

1 Re: Interactive Comment by Colin Cotter

Comment: *p1543 “This reduced communication requirement implies better overall scalability on large-scale parallel systems.” I worry about remarks like this because once they appear in the literature they tend to get cited out of context. Since a scaling factor is applied to the number of communications, the scaling itself wont be affected, but the limit to strong scaling might be. However, DG will hit the wall in strong scaling almost as soon as CG does.*

Response: Agreed. Statements such as these have been circulated in the literature for quite some time, but I’m not sure there is any observable effect until one gets to the scale of one element per processor.

Modification: This sentence has been removed.

Comment: *p1543 “I dont think that Thuburn (2008) actually says anywhere that the non-conservative form is better for conserving potential enstrophy and angular momentum. Potential enstrophy can be exactly conserved if one specifically uses the vector invariant form of the equations together with a careful treatment of the discretisation, but I dont think this would necessarily be a generic property of non-conservative formulations. Angular momentum can only be exactly preserved if the grid has rotational symmetries.*

Response: The phrasing of this sentence is understandably confusing: I mean to say that Thuburn (2008) describes angular momentum and potential enstrophy as quantities which are relevant for atmospheric motions. This form of the equations is meant to be contrasted with the conservative form, which evolves the momentum hu^d . In conservative form / flux-form it is unclear how these secondary quantities can be conserved (without choosing them as prognostic variables explicitly).

Modification: Angular momentum and potential enstrophy are particularly relevant to atmospheric motions (Thuburn, 2008) and can be easily conserved under a non-conservative formulation of the shallow-water equations (Taylor and Fournier, 2010). Conservation of these quantities is more difficult when they are diagnosed from the flux-form prognostic variables. The non-conservative formulation also has the advantage of leading to a more accurate treatment of wave-like motion when formulated appropriately (Thuburn and Woollings, 2005).

Comment: *p5148: please provide more clues as to why this discretisation produces identical results to nodal spectral element for CG spaces i.e. with direct stiffness summation.*

Response: Added.

Modification: An appendix has been added with the proof of this result.

Comment: *p5150: “A stabilization operator is necessary for finite element methods to avoid dispersive errors associated with spectral ringing.” Again, Im worried about this being cited out of context. In the compatible finite element setting we have produced finite element*

discretisations that are stabilised purely by the stable vorticity advection operator, with no need for an explicit stabilization operator. We dont need any viscosity or hyperviscosity. It is also the case that DG methods are stabilised purely from the upwinding. Please have a go at narrowing down the language here.

Response: Agreed.

Modification: Manuscript has been modified to read “Stabilization is typically needed for co-located (or unstaggered) finite element methods, whether implicitly in the form of upwinding or explicitly in the form of a diffusive operator, to avoid high-frequency dispersive errors associated with spectral ringing.”

Comment: *p5151: The viscosity operator here feels a bit like you are mixing your drinks in that the non-diffusive part doesnt rely on any test functions, but the viscosity part does. I think this just needs a bit more careful explanation to explain how you are obtaining the operator by dividing by the (diagonal) mass matrix. Is this viscosity operator actually a stable discretisation of the Laplacian i.e. does it have spurious eigenvalues? The flux reconstruction people normally have to resort to LDG/CDG-style operators for this.*

Response: I agree that it’s a bit strange to have a non-diffusive operator formulated in differential form and a diffusive operator formulated in variational form. Unfortunately, although using discrete derivatives is a very intuitive way to write the first derivatives in a finite element method (equivalence of the differential and variational form has been demonstrated above), it is not immediately clear how this formulation can be extended for second derivatives. In particular, consider a second-order spectral element method: the second derivative of any linear test function ϕ is zero within an element – but under the variational formulation the second derivative is actually correctly computed. This result extends directly to higher-order methods: A Laplacian-type diffusive operator based on the differential formulation yields spurious eigenvalues (wave modes which are not properly damped). These spurious eigenvalues do not seem to appear for a Laplacian formulated using the variational approach (this is definitely true for SE, but admittedly I have not analyzed the DG Laplacian operator described here). Note that the spurious eigenvalues do not appear under the LDG formulation since only first derivatives are used.

It seems, in this case at least, that the mixed drink actually tastes somewhat decent: The differential formulation is used for first derivatives, where it can be used to discretize the non-conservative equations, and the variational formulation is used for the diffusive terms where it appears best suited.

Modification: No modification.

Comment: *Please provide a bit more detail on how you obtained the timestep sizes for your numerical calculations, I think this is important as it is the main assessment we can make of computational cost here. It would also be good if you could make some remarks about the relative computational time for one timestep between CG and DG for the same polynomial degree (I know this is tricky since implementation details vary). Its also good to remind the reader how CG and DG DOFs scale with number of elements as a function of polynomial*

degree (unless I missed this somewhere?).

Response: For all test cases, time step sizes were chosen to be as large as possible without observing instability over the simulation period. For steady state geostrophically balanced flow (section 6.1), $dt = 2300s$, $dt = 900s$ and $dt = 500s$ were found to be unstable for continuous, g_2 discontinuous and g_1 discontinuous elements, respectively ($n_e = 4$). For zonal flow over an isolated mountain, $dt = 520s$, $dt=260s$ and $dt=130s$ were found to be unstable for continuous, g_2 discontinuous and g_1 discontinuous elements, respectively ($n_e=16$). For barotropic instability, $dt=160s$, $dt=80s$ and $dt=55s$ were found to be unstable for continuous, g_2 discontinuous and g_1 discontinuous elements, respectively ($n_e=32$).

The discontinuous code was not as thoroughly optimized as the continuous code, although based on formulation alone it is clear that the discontinuous code should be more computationally intensive. Overall the discontinuous code was observed to be about 30% slower than the continuous code for the same simulation duration and time step size on a low processor count.

Overall: Although I don't want to make an explicit judgement, it seems that the discontinuous formulation isn't worth it.

Modification: A section on "implementation considerations" has been added to the text. It reads:

On the cubed-sphere grid, the discontinuous method has $6n_e^2n_p^2$ degrees of freedom compared to $8+8(n_e(n_p-1)-1)+6(n_e(n_p-1)-1)^2$ for the continuous method. In the limit as $n_e \rightarrow \infty$ this yields a ratio of $(n_p-1)^2/n_p^2$ degrees of freedom for the continuous formulation versus the discontinuous formulation. Note that in practice, the continuous formulation typically stores redundant degrees of freedom in order to reduce computational expense associated with indexing and so memory requirements are typically identical.

The primary computational difference between the continuous and discontinuous formulations is due to the evaluation of the penalty terms (18)-(19). Note that although the robust differentiation operation (15) does require additional computation for discontinuous methods, the cost of evaluating the discontinuous terms in this expression is roughly equivalent to the computational cost of the direct stiffness summation operation needed for continuous elements. Nonetheless, from numerical experiments the discontinuous method has an approximately 30% overhead compared with a continuous method (when run with the same time step size).

Further, the following sentences were added to the Results section:

For the steady-state geostrophically balanced flow test, "Increasing the time step by 100 s led to an unstable simulation."

For the zonal flow over an isolated mountain test, "Increasing the time step by 20 s led to an unstable simulation."

For the barotropic instability test, “Increasing the time step by 10 s led to an unstable simulation.”

2 Re: Interactive Comment by Anonymous Referee

Comment: *Although the high-order models give many impressive results in the smooth cases, an effective limiter is necessary for any high-order model to correctly simulate the discontinuities even for atmospheric flows. The artificial viscosity was introduced in this study and may improve the results near the discontinuities to some extent. Is there any further plan to develop other better methods to deal with the discontinuities?*

Response: There is significant interest in pursuing the problem of limiting finite element calculations for tracer transport (see for instance Guba (2014)), but the interest is far less for dynamics calculations. High resolution calculations with the spectral element dynamical core in CESM (which runs at $n_p = 4$) have shown that topography needs to be greatly smoothed to reduce noise in the vicinity of sharp elevation changes. This effect clearly reflects the issue you have pointed out, although it is not clear what kind of effect monotization would have on the simulated climate.

Reference: Guba, Oksana, Mark Taylor, and Amik St-Cyr. “Optimization-based limiters for the spectral element method.” *Journal of Computational Physics* 267 (2014): 176-195.

Modification: No modification.

Comment: *I would like to see the numerical results of the Rossby-Haurwitz wave (Williamson's test case 6). To my knowledge, this test is sensitive to the numerical viscosity. Please check this test using different schemes with and without the artificial viscosity.*

Response: Added.

Modification: The Rossby-Haurwitz test has been added to the results section.

Comment: *FR method is a general framework for the arbitrary high-order schemes. Why do you chose $n_p=4$? Have you ever tried the higher order schemes?*

Response: A thorough investigation of different values of n_p would greatly extend the length of the manuscript, so I settled on one value of n_p , chosen in accordance with the current value of n_p for the Community Atmosphere Model spectral element dynamical core and Ullrich (2013). This medium-order choice has the advantage of greatly improving on the treatment of waves over lower-order methods, but does not suffer from the increased computational expense of much higher-order methods.

Reference: Ullrich, P.A. (2013) “Understanding the treatment of waves in atmospheric models, Part I: The shortest resolved waves of the 1D linearized shallow water equations.” *Quart.*

Modification: The following paragraph has been added to the Results section:

All simulations are performed with $n_p = 4$. A thorough investigation of different values of n_p would greatly extend the length of the manuscript, so n_p was chosen in accordance with the Community Atmosphere Model spectral element dynamical core. As argued by Ullrich (2013), this choice is also “optimal” when considering the accurate treatment of linear waves.

Comment: *Please check the balanced setup of barotropic instability test. In Chen et al., JCP, 2014, the evolution patterns of l1 errors are different on three different grids (see their Fig.23). On cubed-sphere grid, at the beginning of the simulation the l1 error becomes much larger compared with other two grids and will not decrease on the refined grids. Will you find the similar issues using your models?*

Response: I find that in this model setup there is convergence with resolution of all three error norms (at fourth-order), but this convergence does not occur until at least $n_e = 40$. Presumably there are fine-scale features in the barotropic instability test which are not resolved on the grid scale until this resolution is reached. Similar results have been observed in an analogous fourth-order cubed-sphere finite-volume code that is described in an upcoming manuscript in CAMCoS.

Modification: No modification.

3 Other Modifications

Some of the text has been clarified and compacted to address excessive wordiness and two small typos have been corrected. No modifications have been made to change the scientific results.

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A global finite-element shallow-water model supporting continuous and discontinuous elements

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Abstract

This paper presents a novel nodal finite element method for either continuous and discontinuous elements, as applied to the 2D shallow-water equations on the cubed-sphere. The cornerstone of this method is the construction of a robust derivative operator ~~which~~ that can be applied to compute discrete derivatives even over a discontinuous function space. A key advantage of the robust derivative is that it can be applied to partial differential equations in either conservative or non-conservative form. However, it is also shown that discontinuous penalization is required to recover the correct order of accuracy for discontinuous elements. Two versions with discontinuous elements are examined, using either the g_1 and g_2 flux correction function for distribution of boundary fluxes and penalty across nodal points. Scalar and vector hyperviscosity operators valid for both continuous and discontinuous elements are also derived for stabilization and removal of grid-scale noise. This method is validated using ~~three~~ four standard shallow-water test cases, including geostrophically balanced flow, a mountain-induced Rossby wave train, the Rossby-Haurwitz wave and a barotropic instability. The results show that although the discontinuous basis requires a smaller time step size than that required for continuous elements, the method exhibits better stability and accuracy properties in the absence of hyperviscosity.

1 Introduction

Modeling of the 2D shallow-water equations is an important ~~first~~ step in understanding the behavior of a numerical discretization for atmospheric ~~models. The 2D shallow-water equations capture many of the important properties of the equations of motion for the atmosphere, and so are an important first step in the development of a global atmospheric modeling system modeling.~~ In particular, the dynamical character of the global shallow-water equations is governed by features common with atmospheric motions including nonlinearity, barotropic Rossby waves and inertia-gravity waves, without the added complexity of a vertical dimension.

A comprehensive literature already exists on the development of numerical methods for the global shallow-water equations spanning the past several decades. Examples include the spec-

tral transform method (Jakob-Chien et al., 1995), semi-Lagrangian methods (Ritchie, 1988; Bates et al., 1990; Tolstykh, 2002; Zerroukat et al., 2009; Tolstykh and Shashkin, 2012; Qad-douri et al., 2012), finite-difference methods (Heikes and Randall, 1995; Ronchi et al., 1996), Godunov-type finite-volume methods (Rossmanith, 2006; Ullrich et al., 2010), staggered finite-volume methods (Lin and Rood, 1997; Ringler et al., 2008; Ringler et al., 2011), multi-moment finite-volume methods (Chen and Xiao, 2008; Li et al., 2008; Chen et al., 2013), and finite-element methods (Taylor et al., 1997; Côté and Staniforth, 1990; Thomas and Loft, 2005; Giraldo et al., 2002; Nair et al., 2005; Läuter et al., 2008; Comblen et al., 2009; Bao et al., 2013).

