *Journal of Glaciology*, 57(206):985–1002.
- de Fleurian, B., Gagliardini, O., Zwinger, T., Durand, G., Le Meur, E., Mair, D., and Råback, P. (2014). A double continuum hydrological model for glacier applications. *The Cryosphere*, 8(1):137–153.
- Schoof, C., Hewitt, I. J., and Werder, M. A. (2012). Flotation and free surface flow in a model for subglacial drainage. Part I: Distributed drainage. *Journal of Fluid Mechanics*, 702:126–156.
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