This paper introduces a novel ~~unified formulation for discretizing either conservative or non-conservative formulations of discrete derivative operator that is applied to~~ the shallow-water equations on a manifold using continuous and discontinuous finite elements. This work is motivated by the flux correction methods of Huynh (2007) and Vincent et al. (2011), is an alternative to formulations with discontinuous elements that ~~rely on the conservative form of the discretize the conservative~~ equations of motion with explicit momentum fluxes (Giraldo et al., 2002; Nair et al., 2005), and generalizes both spectral element and discontinuous Galerkin ~~discretizations methods~~. This approach is also quadrature-free, requiring no integral computation. ~~Further, this paper~~ This paper further introduces a general variational discretization of the scalar and vector Laplacian operator which is valid for ~~either choice of continuous or discontinuous~~ elements and only requires one communication per application of the Laplacian. ~~The discretization presented in this paper reduces to a traditional discontinuous Galerkin scheme if applied to the conservative form of the shallow-water equations.~~

~~There are several reasons why discontinuous~~ Discontinuous elements are potentially more desirable than continuous elements for several reasons: First, discontinuous elements only require parallel communication along coordinate axes, whereas continuous elements also require parallel communication along diagonals, a doubling of the total number of communications in 2D. ~~This reduced communication requirement implies better overall scalability on large-scale parallel systems.~~ Second, discontinuous elements provide a natural mechanism to enforce stabilization via discontinuous penalization (or Riemann solvers, for equations in conservation form). Third, discontinuous elements can be used in conjunction with upwind methods, which

are generally better for tracer transport and associated problems. However, discontinuous elements also have a number of disadvantages, including higher storage requirements (for the same order of accuracy), a maximum time step size which is typically smaller than that imposed for continuous elements (Ullrich, 2013), and added computational expense for many hyperbolic operations. ~~Nonetheless, it is a worthwhile to explore the differences between these two formulations for a real global modeling system.~~

The outline of this paper is as follows. Section 2 presents the shallow-water equations on a manifold. The cubed-sphere grid, which will be used for simulations on the sphere, is described in section 3. The ~~discretization~~ discretizations of the dynamics and hyperviscosity are then presented in sections 4 and 5 respectively. Results from ~~three~~ four standard shallow-water test cases are ~~shown~~ given in section 6 and conclusions ~~given~~ follow in section 7.

2 The shallow-water equations on a manifold

The 2D shallow-water equations in on a Riemannian manifold with coordinate indices $x^s = \{\alpha, \beta\}$ can be written as

$$\frac{\partial u^\alpha}{\partial t} + u^s \nabla_s u^\alpha + g^{\alpha s} \frac{\partial}{\partial x^s} (g_c H) + f(\mathbf{k} \times \mathbf{u})^\alpha = 0, \quad (1)$$

$$\frac{\partial u^\beta}{\partial t} + u^s \nabla_s u^\beta + g^{\beta s} \frac{\partial}{\partial x^s} (g_c H) + f(\mathbf{k} \times \mathbf{u})^\beta = 0, \quad (2)$$

$$\frac{\partial H}{\partial t} + \nabla_s (h u^s) = 0. \quad (3)$$

The prognostic variables are free surface height H and vector velocity $\mathbf{u} = u^\alpha \mathbf{g}_\alpha + u^\beta \mathbf{g}_\beta$, where $\mathbf{g}_\alpha = \partial \mathbf{x} / \partial \alpha$ and $\mathbf{g}_\beta = \partial \mathbf{x} / \partial \beta$ are the natural basis vectors on the manifold. ~~Two other important quantities are the~~ The fluid height h and height of the bottom topography z , which are related to the free surface height via $H = h + z$. Here g^{rs} denotes the contravariant metric with covariant inverse g_{rs} , $J = \sqrt{\det g_{rs}}$ is the metric Jacobian, g_c is gravity, f is the Coriolis parameter, and \mathbf{k} is the vertical basis vector of unit length. Einstein summation notation (im-

plied summation) is used for repeated indices. These equations further make use of the covariant derivative ∇_s , which expands as

$$u^s \nabla_s u^d = u^\alpha \frac{\partial u^d}{\partial \alpha} + u^\beta \frac{\partial u^d}{\partial \beta} + \Gamma_{sr}^d u^s u^r, \quad (4)$$

$$\nabla_s (hu^s) = \frac{1}{J} \frac{\partial}{\partial \alpha} (Jhu^\alpha) + \frac{1}{J} \frac{\partial}{\partial \beta} (Jhw^\beta), \quad (5)$$

where Γ_{sr}^d denotes the Christoffel symbols of the second kind associated with the coordinate transform (again with summation over repeated indices s and r implied).

~~Observe that (The mass equation (3) has been kept in conservative form to enforce strict mass conservation. On the other hand, (1)-(2) are given in a non-conservative form; this formulation is selected over the conservative formulation flux-form equations (where hu^α and hw^β are prognostic variables) since it can more readily conserve quantities more relevant to atmospheric motion, such as angular. Angular momentum and potential enstrophy (Thuburn, 2008), and (depending on the discretization) can lead are particularly relevant to atmospheric motions (Thuburn, 2008) and can be easily conserved under a non-conservative formulation of the shallow-water equations (Taylor and Fournier, 2010). Conservation of these quantities is more difficult when they are diagnosed from the flux-form prognostic variables. The non-conservative formulation also has the advantage of leading to a more accurate treatment of wave-like motions (Thuburn and Woollong, 2005). The mass equation (3) has been kept in conservative form to enforce strict mass conservation. motion when formulated appropriately (Thuburn and Woollong, 2005).~~

3 The cubed-sphere grid

The equations (1)-(3) are now applied to a particular choice of coordinate system. The cubed-sphere grid (Sadourny, 1972; Ronchi et al., 1996) consists of a cube with six Cartesian patches arranged along each face, which is then inflated onto a tangent spherical shell, as shown in Fig. 1. The cubed-sphere is a *quasi-uniform spherical grid*, that is, it is in the class of grids that provide an approximately uniform tiling of the sphere (see Staniforth and Thuburn (2012),

for example, for a review of different options for global grids). On the equiangular cubed-sphere grid, coordinates are given as (α, β, p) , with central angles $\alpha, \beta \in [-\frac{\pi}{4}, \frac{\pi}{4}]$ and panel index $p \in \{1, 2, 3, 4, 5, 6\}$. By convention, we choose panels 1–4 to be along the equator and panels 5 and 6 to be centered on the northern and southern pole, respectively. With uniform grid spacing, each panel consists of a square array of $n_e \times n_e$ elements.

The contravariant 2D metric on the equiangular cubed-sphere of radius a is given by

$$g^{rs} = \frac{\delta^2}{a^2(1 + \tan^2 \alpha)(1 + \tan^2 \beta)} \begin{pmatrix} 1 + \tan^2 \beta \tan \alpha \tan \beta \\ \tan \alpha \tan \beta & 1 + \tan^2 \alpha \end{pmatrix}, \quad (6)$$

where $\delta = \sqrt{1 + \tan^2 \alpha + \tan^2 \beta}$. The Jacobian on the manifold, denoted by J , is then

$$J = \sqrt{\det(g_{rs})} = \frac{a^2(1 + \tan^2 \alpha)(1 + \tan^2 \beta)}{\delta^3}, \quad (7)$$

and induces the infinitesimal area element $dA = J d\alpha d\beta$. The Christoffel symbols of the second kind are given by

$$\Gamma_{ij}^\alpha = \begin{pmatrix} \frac{2 \tan \alpha \tan^2 \beta}{\delta^2} & \frac{-\tan \beta (1 + \tan^2 \beta)}{\delta^2} \\ \frac{-\tan \beta (1 + \tan^2 \beta)}{\delta^2} & 0 \end{pmatrix}, \quad (8)$$

$$\Gamma_{ij}^\beta = \begin{pmatrix} 0 & \frac{-\tan \alpha (1 + \tan^2 \alpha)}{\delta^2} \\ \frac{-\tan \alpha (1 + \tan^2 \alpha)}{\delta^2} & \frac{2 \tan^2 \alpha \tan \beta}{\delta^2} \end{pmatrix}. \quad (9)$$

Spherical coordinates (λ, ϕ) for longitude $\lambda \in [0, 2\pi]$ and latitude $\phi \in [-\pi/2, \pi/2]$ will also be used for plotting and specification of tests. Coordinate transforms between spherical and equiangular coordinates can be found in Ullrich and Jablonowski (2012) Appendix A.

4 Nodal Finite Element Discretization

4.1 The Nodal Basis

A nodal finite element method is employed (Taylor et al., 1997; Giraldo et al., 2002; Hesthaven and Warburton, 2007). The 1D reference element is defined as the interval $x \in [-1, 1]$ along with a set of test functions $\hat{\phi}_{(i)}(x)$. The test functions are defined such that test function $\hat{\phi}_{(i)}(x)$ is the unique polynomial of degree n_p that is 1 at the i^{th} Gauss-Lobatto-Legendre (GLL) node ($i \in (0, \dots, n_p - 1)$) and 0 at all other GLL nodes. Each basis polynomial then has a corresponding weight, defined by

$$w_i = \int_{-1}^1 \hat{\phi}_{(i)}(x) dx. \quad (10)$$

The 2D element $\mathcal{Z} = [\alpha_1, \alpha_2] \times [\beta_1, \beta_2]$ (with boundary $\partial\mathcal{Z}$) has accompanying 1D basis functions

$$\tilde{\phi}_{(i)}(\alpha) = \hat{\phi}_{(i)}\left(\frac{2(\alpha - \alpha_1)}{\Delta\alpha} - 1\right), \quad \tilde{\phi}_{(j)}(\beta) = \hat{\phi}_{(j)}\left(\frac{2(\beta - \beta_1)}{\Delta\beta} - 1\right), \quad (11)$$

where $\Delta\alpha = \alpha_2 - \alpha_1$ and $\Delta\beta = \beta_2 - \beta_1$. The accompanying 2D tensor-product basis is then defined by

$$\phi_{(i,j)}(\alpha, \beta) = \tilde{\phi}_{(i)}(\alpha)\tilde{\phi}_{(j)}(\beta). \quad (12)$$

Figure 2 provides a depiction of GLL nodes within a single element. For vector quantities (such as velocity \mathbf{u}), test functions are instead vector fields. Uniqueness of the variational system is retained if exactly two degrees of freedom are allowed at each nodal location for the vector test function ϕ . As we shall see, the most natural choice is test functions $\phi_{(i,j)}^{(\alpha)}$ and $\phi_{(i,j)}^{(\beta)}$ with covariant components

$$\phi_{(i,j)\alpha}^{(\alpha)} = \phi_{(i,j)}, \quad \phi_{(i,j)\beta}^{(\alpha)} = 0, \quad \phi_{(i,j)\alpha}^{(\beta)} = 0, \quad \phi_{(i,j)\beta}^{(\beta)} = \phi_{(i,j)}. \quad (13)$$

4.2 Robust Differentiation

A robust differentiation operator is now constructed for both continuous and discontinuous finite elements. Let $f : (\alpha, \beta) \rightarrow \mathbb{R}$ be defined and continuous on $\mathcal{Z} \cup \partial\mathcal{Z}$ with basis $\phi_{(i,j)}$,

$$f(\alpha, \beta) = \sum_{p=0}^{n_p-1} \sum_{q=0}^{n_p-1} f_{(p,q)} \phi_{(p,q)}(\alpha, \beta), \quad (14)$$

with for coefficients $f_{(p,q)} \in \mathbb{R}$. Further, let $\tilde{f} : (\alpha, \beta) \rightarrow \mathbb{R}$ be defined and continuous on $\partial\mathcal{Z}$. Here \tilde{f} represents the evaluation of f in neighboring elements. Note that for a continuous finite element method, f and \tilde{f} must satisfy $\tilde{f}(\alpha, \beta) = f(\alpha, \beta)$ on $\partial\mathcal{Z}$, whereas no such restriction is imposed for discontinuous finite elements. Following Huynh (2007), robust differentiation in the α direction is defined at GLL nodes via

$$D_\alpha f(\alpha_i, \beta_j) = \sum_{p=0}^{n_p-1} f_{(p,j)} \frac{\partial \tilde{\phi}_{(p)}}{\partial \alpha}(\alpha_i) + \frac{dg_R}{d\alpha}(\alpha_i) (\overline{f}_{(n_p-1,j)} - f_{(n_p-1,j)}) + \frac{dg_L}{d\alpha}(\alpha_i) (\overline{f}_{(0,j)} - f_{(0,j)}), \quad (15)$$

where the overline denotes the co-located average of f and \tilde{f} ,

$$\overline{f}_{(n_p-1,j)} = \frac{f(\alpha_{n_p-1}, \beta_j) + \tilde{f}(\alpha_{n_p-1}, \beta_j)}{2}, \quad \overline{f}_{(0,j)} = \frac{f(\alpha_0, \beta_j) + \tilde{f}(\alpha_0, \beta_j)}{2}. \quad (16)$$

An analogous definition holds in the β direction. Here g_L and g_R are the local *flux correction functions*, which are chosen to satisfy

$$g_L(\alpha_0) = 1, \quad g_L(\alpha_{n_p-1}) = 0, \quad g_R(\alpha_0) = 0, \quad g_R(\alpha_{n_p-1}) = 1, \quad (17)$$

and otherwise are chosen to approximate zero throughout $[\alpha_0, \alpha_{n_p-1}]$. A number of Several options for g_L and g_R exist will lead to a stable discretization, including g_1 (Radau polynomials),

which will lead to the discontinuous Galerkin method, and g_2 , which will lead to the mass-lumped discontinuous Galerkin method (Huynh, 2007). Hereafter discontinuous elements with the g_1 flux correction function will be referred to as “discontinuous g_1 elements,” whereas elements using of the g_2 flux correction function will be referred to as “discontinuous g_2 elements.”

5 ~~An analogous procedure is used to construct a derivative operator in the β direction.~~ Observe that for continuous finite elements, the rightmost two terms in (15) are exactly zero.

With the definition of a robust discrete derivative (15), discretization of the shallow-water system (1)-(3) is straightforward. Note that for continuous finite elements, this discretization is identical to the approach of Taylor et al. (1997) when applied in conjunction with *Direct Stiffness Summation* (that is, projection into the space of continuous functions) (see Appendix A). If the conservative form of the shallow-water equations were employed, this discretization ~~would be is~~ the same as Giraldo et al. (2002) when mass lumping is not employed (discontinuous g_1) and Nair et al. (2005) if mass lumping is applied (discontinuous g_2). To the best of the author’s knowledge, no previous work has used both discontinuous elements and a non-conservative
15 form of the shallow-water system.

4.3 Discontinuous Penalization

At element boundaries, the use of one-sided derivatives will cause the discontinuity between neighboring elements to exhibit an error with magnitude $\mathcal{O}(\Delta x^{n_p-1})$, an effective loss of one order of accuracy from the expected convergence rate. To reduce errors associated with the
20 discontinuity, a penalization term is added in each coordinate direction. In the α direction this

term reads

$$\begin{aligned} \frac{\partial H}{\partial t}(\alpha_i, \beta_j) = & \dots + \frac{\partial g_R}{\partial \alpha}(\alpha_i) \frac{|\lambda(\alpha_{n_p-1}, \beta_j)|}{2} \left[\tilde{H}(\alpha_{n_p-1}, \beta_j) - H(\alpha_{n_p-1}, \beta_j) \right] \frac{J(\alpha_{n_p-1}, \beta_j)}{J(\alpha_i, \beta_j)} \\ & + \frac{\partial g_L}{\partial \alpha}(\alpha_i) \frac{|\lambda(\alpha_0, \beta_j)|}{2} \left[H(\alpha_0, \beta_j) - \tilde{H}(\alpha_0, \beta_j) \right] \frac{J(\alpha_0, \beta_j)}{J(\alpha_i, \beta_j)}, \end{aligned} \quad (18)$$

$$\begin{aligned} \frac{\partial u^d}{\partial t}(\alpha_i, \beta_j) = & \dots + \frac{\partial g_R}{\partial \alpha}(\alpha_i) \frac{|\lambda(\alpha_{n_p-1}, \beta_j)|}{2} \left[\tilde{u}^d(\alpha_{n_p-1}, \beta_j) - u^d(\alpha_{n_p-1}, \beta_j) \right] \\ & + \frac{\partial g_L}{\partial \alpha}(\alpha_i) \frac{|\lambda(\alpha_0, \beta_j)|}{2} \left[u^d(\alpha_0, \beta_j) - \tilde{u}^d(\alpha_0, \beta_j) \right]. \end{aligned} \quad (19)$$

where $\lambda(\alpha, \beta) = |u^\alpha| + \sqrt{gh}/a$ represents the maximum local wave speed in the α direction. An analogous term is added in the β direction. Note that with this choice of penalization the evolution equation for \tilde{H} is identical to the evolution equation that would arise from a traditional conservative discontinuous Galerkin method with local Lax-Friedrichs flux. Since the penalization term is equivalent to upwinding, it is weakly diffusive and so allows the discontinuous scheme to maintain stability even in the absence of explicit viscosity.

4.4 Implementation considerations

On the cubed-sphere grid, the discontinuous method has $6n_e^2 n_p^2$ degrees of freedom compared to $8 + 8(n_e(n_p - 1) - 1) + 6(n_e(n_p - 1) - 1)^2$ for the continuous method. In the limit as $n_e \rightarrow \infty$ this yields a ratio of $(n_p - 1)^2/n_p^2$ degrees of freedom for the continuous formulation versus the discontinuous formulation. Note that in practice, the continuous formulation typically stores redundant degrees of freedom in order to reduce computational expense associated with indexing and so memory requirements are typically identical.

The primary computational difference between the continuous and discontinuous formulations is due to the evaluation of the penalty terms (18)-(19). Note that although the robust differentiation operation (15) does require additional computation for discontinuous methods, the cost of evaluating the discontinuous terms in this expression is roughly equivalent to the computational

cost of the direct stiffness summation operation needed for continuous elements. Nonetheless, from numerical experiments the discontinuous method has an approximately 30% overhead compared with a continuous method (when run with the same time step size).

5 Viscosity and Hyperviscosity

5 ~~A stabilization operator is necessary for~~ Stabilization is typically needed for co-located (or unstaggered) finite element methods ~~to avoid~~, whether implicitly in the form of upwinding or explicitly in the form of a diffusive operator, to avoid high-frequency dispersive errors associated with spectral ringing. In general, it is preferred that this operator is consistent with the underlying geometry of the Riemannian manifold, which precludes, for example, the Boyd-Vandeven filter (Boyd, 1996). There has been considerable success with the use of hyperviscosity in the spectral element method (Dennis et al., 2011), which maintains geometric consistency by mimicking the natural fourth-order hyperviscosity operator. Previously, it has not been clear how to extend this operator to a discontinuous function space. However, the robust derivative (15) provides a direct mechanism by which the hyperviscosity operator can be constructed. The viscosity operator for both the continuous and discontinuous function space will be discussed here.

Note that any viscosity operator will lead to a loss of energy conservation of the underlying numerical method. This loss is exhibited in two obvious ways: First, for geostrophically balanced flows the error will tend to grow over time. Second, energy conservation is lost leading to a decay in the total energy content of the system over time.

5.1 Scalar Viscosity

For stabilization of the method, diffusion is added in the form of either viscosity or hyperviscosity, which corresponds to multiple applications of the viscosity operator. A scalar viscosity operator is constructed to satisfy

$$25 \quad \mathcal{H}(\nu)\psi \approx \nu \nabla^2 \psi, \tag{20}$$

where $\nabla^2 = \nabla \cdot \nabla$ is the usual scalar Laplacian. The operator is defined implicitly via a variational construction. If $f = \mathcal{H}(\nu)\psi$ then, multiplying (20) by a test function and applying integration by parts, one obtains

$$5 \quad \iint f \phi_{(i,j)} dA = \nu \left[\oint_{\partial Z} \phi_{(i,j)} \nabla \psi \cdot d\mathbf{S} - \iint_Z \nabla \phi_{(i,j)} \cdot \nabla \psi dA \right], \quad (21)$$

where $d\mathbf{S}$ is the infinitesimal line element along the boundary of Z and dA is the infinitesimal area element. The two terms on the right-hand side of this expression correspond to the viscosity flux through element boundaries and the Laplacian within the element. Under a continuous element formulation, only the rightmost term is retained and fluxes are instead accounted for via *Direct Stiffness Summation* (**DSS**) [direct stiffness summation](#). Under a discontinuous formulation, both terms are retained and discretized. The discrete equation satisfied by $f_{(i,j)}$ that follows from (21) is written as

$$f_{(i,j)} = f_{(i,j)}^B + f_{(i,j)}^A, \quad (22)$$

15 where $f_{(i,j)}^B$ denotes the discretization of the boundary integral and $f_{(i,j)}^A$ denotes the discretization of the area integral. After a lengthy derivation (see Appendix [B](#)), these discretizations read

$$f_{(i,j)}^A = -\frac{\nu}{w_i J(\alpha_i, \beta_j)} \sum_{m=0}^{n_p-1} \frac{\partial \tilde{\phi}^{(i)}}{\partial \alpha} \nabla^\alpha \psi J w_m \Big|_{\alpha=\alpha_m, \beta=\beta_j} \\ - \frac{\nu}{w_j J(\alpha_i, \beta_j)} \sum_{n=0}^{n_p-1} \frac{\partial \tilde{\phi}^{(j)}}{\partial \beta} \nabla^\beta \psi J w_n \Big|_{\alpha=\alpha_i, \beta=\beta_n}, \quad (23)$$

and

$$f_{(i,j)}^B = \nu \left[\underbrace{\frac{\delta_{i,n_p-1}}{w_i \Delta \alpha} \nabla^\alpha \psi}_{\text{Right}} + \underbrace{\frac{\delta_{j,n_p-1}}{w_j \Delta \beta} \nabla^\beta \psi}_{\text{Top}} - \underbrace{\frac{\delta_{i,0}}{w_i \Delta \alpha} \nabla^\alpha \psi}_{\text{Left}} - \underbrace{\frac{\delta_{j,0}}{w_j \Delta \beta} \nabla^\beta \psi}_{\text{Bottom}} \right], \quad (24)$$

where $\delta_{i,j}$ is the Kröner delta. Here the contravariant derivative of ψ has been used, defined via

$$\nabla^p \psi = g^{pq} \nabla_q \psi = g^{p\alpha} \frac{\partial \psi}{\partial \alpha} + g^{p\beta} \frac{\partial \psi}{\partial \beta}. \quad (25)$$

Note that the contravariant derivatives $\nabla^p \psi$ are multivalued along this edge, and so must be adjusted using the robust derivative operator (15).

5.2 Vector Viscosity

Vector viscosity is used for damping of the velocity field, and takes the form

$$\mathcal{H}(\nu_d, \nu_v) \mathbf{u} \approx \nu_d \nabla (\nabla \cdot \mathbf{u}) - \nu_v \nabla \times (\nabla \times \mathbf{u}). \quad (26)$$

Observe that if $\nu = \nu_d = \nu_v$ then this expression is exactly the standard vector Laplacian operator $\nabla^2 \mathbf{u}$, with coefficient ν . By writing the vector Laplacian as (26), the combined operator separates into two distinct operators that effect divergence damping (with coefficient ν_d) and vorticity damping (with coefficient ν_v). This result can be quickly verified by taking the divergence and curl of (26),

$$\nabla \cdot \mathcal{H}(\nu_d, \nu_v) \mathbf{u} = \nu_d \nabla^2 (\nabla \cdot \mathbf{u}), \quad (27)$$

$$\nabla \times \mathcal{H}(\nu_d, \nu_v) \mathbf{u} = -\nu_v \nabla \times (\nabla \times (\nabla \times \mathbf{u})) = \nu_v \nabla^2 (\nabla \times \mathbf{u}) \quad (28)$$

For simplicity of calculation, we treat divergence damping and vorticity damping separately. For divergence damping, the operator is constructed by taking the inner product of $\mathbf{f} = \mathcal{H}(\nu_d, \nu_v) \mathbf{u}$ with the vector test function ϕ , integrating over \mathcal{Z} and applying integration by parts,

$$\begin{aligned} \nu_d \iint_{\mathcal{Z}} \phi \cdot \mathbf{f} dA &= \nu_d \iint_{\mathcal{Z}} \phi \cdot \nabla (\nabla \cdot \mathbf{u}), \\ &= \nu_d \left[\oint_{\partial \mathcal{Z}} (\nabla \cdot \mathbf{u}) \phi \cdot d\mathbf{S} - \iint_{\mathcal{Z}} (\nabla \cdot \phi) (\nabla \cdot \mathbf{u}) dV \right]. \end{aligned} \quad (29)$$

For vorticity damping an analogous procedure leads to

$$\begin{aligned}
 \nu_v \iint_{\mathcal{Z}} \boldsymbol{\phi} \cdot \mathbf{f} dA &= -\nu_v \iint_{\mathcal{Z}} \boldsymbol{\phi} \cdot \nabla \times (\nabla \times \mathbf{u}) dV, \\
 &= -\nu_v \left[\oint_{\partial\mathcal{Z}} (\nabla \times \mathbf{u}) \times \boldsymbol{\phi} \cdot d\mathbf{S} + \iint_{\mathcal{Z}} (\nabla \times \boldsymbol{\phi}) \cdot (\nabla \times \mathbf{u}) dV \right]
 \end{aligned} \tag{30}$$

Note that for shallow-water flows, only the radial component of the vorticity is relevant for this calculation. The discrete value of $f_{(i,j)}^\alpha$ and $f_{(i,j)}^\beta$ that arises from this calculation then has contributions from the area integral via $f_{(i,j)}^{A,d}$ and boundary integral via $f_{(i,j)}^{B,d}$,

$$f_{(i,j)}^\alpha = f_{(i,j)}^{B,\alpha} + f_{(i,j)}^{A,\alpha}, \quad f_{(i,j)}^\beta = f_{(i,j)}^{B,\beta} + f_{(i,j)}^{A,\beta}. \tag{31}$$

Following another lengthy derivation (see Appendix [B](#)) the area integral term appears as

$$\begin{aligned}
 f_{(i,j)}^{A,\alpha} &= -\frac{\nu_d}{J(\alpha_i, \beta_j) w_i} \sum_{m=0}^{n_p-1} J g^{\alpha\alpha} \frac{d\tilde{\phi}_{(i)}}{d\alpha} (\nabla \cdot \mathbf{u}) w_m \Big|_{\alpha=\alpha_m, \beta=\beta_j} \\
 &\quad - \frac{\nu_d}{J(\alpha_i, \beta_j) w_j} \sum_{n=0}^{n_p-1} J g^{\beta\alpha} \frac{d\tilde{\phi}_{(j)}}{d\beta} (\nabla \cdot \mathbf{u}) w_n \Big|_{\alpha=\alpha_i, \beta=\beta_n} \\
 &\quad + \frac{\nu_v}{J(\alpha_i, \beta_j) w_j} \sum_{n=0}^{n_p-1} \frac{d\tilde{\phi}_{(j)}}{d\beta} (\nabla \times \mathbf{u})_r w_n \Big|_{\alpha=\alpha_i, \beta=\beta_n},
 \end{aligned} \tag{32}$$

and

$$\begin{aligned}
 f_{(i,j)}^{A,\beta} = & -\frac{\nu_d}{J(\alpha_i, \beta_j)w_i} \sum_{m=0}^{n_p-1} Jg^{\alpha\beta} \frac{d\tilde{\phi}^{(i)}}{d\alpha} (\nabla \cdot \mathbf{u}) w_m \Big|_{\alpha=\alpha_m, \beta=\beta_j} \\
 & -\frac{\nu_d}{J(\alpha_i, \beta_j)w_j} \sum_{n=0}^{n_p-1} Jg^{\beta\beta} \frac{d\tilde{\phi}^{(j)}}{d\beta} (\nabla \cdot \mathbf{u}) w_n \Big|_{\alpha=\alpha_i, \beta=\beta_n} \\
 & -\frac{\nu_v}{J(\alpha_i, \beta_j)w_i} \sum_{m=0}^{n_p-1} \frac{d\tilde{\phi}^{(i)}}{d\alpha} (\nabla \times \mathbf{u})_r w_m \Big|_{\alpha=\alpha_m, \beta=\beta_j}, \tag{33}
 \end{aligned}$$

whereas the boundary integral term is

$$\begin{aligned}
 f_{(i,j)}^{B,\alpha} = & \nu_d \left[\underbrace{\frac{\delta_{i,n_p-1} g^{\alpha\alpha} (\nabla \cdot \mathbf{u})}{w_i \Delta\alpha}}_{\text{Right}} + \underbrace{\frac{\delta_{j,n_p-1} g^{\alpha\beta} (\nabla \cdot \mathbf{u})}{w_j \Delta\beta}}_{\text{Top}} - \underbrace{\frac{\delta_{i,0} g^{\alpha\alpha} (\nabla \cdot \mathbf{u})}{w_i \Delta\alpha}}_{\text{Left}} - \underbrace{\frac{\delta_{j,0} g^{\alpha\beta} (\nabla \cdot \mathbf{u})}{w_j \Delta\beta}}_{\text{Bottom}} \right]_{\alpha=\alpha_i, \beta=\beta_j} \\
 & + \nu_v \left[\underbrace{-\frac{\delta_{j,n_p-1} (\nabla \times \mathbf{u})_r}{Jw_j \Delta\beta}}_{\text{Top}} + \underbrace{\frac{\delta_{j,0} (\nabla \times \mathbf{u})_r}{Jw_j \Delta\beta}}_{\text{Bottom}} \right]_{\alpha=\alpha_i, \beta=\beta_j}. \tag{34}
 \end{aligned}$$

Applying an analogous procedure for test function $\phi_{(i,j)}^{(\beta)}$,

$$\begin{aligned}
 f_{(i,j)}^{B,\beta} = & \nu_d \left[\underbrace{\frac{\delta_{i,n_p-1} g^{\beta\alpha} (\nabla \cdot \mathbf{u})}{w_i \Delta \alpha}}_{\text{Right}} + \underbrace{\frac{\delta_{j,n_p-1} g^{\beta\beta} (\nabla \cdot \mathbf{u})}{w_j \Delta \beta}}_{\text{Top}} - \underbrace{\frac{\delta_{i,0} g^{\beta\alpha} (\nabla \cdot \mathbf{u})}{w_i \Delta \alpha}}_{\text{Left}} - \underbrace{\frac{\delta_{j,0} g^{\beta\beta} (\nabla \cdot \mathbf{u})}{w_j \Delta \beta}}_{\text{Bottom}} \right]_{\alpha=\alpha_i, \beta=\beta_j} \\
 & + \nu_v \left[\underbrace{\frac{\delta_{i,n_p-1} (\nabla \times \mathbf{u})_r}{J w_i \Delta \alpha}}_{\text{Right}} - \underbrace{\frac{\delta_{i,0} (\nabla \times \mathbf{u})_r}{J w_i \Delta \alpha}}_{\text{Left}} \right]_{\alpha=\alpha_i, \beta=\beta_j} . \tag{35}
 \end{aligned}$$

The divergence and curl, which are needed for evaluation of the Laplacian, are computed via

$$\nabla \cdot \mathbf{u} = \nabla_p u^p = \nabla_\alpha u^\alpha + \nabla_\beta u^\beta \tag{36}$$

$$(\nabla \times \mathbf{u})_r = \epsilon_{r pq} g^{ps} \nabla_s u^q = J \left[g^{\alpha\alpha} \nabla_\alpha u^\beta + g^{\alpha\beta} \nabla_\beta u^\beta - g^{\beta\alpha} \nabla_\alpha u^\alpha - g^{\beta\beta} \nabla_\beta u^\alpha \right], \tag{37}$$

where

$$\nabla_\alpha u^\alpha = \frac{\partial u^\alpha}{\partial \alpha} + \Gamma_{\alpha\alpha}^\alpha u^\alpha + \Gamma_{\alpha\beta}^\alpha u^\beta, \quad \nabla_\alpha u^\beta = \frac{\partial u^\beta}{\partial \alpha} + \Gamma_{\alpha\alpha}^\beta u^\alpha + \Gamma_{\alpha\beta}^\beta u^\beta, \tag{38}$$

$$\nabla_\beta u^\alpha = \frac{\partial u^\alpha}{\partial \beta} + \Gamma_{\beta\alpha}^\alpha u^\alpha + \Gamma_{\beta\beta}^\alpha u^\beta, \quad \nabla_\beta u^\beta = \frac{\partial u^\beta}{\partial \beta} + \Gamma_{\beta\alpha}^\beta u^\alpha + \Gamma_{\beta\beta}^\beta u^\beta. \tag{39}$$

All partial derivatives are evaluated using the robust derivative operator (15).

5.3 Hyperviscosity

For stabilization of a high-order discretization, hyperviscosity is preferred since it retains the order of accuracy of the underlying scheme. In practice, hyperviscosity is implemented by repeated application of the viscosity operator. For instance, for fourth-order hyperviscosity, the

∇^4 operator is approximated as follows

$$\frac{\partial \mathbf{u}}{\partial t} = \mathcal{H}(\nu_d, \nu_v) \mathcal{H}(1, 1) \mathbf{u}, \quad \frac{\partial h}{\partial t} = \mathcal{H}(\nu) \mathcal{H}(1) h. \quad (40)$$

5.4 Computational Considerations

Calculation of hyperviscosity in the form presented here requires one parallel exchange per application of the Laplacian operator. For continuous elements, this communication is manifested through the application of DSS, which averages away any discontinuity that has been generated along element edges. For discontinuous elements, scalar viscosity requires pointwise updates along element edges computed from (24), whereas vector viscosity requires both one-sided values of \mathbf{u} , $(\nabla \cdot \mathbf{u})$ and $(\nabla \times \mathbf{u})_r$, which are in turn used for computing nodal values of $(\nabla \cdot \mathbf{u})$ and $(\nabla \times \mathbf{u})_r$ needed for (32)-(35). This constitutes a doubling of the overall bandwidth requirement relative to continuous elements.

6 Results

In this section selected results are provided to evaluate the relative performance of the [method methods](#) described in this paper. [Three](#) [Four](#) test cases are evaluated: From the Williamson et al. (1992) test case suite, steady-state geostrophically balanced flow [and](#) zonal flow over an isolated mountain [and the Rossby-Haurwitz wave](#) will be analyzed, in addition to the barotropic instability test of Galewsky et al. (2004). For all test cases, time integration is handled via the strong-stability preserving three-stage third-order Runge-Kutta method (Gottlieb et al., 2001). Note that some improvement in the maximum time step size for discontinuous elements can be obtained from the use of the five-stage third-order Runge-Kutta method (Ruuth, 2006), which has a stability region that covers a larger portion of the negative real plane. The time step Δt for each test is chosen to be close to the stability limit in each case (observed empirically). The value of Δt has negligible effect on the results (not shown), suggesting that spatial errors dominate in each case. Further, note that mass conservation is maintained to machine truncation

for all simulations (not shown). From the shallow-water equations, the values of g_c and f for the Earth are used,

$$5 \quad g_c = 9.80616 \text{ m s}^{-2}, \quad f = 2\Omega \sin \phi \quad \Omega = 7.29212 \times 10^{-5} \text{ s}^{-1}. \quad (41)$$

All simulations are performed with $n_p = 4$. A thorough investigation of different values of n_p would greatly extend the length of the manuscript, so n_p was chosen in accordance with the Community Atmosphere Model spectral element dynamical core. As argued by Ullrich (2013), this choice is also “optimal” when considering the accurate treatment of linear waves.

When required, the standard L_2 error measure is calculated via

$$L_2(h) = \sqrt{\frac{I[(h - h_T)^2]}{I[h_T^2]}}, \quad (42)$$

where h_T is the height field at the initial time (which is the analytical solution for steady-state test cases) and I denotes an approximation to the global integral, given by

$$15 \quad I[x] = \sum_{\text{all elements } k} \left[\sum_{m=0}^{n_p-1} \sum_{n=0}^{n_p-1} x_k(\alpha_m, \beta_n) J_k(\alpha_m, \beta_n) w_m w_n \Delta\alpha \Delta\beta \right], \quad (43)$$

where the subscript k denotes the values of x and J within the k^{th} element.

When applied, hyperviscosity uses a single coefficient for both scalar and vector hyperviscosity,

$$\nu = \nu_d = \nu_v = (1.0 \times 10^{15} \text{ m}^4 / \text{s}) \left(\frac{n_e}{30} \right)^{3.2}. \quad (44)$$

20 This choice of scaling for the hyperviscosity coefficient is based on Takahashi et al. (2006).

6.1 Steady-State Geostrophically Balanced Flow

Test case 2 of Williamson et al. (1992) ~~simulates~~ describes a zonally symmetric geostrophically balanced flow. This test utilizes an unstable equilibrium solution to the shallow-water equations which is expected to be exactly maintained over time. However, it is generally true that

only methods that satisfy the curl-grad annihilator property $\nabla \times \nabla \phi = 0$ maintain some sort of discrete equilibrium. Nonetheless, since an analytical solution is known (identical to the initial conditions), this test is effective at measuring the convergence rate of a numerical method. Further, the error fields from this test provide some indication of what effect the grid has on the errors of the underlying method. The analytical height field for this test is given by

$$h = h_0 - \frac{1}{g_c} \left(\Omega u_0 a + \frac{u_0^2}{2} \right) \sin^2 \phi, \quad (45)$$

with background height h_0 and velocity amplitude u_0 chosen to be

$$h_0 = \frac{2.94 \times 10^4 \text{ m}^2 \text{ s}^{-2}}{g_c}, \quad \text{and} \quad u_0 = \frac{\pi a}{6} \text{ day}^{-1}. \quad (46)$$

This height field also serves as the reference solution. The non-divergent velocity field is specified in latitude-longitude (ϕ, λ) coordinates as

$$u_\lambda = u_0 \cos \phi, \quad u_\phi = 0. \quad (47)$$

Figure 3 shows L_2 errors in the height field after a 5 day integration of the model at $n_e = 4$ resolution with $n_p = 4$. Simulations were completed for continuous elements (a) with hyperviscosity and (d) without hyperviscosity, discontinuous elements (b,e) with mass lumping (the g_2 flux correction function), (c,f) without mass lumping (the g_1 flux correction function), (b,c) with discontinuous penalization, and (e,f) without discontinuous penalization. The time step is $\Delta t = 2200$ s for simulations (a,d), $\Delta t = 800$ s for simulations (b,c,e), and $\Delta t = 400$ s for simulation (f). Increasing the magnitude of the time step by 100 s led to simulation instability in each case. Since the addition of hyperviscosity leads to loss of energy conservation there is a slow decay of the geostrophically balanced flow towards a uniform height state, hence leading to a nearly zonally symmetric decay in the height field towards the poles. For all configurations (both continuous and discontinuous elements) visually identical results are observed when hyperviscosity is added, and so these results are not shown. All simulations exhibit a characteristic wavenumber-4 mode triggered by the underlying cubed-sphere, although the specific error

pattern differs throughout. Simulation (d) is exactly mimetic and leads to exact maintenance of geostrophic balance. Simulations (b) and (c) are quasi-mimetic, only losing energy conservation due to the discontinuous penalty term, and so exhibit very slow error growth with time. Simulations (e) and (f), which correspond to discontinuous elements without penalization, show greatly enhanced error norms and substantial imprinting from the $n_e = 4$ pattern.

To understand the growth of error norms associated with each configuration, additional simulations with $n_e = 16$ have been performed and L_2 error norms plotted as a function of time in Fig. 4. All simulations show an expected near-identical growth of errors with time when hyperviscosity is active. With hyperviscosity disabled the results from each simulation disentangle: Continuous elements are oscillatory but show stable error norms, discontinuous elements with penalization show smaller error norms than continuous elements but a very slow growth with time due to the upwinding effect of the discontinuous penalization, and discontinuous elements without penalization show rapid growth in errors (and ~~eventual~~even instability without mass lumping).

To verify that the model exhibits the correct convergence rate, Fig. 5 shows the global error norms associated with simulations with $n_e \in \{4, 8, 16, 32, 64\}$ over a 5 day integration period. At $n_e = 4$, the time step is $\Delta t = 2200$ s for continuous elements, $\Delta t = 800$ s for g_2 discontinuous elements and g_1 discontinuous elements with penalization, and $\Delta t = 400$ s for g_1 discontinuous elements without penalization. Increasing the time step by 100 s led to an unstable simulation. The time step is scaled inversely with increasing resolution. Missing simulations correspond to model instability and divergence prior to simulation completion. The use of hyperviscosity reduces the convergence rate to $\mathcal{O}(\Delta x^{3.2})$, as expected from the choice of hyperviscosity coefficient (44). With hyperviscosity disabled, model simulations converge at $\mathcal{O}(\Delta x^4)$ for continuous elements and discontinuous elements with penalty, and $\mathcal{O}(\Delta x^3)$ for discontinuous elements without penalty. The loss of one order of accuracy is due to one-sided differentiation at co-located nodes along element edges, leading to enhancement of the discontinuity. Similar results (not shown) are observed when changing n_p – that is, continuous elements and discontinuous elements with penalty converge at $\mathcal{O}(\Delta x^{n_p})$, whereas unpenalized discontinuous elements converge at $\mathcal{O}(\Delta x^{n_p-1})$.

6.2 Zonal flow over an isolated mountain

Test case 5 in Williamson et al. (1992) considers zonal flow with underlying topography. The wind and height fields are defined as in section 6.1, except with $h_0 = 5960$ m and $u_0 = 20$ m s⁻¹. A conical mountain is used for the topographic forcing, given by

$$z = z_0(1 - r/R), \quad (48)$$

with $z_0 = 2000$ m, $R = \pi/9$ and $r^2 = \min [R^2, (\lambda - \lambda_c)^2 + (\phi - \phi_c)^2]$. The center of the mountain is at $\lambda_c = 3\pi/2$ and $\phi_c = \pi/6$.

Simulation results for this test case were computed at $n_e = 16$ and $n_p = 4$ after 15 days of integration both with and without hyperviscosity. For discontinuous elements penalization is always used. The time step used for these runs was $\Delta t = 480$ s for continuous elements, $\Delta t = 240$ s for g_2 discontinuous elements and $\Delta t = 120$ s for g_1 discontinuous elements. [Increasing the time step by 20 s led to an unstable simulation.](#) These results are visually indistinguishable, so are instead compared against the continuous element run (with HV) in Fig. 6, where the height field h and height field difference $h - h_c$ is plotted (where h_c is the height field given in (a)). Simulations (b) and (c), corresponding to discontinuous elements with and without mass lumping, are very similar in structure and exhibit smooth differences from the continuous model. With no hyperviscosity applied, continuous elements (d) show significant noise which is not present for discontinuous elements (e,f). These simulations match closely with results from the literature (Nair et al., 2005; Ullrich et al., 2010)

To understand conservation of invariants over time, total energy E and potential enstrophy ξ are computed over the duration of the simulation. Since these quantities are invariant under the shallow-water equations, it would be expected that a perfect simulation would conserve these quantities exactly. They are defined via

$$E = \frac{1}{2}h\mathbf{v} \cdot \mathbf{v} + \frac{1}{2}\underline{G}g_c(H^2 - z^2), \quad \text{and} \quad \xi = \frac{(\zeta + f)^2}{2h}. \quad (49)$$

A time series of energy and potential enstrophy are plotted in Fig. 7. With hyperviscosity (a,b) all simulations exhibit nearly identical conservation properties, suggesting that both the contin-

uous and discontinuous hyperviscosity operators (which are responsible for the loss of energy and potential enstrophy conservation) act in a nearly identical manner over the course of the simulation. Without hyperviscosity (c,d) change in energy and potential enstrophy is much smaller. Continuous elements show initiation of instability at approximately day 6, likely due to high-wavenumber oscillations in the height field caused by nonlinear aliasing. On the other hand, discontinuous elements instead show a slow decay of energy and potential enstrophy driven by the weak diffusivity of the discontinuous penalization.

6.3 Rossby-Haurwitz Wave

Test case 6 in Williamson et al. (1992) consists of a westward-propagating Rossby-Haurwitz wave that exactly solves the barotropic vorticity equation, but only approximately solves the nonlinear shallow water equations. This test is particularly interesting since it is known to be sensitive to the choice of horizontal viscosity.

Results for the Rossby-Haurwitz wave are given in Fig. 8 and 9 for $n_e = 16$ and $n_p = 4$ horizontal resolution after 14 days of integration. The time step used for these runs was $\Delta t = 480$ s for continuous elements, $\Delta t = 200$ s for g_2 discontinuous elements and $\Delta t = 120$ s for g_1 discontinuous elements. Increasing the time step by 20 s led to an unstable simulation. As expected, there are significant differences in the height field which are induced by the addition of the hyperviscosity (although both simulations appear reasonable given the coarse horizontal resolution). Except for this difference, the results are nonetheless analogous to zonal flow over an isolated mountain: SE is unstable without the addition of hyperviscosity, whereas DG with penalization is effective at stabilizing the method for both lumped and non-lumped variants.

6.4 Barotropic instability

The barotropic instability test case of Galewsky et al. (2004) consists of a zonal jet with compact support at a latitude of 45° , with a latitudinal profile roughly analogous to a much stronger version of test case 3 of Williamson et al. (1992). A small height perturbation is added atop the jet which leads to the controlled formation of an instability in the flow. The relative vorticity

of the flow field at day 6 can then be visually compared against a high-resolution numerically computed solution Galewsky et al. (2004); St-Cyr et al. (2008).

5 Simulation results for this test case were computed at $n_e = 32$ and $n_p = 4$ after 12 days of integration with hyperviscosity enabled. The time step used for these runs was $\Delta t = 150$ s for continuous elements, $\Delta t = 75$ s for g_2 discontinuous elements and $\Delta t = 50$ s for g_1 discontinuous elements. Increasing the time step by 10 s led to an unstable simulation. Simulations are again compared against the continuous element run (with HV) in Fig. 10, where the relative vorticity field ζ and relative vorticity field difference $\zeta - \zeta_c$ is plotted (where ζ_c is the height field given in (a)). Due to the presence of sharp frontal activity in this test case and the strong resolution dependence of this problem (Ullrich et al., 2010), differences in ζ are of the same magnitude as the original field. In particular, the simulations without hyperviscosity (d,e,f) all show enhancement near wave fronts which is not apparent in the simulations with hyperviscosity (b,c). Although most differences can be found near sharp fronts, there is also a clear enhancement in the differences near 120E associated with a trailing instability. For continuous elements without hyperviscosity (c), there is also apparent grid-scale noise which is missing from the other simulations, suggesting that this method is under-diffused.

20 Normalized total energy and potential enstrophy are plotted for the barotropic instability in Fig. 11 for a 12 day integration with $n_e = 16$ and $n_p = 4$. With hyperviscosity (a,b) there are small but visible differences in the results associated with changes in the type of elements. Without hyperviscosity (c,d) the simulation with continuous elements exhibit instability around day 6, leading to rapid growth of energy and potential enstrophy. On the other hand, with discontinuous elements there is a steady loss of energy and potential enstrophy over time due to diffusivity from discontinuous penalization. Prior to wave breaking (which occurs around day 4), energy and potential enstrophy loss are significantly reduced compared to the simulations with hyperviscosity. After wave breaking, energy and potential enstrophy loss are of the same order of magnitude for simulations with and without hyperviscosity, associated with the fact that diffusivity is enhanced near the barotropic fronts where discontinuities are large.

7 Conclusions

Following Huynh (2007), a novel nodal finite element method for continuous and discontinuous elements has been constructed using a robust derivative operator and discontinuous penalization. The resulting methodology can be used for straightforward discretization of partial differential equations in either conservative or non-conservative form. A hyperviscosity operator valid for both continuous and discontinuous elements was also presented that would provide a mechanism for numerical stabilization and the removal of grid-scale noise. Two versions with discontinuous elements were studied, using either the g_1 and g_2 flux correction function for distribution of boundary fluxes and penalty across nodal points. The resulting method was then applied to the 2D shallow-water equations in cubed-sphere geometry and tested on a suite of test problems.

From the Williamson et al. (1992) test case suite, steady-state geostrophically balanced flow and zonal flow over an isolated mountain and the Rossby-Haurwitz wave were examined, in addition to the barotropic instability test of Galewsky et al. (2004). The method was shown to be stable and accurate for both continuous and discontinuous elements, with fourth-order convergence being verified for cubic basis functions. Discontinuous penalization was shown to be necessary for stability and for maintaining the correct order of accuracy of the discontinuous method. Overall the discontinuous elements required a smaller time step than for continuous elements, although all methods led to similar error norms when hyperviscosity was active. When hyperviscosity was deactivated, the discontinuous method exhibited better error norms than the continuous approach and discontinuous penalization was shown to be sufficient for stability of the method even for complex flows. Nonetheless, differences between all three approaches appeared minor, and all methods performed well for this suite of tests.

The non-conservative discontinuous element formulation has been shown to be a potential candidate for future atmospheric modeling. It has the advantage of requiring fewer parallel communications than continuous methods, and exhibits stability even when hyperviscosity is not used for explicit stabilization. However, with the reduced time step size it remains unclear

whether the discontinuous formulation would be computationally competitive with continuous element methods.

- 5 The method discussed in this paper has been implemented in the Tempest atmospheric model, available from <https://github.com/paullic/tempestmodel>.

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Equivalence of differential and variational forms

5 In this appendix equivalence of the variational formulation of the spectral element method and the differential formulation using the robust derivative is demonstrated. For continuous elements, $\tilde{f} = f$ and (15) reduces to

$$D_\alpha f(\alpha_i, \beta_j) = \sum_{p=0}^{n_p-1} f_{(p,j)} \frac{\partial \tilde{\phi}_{(p)}}{\partial \alpha}(\alpha_i), \tag{A1}$$

which is simply the derivative of the continuous analogue to the nodal values along $\beta = \beta_j$.

10 For simplicity consider a single quadrilateral spectral element with test functions ϕ_{ij} located at nodal points $(\alpha_i, \beta_j), (i, j) \in [0, \dots, n_p - 1]^2$. The result is shown for an arbitrary 2D conservation law,

$$\frac{\partial \psi}{\partial t} + \nabla \cdot \mathbf{F} = 0. \tag{A2}$$

Using the derivative operator (A1) this equation reads

$$15 \frac{\partial \psi_{ij}}{\partial t} + \frac{1}{J_{ij}} D_\alpha (JF^\alpha) + \frac{1}{J_{ij}} D_\beta (JF^\beta) = 0, \tag{A3}$$

whereas under the variational formulation (A2) is formulated as

$$\int \frac{\partial \psi}{\partial t} \phi_{ij} dA + \int \phi_{ij} \nabla \cdot \mathbf{F} dA = 0. \tag{A4}$$

Then using integration by parts,

$$\sum_{m,n} \left(\int \phi_{ij} \phi_{mn} dA \right) \frac{\partial \psi_{mn}}{\partial t} + B - \int \nabla \phi_{ij} \cdot \mathbf{F} dA = 0, \quad (\text{A5})$$

5 where B is the contribution due to the boundary which disappears under DSS. Introducing coordinates (α, β) with integration on GLL nodes,

$$\int f dA = \sum_{s=0}^{n_p-1} \sum_{t=0}^{n_p-1} f_{st} J_{st} w_s w_t \Delta \alpha \Delta \beta, \quad (\text{A6})$$

and so the first term of (A5) reads

$$\begin{aligned} \sum_{m,n} \left(\int \phi_{ij} \phi_{mn} dA \right) \frac{\partial \psi_{mn}}{\partial t} &= \sum_{m,n} (\delta_{i,m} \delta_{j,n} J_{ij} w_i w_j \Delta \alpha \Delta \beta) \frac{\partial \psi_{mn}}{\partial t} \\ &= J_{ij} w_i w_j \Delta \alpha \Delta \beta \frac{\partial \psi_{ij}}{\partial t}. \end{aligned} \quad (\text{A7})$$

For the last term, observe that on a manifold

$$\nabla \phi_{ij} \cdot \mathbf{F} = g_{pq} F^p \left(g^{qr} \frac{\partial \phi}{\partial x^r} \right) = F^\alpha \frac{\partial \phi}{\partial \alpha} + F^\beta \frac{\partial \phi}{\partial \beta}, \quad (\text{A8})$$

and so

$$\int \nabla \phi_{ij} \cdot \mathbf{F} dA = \sum_{s=0}^{n_p-1} \sum_{t=0}^{n_p-1} \left[F^\alpha \frac{\partial \phi_{ij}}{\partial \alpha} + F^\beta \frac{\partial \phi_{ij}}{\partial \beta} \right]_{\alpha=\alpha_s, \beta=\beta_t} J_{st} w_s w_t \Delta \alpha \Delta \beta. \quad (\text{A9})$$

But, by construction,

$$\frac{\partial \phi_{ij}}{\partial \alpha} = \frac{\partial \tilde{\phi}_{(i)}}{\partial \alpha} \tilde{\phi}_{(j)}, \quad (\text{A10})$$

5 and $\tilde{\phi}_{(j)}(\beta_t) = \delta_{jt}$. This leads to

$$\int \nabla \phi_{ij} \cdot \mathbf{F} dA = \left[\sum_{s=0}^{n_p-1} F_{sj}^\alpha \frac{\partial \tilde{\phi}_{(i)}}{\partial \alpha}(\alpha_s) J_{sj} w_s w_j + \sum_{t=0}^{n_p-1} F_{it}^\beta \frac{\partial \tilde{\phi}_{(j)}}{\partial \beta}(\beta_t) J_{it} w_i w_t \right] \Delta \alpha \Delta \beta. \quad (\text{A11})$$

Then in conjunction with (A7) this can be written as

$$\frac{\partial \psi_{ij}}{\partial t} - \frac{1}{J_{ij}} \sum_{s=0}^{n_p-1} J_{sj} F_{sj}^\alpha \frac{\partial \tilde{\phi}_{(i)}}{\partial \alpha}(\alpha_s) \frac{w_s}{w_i} - \frac{1}{J_{ij}} \sum_{t=0}^{n_p-1} J_{it} F_{it}^\beta \frac{\partial \tilde{\phi}_{(j)}}{\partial \beta}(\beta_t) \frac{w_t}{w_j} = 0. \quad (\text{A12})$$

10 Equivalence of this equation with (A3) follows for a formulation on GLL nodes (Boyd, 2001, Appendix) since these basis functions satisfy the property

$$\frac{\partial \tilde{\phi}_{(i)}}{\partial \alpha}(\alpha_s) w_s = - \frac{\partial \tilde{\phi}_{(s)}}{\partial \alpha}(\alpha_i) w_i. \quad (\text{A13})$$

Appendix B

Derivation of the Viscosity Operator

In this appendix the derivation of the discrete viscosity operator is provided for scalar and vector hyperviscosity on a Riemannian manifold.

B1 Scalar Viscosity

From the natural quadrature rule that arises from the nodal finite element formulation, the left-hand-side of (21) is discretized as

$$\iint f \phi_{(i,j)} dA = \iint f \tilde{\phi}_{(i)}(\alpha) \tilde{\phi}_{(j)}(\beta) dA = f_{(i,j)} w_i w_j J \Delta \alpha \Delta \beta, \quad (\text{B1})$$

and so, pointwise, the \mathcal{H} operator is applied via

$$f_{(i,j)} = \frac{\nu}{w_i w_j \Delta \alpha \Delta \beta J(\alpha_i, \beta_j)} \left[\oint_{\partial \mathcal{Z}} \phi_{(i,j)} \nabla \psi \cdot d\mathbf{S} - \iint_{\mathcal{Z}} \nabla \phi_{(i,j)} \cdot \nabla \psi dA \right]. \quad (\text{B2})$$

The area integral term in (B2) is then computed:

$$\begin{aligned} \iint \nabla \phi_{(i,j)} \cdot \nabla \psi dA &= \iint \nabla_p \phi \nabla^p \psi dA = \iint \frac{\partial \phi_{(i,j)}}{\partial \alpha} \nabla^\alpha \psi + \frac{\partial \phi_{(i,j)}}{\partial \beta} \nabla^\beta \psi dA, \\ &= \Delta \alpha \Delta \beta \sum_{m=0}^{n_p-1} \sum_{n=0}^{n_p-1} \tilde{\phi}_{(j)} \frac{\partial \tilde{\phi}_{(i)}}{\partial \alpha} \nabla^\alpha \psi J w_m w_n \Big|_{\alpha=\alpha_m, \beta=\beta_n} \\ &\quad + \Delta \alpha \Delta \beta \sum_{m=0}^{n_p-1} \sum_{n=0}^{n_p-1} \tilde{\phi}_{(i)} \frac{\partial \tilde{\phi}_{(j)}}{\partial \beta} \nabla^\beta \psi J w_m w_n \Big|_{\alpha=\alpha_m, \beta=\beta_n} \\ &= \Delta \alpha \Delta \beta w_j \sum_{m=0}^{n_p-1} \frac{\partial \tilde{\phi}_{(i)}}{\partial \alpha} \nabla^\alpha \psi J w_m \Big|_{\alpha=\alpha_m, \beta=\beta_j} \\ &\quad + \Delta \alpha \Delta \beta w_i \sum_{n=0}^{n_p-1} \frac{\partial \tilde{\phi}_{(j)}}{\partial \beta} \nabla^\beta \psi J w_n \Big|_{\alpha=\alpha_i, \beta=\beta_n} \end{aligned} \quad (\text{B3})$$

From (B2), (23) then follows. The boundary integral term in (21) takes the form

$$\oint_{\partial Z} \phi_{(i,j)} \nabla \psi \cdot d\mathbf{S} = \int_{\partial Z_R} \phi_{(i,j)} \nabla \psi \cdot d\mathbf{S} + \int_{\partial Z_T} \phi_{(i,j)} \nabla \psi \cdot d\mathbf{S} + \int_{\partial Z_L} \phi_{(i,j)} \nabla \psi \cdot d\mathbf{S} + \int_{\partial Z_B} \phi_{(i,j)} \nabla \psi \cdot d\mathbf{S}, \quad (\text{B4})$$

where R , T , L and B denote the right, top, left and bottom edges, respectively. The quantity $d\mathbf{S} = \mathbf{N}d\ell$ denotes the normal vector to the edge with magnitude equal to the infinitesimal length element. Only the covariant components of the face normals need to be known, at each edge given by

$$N_p^R = \left(\frac{1}{\sqrt{g^{\alpha\alpha}}}, 0 \right), \quad N_p^T = \left(0, \frac{1}{\sqrt{g^{\beta\beta}}} \right), \quad N_p^L = \left(-\frac{1}{\sqrt{g^{\alpha\alpha}}}, 0 \right), \quad N_p^B = \left(0, -\frac{1}{\sqrt{g^{\beta\beta}}} \right), \quad (\text{B5})$$

The infinitesimal length element along each edge is given by the covariant metric,

$$d\ell^R = \sqrt{g_{\beta\beta}}d\beta, \quad d\ell^T = \sqrt{g_{\alpha\alpha}}d\alpha, \quad d\ell^L = \sqrt{g_{\beta\beta}}d\beta, \quad d\ell^B = \sqrt{g_{\alpha\alpha}}d\alpha. \quad (\text{B6})$$

Then along the right edge, using the nodal discretization of the boundary integral,

$$\begin{aligned} \int_{\partial Z_R} \phi_{(i,j)} \nabla \psi \cdot d\mathbf{S} &= \delta_{i,n_p-1} \sum_{n=0}^{n_p-1} \tilde{\phi}_{(j)}(\beta) \nabla^\alpha \psi N_\alpha^R w_n \sqrt{g_{\beta\beta}} \Delta\beta \Big|_{\alpha=\alpha_{n_p-1}, \beta=\beta_n} \\ &= \delta_{i,n_p-1} w_j \Delta\beta J \nabla^\alpha \psi \Big|_{\alpha=\alpha_{n_p-1}, \beta=\beta_j}, \end{aligned} \quad (\text{B7})$$

where we have used $g_{\beta\beta} = J^2 g^{\alpha\alpha}$. Repeating for all edges and using (B2) then yields (24).

B2 Vector Viscosity

The area integral that appears on the left-hand-side of (29) and (30) takes the form

$$5 \quad \iint_{\underline{z}} \mathbf{f} \cdot \boldsymbol{\phi}_{(i,j)}^{(\alpha)} dA = \iint_{\underline{z}} f^\alpha \tilde{\phi}_{(i)}(\alpha) \tilde{\phi}_{(j)}(\beta) dA = f_{(i,j)}^\alpha w_i w_j J \Delta \alpha \Delta \beta, \quad (\text{B8})$$

$$\iint_{\underline{z}} \mathbf{f} \cdot \boldsymbol{\phi}_{(i,j)}^{(\beta)} dA = \iint_{\underline{z}} f^\beta \tilde{\phi}_{(i)}(\alpha) \tilde{\phi}_{(j)}(\beta) dA = f_{(i,j)}^\beta w_i w_j J \Delta \alpha \Delta \beta. \quad (\text{B9})$$

B2.1 Discretization of the area integral

In nodal form, the divergence expands as

$$10 \quad (\nabla \cdot \boldsymbol{\phi}_{(i,j)}^{(\alpha)}) = \frac{1}{J} \frac{\partial}{\partial \alpha} (J g^{\alpha\alpha} \phi_{(i,j)\alpha}) + \frac{1}{J} \frac{\partial}{\partial \beta} (J g^{\beta\alpha} \phi_{(i,j)\alpha}) \quad (\text{B10})$$

$$= \frac{\tilde{\phi}_{(j)}(\beta)}{J} \frac{\partial}{\partial \alpha} (J g^{\alpha\alpha} \tilde{\phi}_{(i)}(\alpha)) + \frac{\tilde{\phi}_{(i)}(\alpha)}{J} \frac{\partial}{\partial \beta} (J g^{\beta\alpha} \tilde{\phi}_{(j)}(\beta)), \quad (\text{B11})$$

and so

$$\begin{aligned}
& \iint_{\mathcal{Z}} (\nabla \cdot \boldsymbol{\phi}_{(i,j)}) (\nabla \cdot \mathbf{u}) dA \\
5 \quad &= \Delta\alpha\Delta\beta \sum_{m=0}^{n_p-1} \sum_{n=0}^{n_p-1} \left[\frac{\tilde{\phi}_{(j)}(\beta_n)}{J} \frac{\partial}{\partial\alpha} \left(Jg^{\alpha\alpha} \tilde{\phi}_{(i)}(\alpha) \right) \right. \\
& \quad \left. + \frac{\tilde{\phi}_{(i)}(\alpha_m)}{J} \frac{\partial}{\partial\beta} \left(Jg^{\beta\alpha} \tilde{\phi}_{(j)}(\beta) \right) \right] (\nabla \cdot \mathbf{u}) Jw_m w_n \\
&= \Delta\alpha\Delta\beta w_j \sum_{m=0}^{n_p-1} Jg^{\alpha\alpha} \frac{d\tilde{\phi}_{(i)}}{d\alpha} (\nabla \cdot \mathbf{u}) w_m \Big|_{\alpha=\alpha_m, \beta=\beta_j} \\
& \quad + \Delta\alpha\Delta\beta w_i \sum_{n=0}^{n_p-1} Jg^{\beta\alpha} \frac{d\tilde{\phi}_{(j)}}{d\beta} (\nabla \cdot \mathbf{u}) w_n \Big|_{\alpha=\alpha_i, \beta=\beta_n} \tag{B12}
\end{aligned}$$

10 Further, the radial component of the vorticity expands as

$$(\nabla \times \boldsymbol{\phi}_{(i,j)})^r = -\frac{1}{J} \frac{\partial \phi_{(i,j)\alpha}}{\partial\beta} = -\frac{\tilde{\phi}_{(i)}}{J} \frac{d\tilde{\phi}_{(j)}}{d\beta} \tag{B13}$$

and so

$$\begin{aligned}
& \iint_{\mathcal{Z}} (\nabla \times \boldsymbol{\phi}_{(i,j)})^r (\nabla \times \mathbf{u})_r dA \\
15 \quad &= \Delta\alpha\Delta\beta \sum_{m=0}^{n_p-1} \sum_{n=0}^{n_p-1} \left[-\frac{\tilde{\phi}_{(i)}(\alpha_m)}{J} \frac{d\tilde{\phi}_{(j)}}{d\beta} \right] (\nabla \times \mathbf{u})_r Jw_m w_n \Big|_{\alpha=\alpha_m, \beta=\beta_n} \\
&= -\Delta\alpha\Delta\beta w_i \sum_{n=0}^{n_p-1} \frac{d\tilde{\phi}_{(j)}}{d\beta} (\nabla \times \mathbf{u})_r w_n \Big|_{\alpha=\alpha_i, \beta=\beta_n} \tag{B14}
\end{aligned}$$

Combining (B8), (B12) and (B14) then gives (32). An analogous procedure for β leads to (33).

B2.2 Discretization of the boundary integral

5 Using (B5)-(B6) and $\sqrt{g_{\beta\beta}} = J\sqrt{g^{\alpha\alpha}}$, the contour integral in (29) along the right edge becomes

$$\int_{\partial\mathcal{Z}_R} (\nabla \cdot \mathbf{u}) \phi_{(i,j)}^{(\alpha)} \cdot d\mathbf{S} = \delta_{i,n_p-1} (\nabla \cdot \mathbf{u}) g^{\alpha\alpha} J w_j \Delta\beta \Big|_{\alpha=\alpha_{n_p-1}, \beta=\beta_j}, \quad (\text{B15})$$

and along the top edge, also using $\sqrt{g_{\alpha\alpha}} = J\sqrt{g^{\beta\beta}}$,

$$\int_{\partial\mathcal{Z}_T} (\nabla \cdot \mathbf{u}) \phi_{(i,j)}^{(\alpha)} \cdot d\mathbf{S} = \delta_{j,n_p-1} (\nabla \cdot \mathbf{u}) g^{\alpha\beta} J w_i \Delta\alpha \Big|_{\alpha=\alpha_i, \beta=\beta_{n_p-1}} \quad (\text{B16})$$

Repeating for all edges and using (B8), the complete boundary integral for divergence damping then leads to the divergence damping contribution to (34). An analogous procedure for test function $\phi_{(i,j)}^{(\beta)}$ leads to (35).

For vorticity damping, along the right edge (30) reads

$$\int_{\partial\mathcal{Z}_R} (\nabla \times \mathbf{u}) \times \phi \cdot d\mathbf{S} = \delta_{i,n_p-1} e^{\beta r \alpha} (\nabla \times \mathbf{u})_r \phi_{(i,j)\alpha} N_\beta w_j \sqrt{g_{\beta\beta}} \Delta\beta \Big|_{\alpha=\alpha_{n_p-1}, \beta=\beta_j} = 0.$$

and along the top edge,

$$\begin{aligned} \int_{\partial\mathcal{Z}_T} (\nabla \times \mathbf{u}) \times \phi \cdot d\mathbf{S} &= \delta_{j,n_p-1} e^{\beta r \alpha} (\nabla \times \mathbf{u})_r \phi_{(i,j)\alpha} N_\beta w_i \sqrt{g_{\alpha\alpha}} \Delta\alpha \Big|_{\alpha=\alpha_i, \beta=\beta_{n_p-1}}, \\ &= \delta_{j,n_p-1} (\nabla \times \mathbf{u})_r w_i \Delta\alpha. \end{aligned}$$

Repeating for all edges and using (B8) then leads to the vorticity damping contribution to (34). An analogous procedure for test function $\phi_{(i,j)}^{(\beta)}$ leads to (35).

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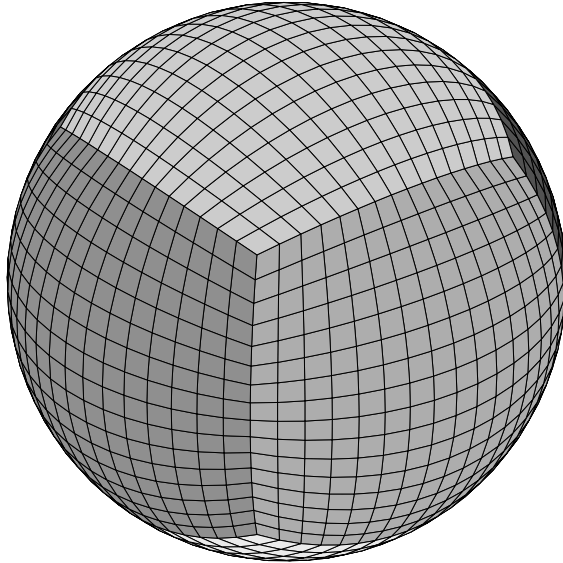


Fig. 1. A 3D view of the cubed-sphere grid shown here with $n_e = 16$. Cubed sphere faces are individually shaded.

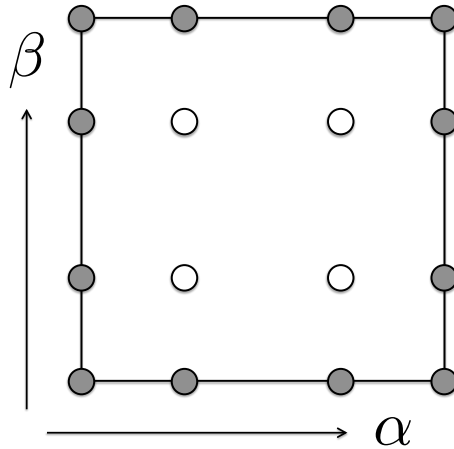


Fig. 2. A depiction of the nodal grid for a reference element on GLL nodes for $n_p = 4$. Boundary nodes, which are connected to neighboring elements, are shaded.

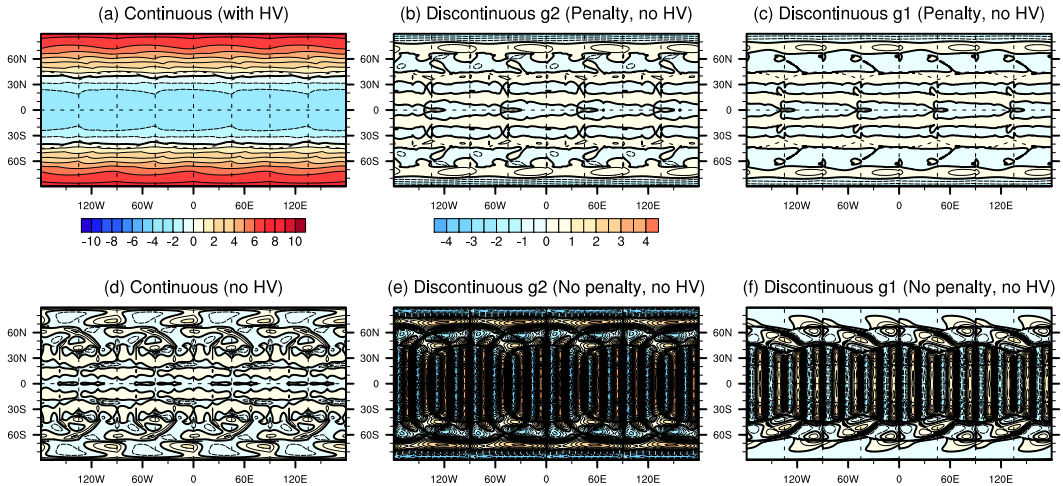


Fig. 3. L_2 errors in Williamson et al. (1992) Test Case 2, steady-state geostrophically balanced flow, for $n_e = 4$ and $n_p = 4$ after a 5 day integration period. Contour spacing for plot (a) is 1 meter. Contour spacing for all other plots is 0.5 meter. The zero line is enhanced. Long dashed lines show the cubed-sphere grid.

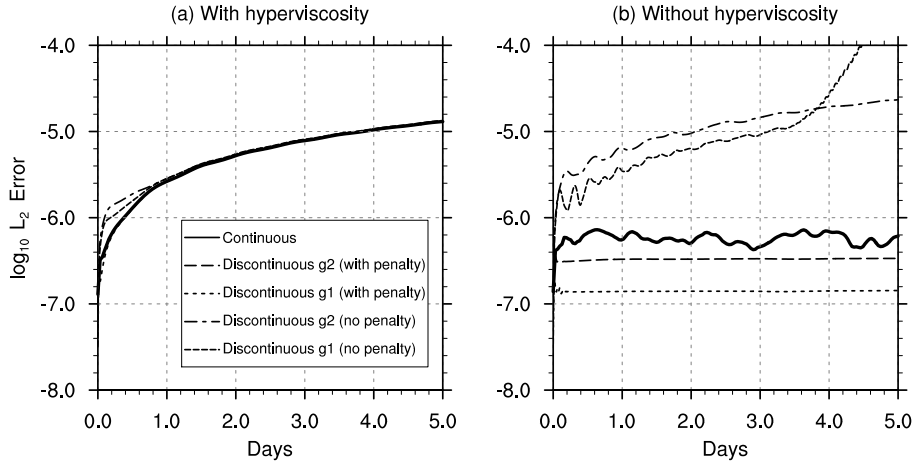


Fig. 4. L_2 error time series for geostrophically balanced flow on the cubed-sphere for $n_e = 16$ and $n_p = 4$ over a 5 day integration period for all continuous and discontinuous schemes.

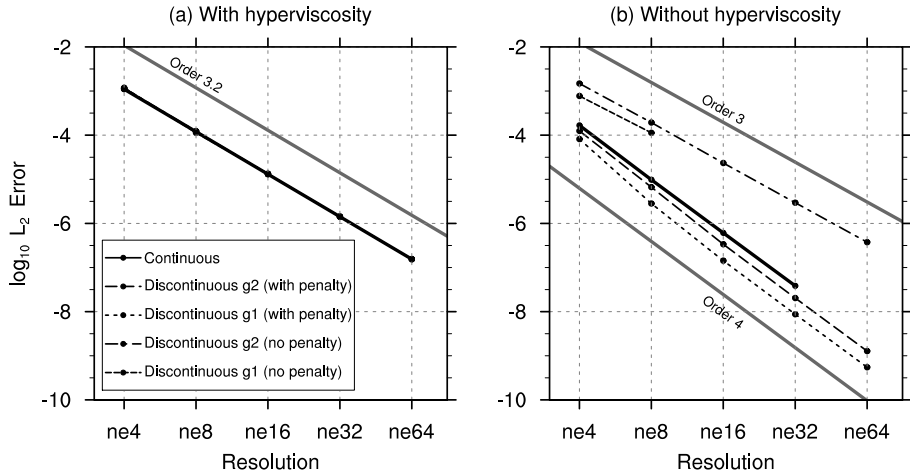


Fig. 5. L_2 errors for geostrophically balanced flow on the cubed-sphere at various resolutions with $n_p = 4$ over a 5 day integration period. In (a) errors due to hyperviscosity dominate and so all simulations have approximately equal error leading to coincident lines. In (b) unstable simulations have been removed.

Zonal Flow over an Isolated Mountain - Height (m) at 15 days

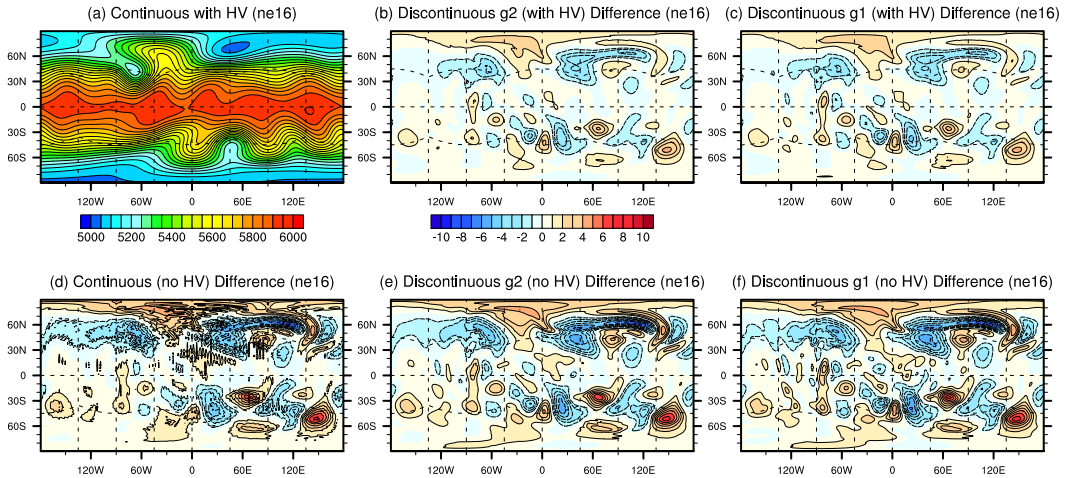


Fig. 6. Height field with $n_e = 16$ and $n_p = 4$ at day 15 for zonal flow over an isolated mountain with (a) continuous elements and hyperviscosity (reference solution). Height difference plot from reference solution with $n_e = 16$ at day 15 for (b) discontinuous g_2 elements with hyperviscosity, (c) discontinuous g_1 elements with hyperviscosity, (d) continuous elements without hyperviscosity, (e) discontinuous g_2 elements without hyperviscosity and (f) discontinuous g_1 elements without hyperviscosity. The time step used for these runs was (a,d) $\Delta t = 480$ s, (b,e) $\Delta t = 240$ s and (c,f) $\Delta t = 120$ s. Discontinuous penalization was used for both discontinuous schemes. Contour spacing is 1 m in all difference plots with the zero line removed. Long dashed lines show the cubed-sphere grid.

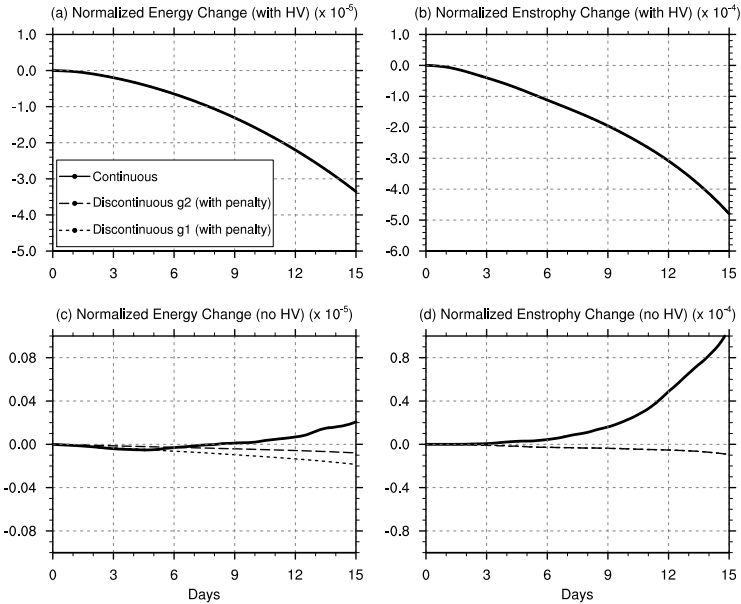


Fig. 7. Normalized total energy and potential enstrophy change for the zonal flow over an isolated mountain test with $n_e = 16$ and $n_p = 4$ over a 15 day simulation. In (a) all simulations show roughly equivalent energy and enstrophy loss and so all lines are coincident. In (c) and (d) the simulation with continuous elements is beginning to experience instability, leading to total energy and enstrophy growth after approximately 6 days simulation time.

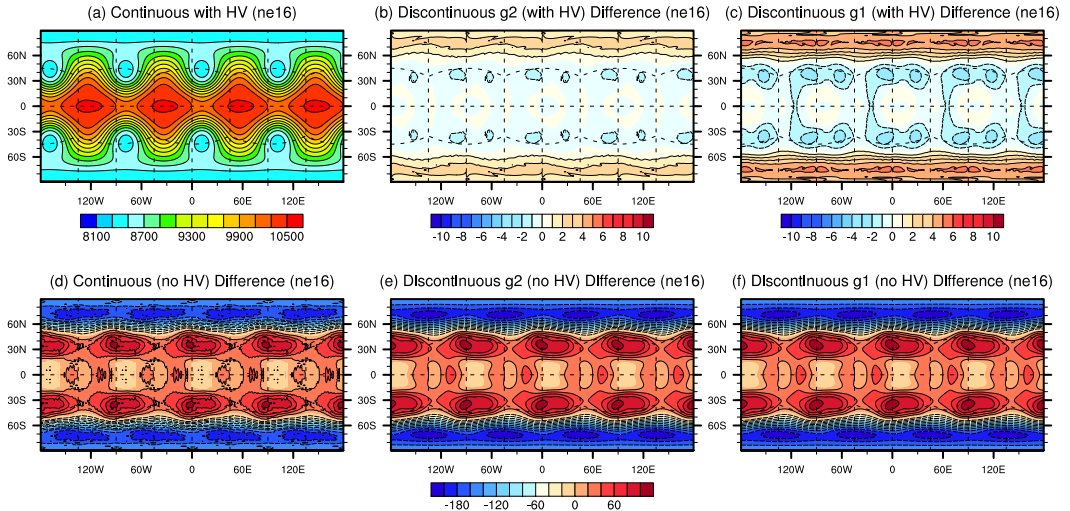


Fig. 8. Height field with $n_e = 16$ and $n_p = 4$ at day 14 for the Rossby-Haurwitz wave with (a) continuous elements and hyperviscosity (reference solution). Height difference plot from reference solution with $n_e = 16$ at day 14 for (b) discontinuous g_2 elements with hyperviscosity, (c) discontinuous g_1 elements with hyperviscosity, (d) continuous elements without hyperviscosity, (e) discontinuous g_2 elements without hyperviscosity and (f) discontinuous g_1 elements without hyperviscosity. The time step used for these runs was (a,d) $\Delta t = 480$ s, (b,e) $\Delta t = 200$ s and (c,f) $\Delta t = 120$ s. Discontinuous penalization was used for both discontinuous schemes. Contour spacing is 1 m in plots (b) and (c) and 20 m in plots (d), (e) and (f). Long dashed lines show the cubed-sphere grid.

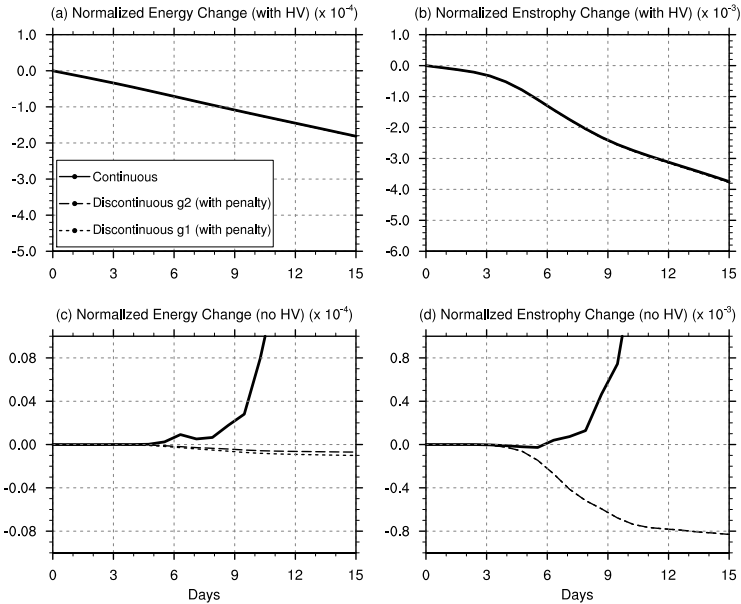


Fig. 9. Normalized total energy and potential enstrophy change for the Rossby-Haurwitz wave test with $n_e = 16$ and $n_p = 4$ over a 15 day simulation. In (a) and (b) all simulations show roughly equivalent energy and enstrophy loss and so all lines are coincident. In (c) and (d) the simulation with continuous elements is beginning to experience instability, leading to total energy and enstrophy growth after approximately 6 days simulation time.

Barotropic Instability Relative Vorticity ($\times 10^{-5} \text{ s}^{-1}$)

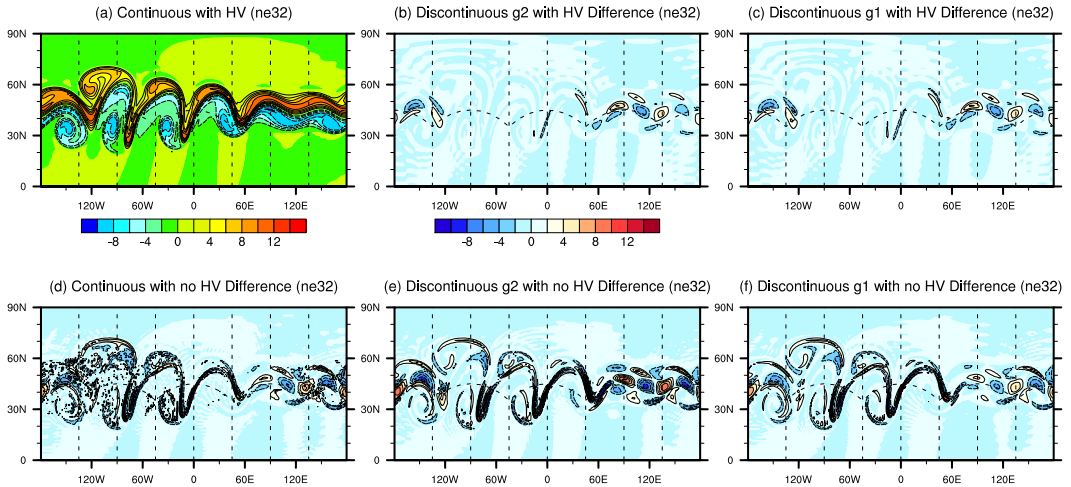


Fig. 10. Relative vorticity field with $n_e = 32$ and $n_p = 4$ at day 6 for the barotropic instability test with (a) continuous elements and hyperviscosity (reference solution). Relative vorticity difference plot from reference solution with $n_e = 16$ at day 6 for (b) discontinuous g_2 elements with hyperviscosity, (c) discontinuous g_1 elements with hyperviscosity, (d) continuous elements without hyperviscosity, (e) discontinuous g_2 elements without hyperviscosity and (f) discontinuous g_1 elements without hyperviscosity. The time step used for these runs was (a,d) $\Delta t = 150$ s, (b,e) $\Delta t = 75$ s and (c,f) $\Delta t = 50$ s. Discontinuous penalization was used for both discontinuous schemes. Contour spacing in all plots is $2 \times 10^{-5} \text{ s}^{-1}$ with the zero line removed. Long dashed lines show the cubed-sphere grid.

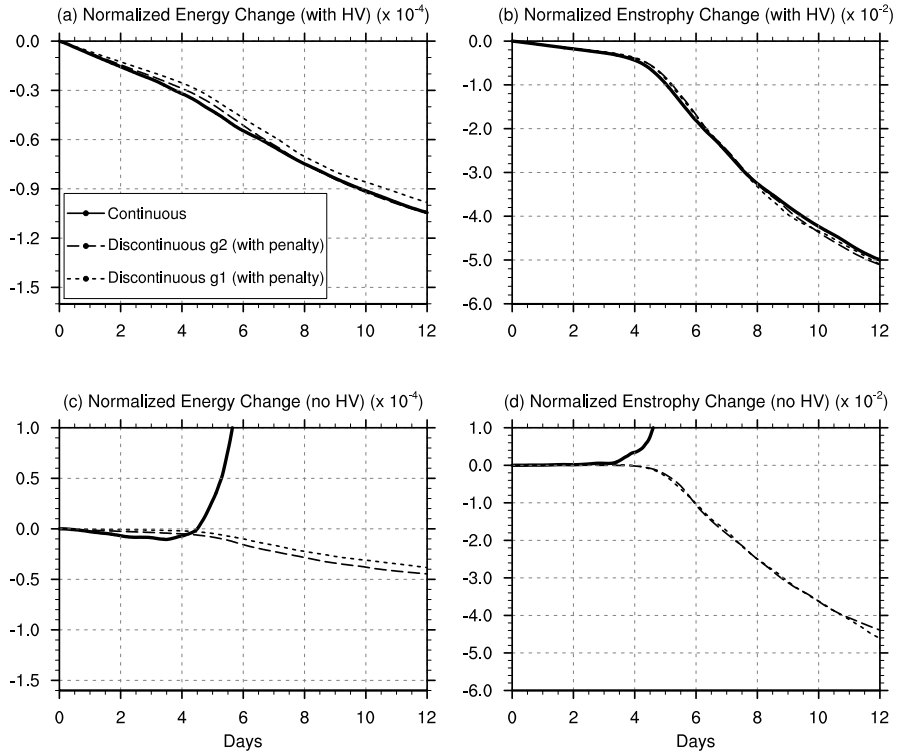


Fig. 11. Normalized total energy and enstrophy change for the barotropic instability test with $n_e = 16$ and $n_p = 4$ over a 12 day simulation. In (c) and (d) the continuous element simulation fails after approximately 6 days, leading to unbounded growth in energy and enstrophy. The time step used for these runs was (a,d) $\Delta t = 300$ s, (b,e) $\Delta t = 150$ s and (c,f) $\Delta t = 75$ s. Discontinuous penalization was used for both discontinuous schemes.