Representing Icebergs in the iLOVECLIM Model (version 1.0)– A Sensitivity Study

M. Bügelmayer[1], D. M. Roche[1,2], H. Renssen[1]

m.bugelmayer@vu.nl, didier.roche@vu.nl, h.renssen@vu.nl

[1] Earth and Climate Cluster, Faculty of Earth and Life Sciences, Vrije Universiteit Amsterdam, Amsterdam, The Netherlands

[2] Laboratoire des Sciences du Climat et de l'Environnement (LSCE), CEA/CNRS-INSU/UVSQ, Gif-sur-Yvette Cedex, France

Recent modelling studies have indicated that icebergs play an active role in the climate system as they interact with the ocean and the atmosphere. The icebergs' impact is due to their slowly released melt water, which freshens and cools the ocean and consequently alters the ocean stratification and the sea ice conditions. The spatial distribution of the icebergs and their melt water depends on the atmospheric and oceanic forces acting on them as well as on the icebergs' size. The studies conducted so far have in common that the icebergs were moved by reconstructed or modelled forcing fields and that the initial size distribution of the icebergs was prescribed according to present day observations. To study the sensitivity of the modelled iceberg distribution to initial and boundary conditions, we performed 15 sensitivity experiments using the climate model *i*LOVECLIM that includes actively coupled ice-sheet and iceberg modules, to analyse 1) the impact of the atmospheric and oceanic forces on the icebergs' distribution and melt flux, and 2) the effect of the used initial iceberg size on the resulting Northern hemisphere climate as well as on the ice sheet, due to feedback mechanisms such as altered atmospheric temperatures, under different climate conditions (pre-industrial, high/low radiative forcing). Our results show that, under equilibrated pre-industrial conditions, the oceanic currents cause the bergs to stay close to the Greenland and North American coast, whereas the atmospheric forcing quickly distributes them further away from their calving site. These different characteristics strongly affect the lifetime of icebergs, since the wind – driven icebergs melt up to two years faster as they are quickly distributed into the relatively warm North Atlantic waters. Moreover, we find that local variations in the spatial distribution due to different iceberg sizes do not result in different climate states and Greenland ice sheet volume, independent of the prevailing climate conditions (pre-industrial, warming or cooling climate). Therefore, we conclude that local differences in the distribution of their melt flux do not alter the prevailing Northern Hemisphere climate and ice sheet under equilibrated conditions and continuous supply of icebergs. Furthermore, our results suggest that the applied radiative forcing scenarios have a stronger impact on climate than the used initial size distribution of the icebergs.

1 *Introduction*

Icebergs are an important part of the climate system as they interact with the ocean, atmosphere and cryosphere (e.g. Hemming, 2004; Smith et al., 2011; Tournadre et al., 2012). Most importantly, icebergs play an important part in the global fresh water cycle since currently up to half of the mass loss of the Antarctic (Rignot et al., 2013) and Greenland ice sheets is due to calving (approx. 0.01 Sv, 1 Sv = $1*10^6$ m^3s^{-1}, Hooke et al., 2005). As icebergs are melting, they affect the upper ocean not only by freshening, but also by cooling due to their uptake of latent heat. Several studies have revealed that the freshening and cooling have opposing effects on ocean stratification, as the cooling enhances the surface density, promoting deep mixing, whereas the freshening decreases the water density, stabilizing the water column (Jongma et al., 2009, 2013, Green et al., 2011).

Moreover, the implementation of dynamical icebergs in climate models has revealed that icebergs enhance the formation of sea ice (Jongma et al., 2009, 2013; Wiersma and Jongma 2010; Bügelmayer et al., 2014), which forms a barrier between the ocean and the atmosphere. Therefore, on the one hand sea ice shields the ocean from being stirred by atmospheric winds, and on the other hand from losing heat to the relatively cold atmosphere, consequently, reducing mixing of the upper water column. Further, this reduced oceanic heat loss leads, in combination with an increase in surface albedo, to a changed atmospheric state (Bügelmayer et al., 2014). Thus, icebergs indirectly alter the ice sheet’s mass balance through their effect on the air temperature and precipitation (Bügelmayer et al., 2014).

The amount of icebergs calved and their effects on climate depend on the calving flux provided by the ice sheets, which is altered by the prevailing climate conditions. For instance, in the relatively cold climate of the last glacial episodic discharges of icebergs into the North Atlantic Ocean, so-called Heinrich events, have been recorded in distinct layers of ice rafted debris (Andrews 1998; Hemming 2004). These periods of enhanced ice discharge have been proposed to be caused by ice shelf collapses (e.g. MacAyeal, 1993; Hulbe et al., 2004; Alvarez-Solas et al., 2011) and happened during periods of a (partial) collapse of the thermohaline circulation (Broecker et al., 1993; McManus et al. 2004; Gherardi et al., 2005; Kageyama et al., 2010). It has been suggested that the collapse was caused by the long duration (Marcott et al., 2011) and the increased amount of freshwater released (0.04 up to 0.4 Sv, Roberts et al., 2014) and affected the global climate.

So far, different approaches have been taken to incorporate icebergs from the Antarctic and Greenland ice sheets into numerical models for different time periods. Bigg et al. (1996, 1997) presented an iceberg module, which was fed with present-day atmospheric and oceanic input fields. The forcing was provided off-line by atmospheric and oceanic models to investigate the drift patterns of icebergs in the Northern Hemisphere. Their approach was further developed for the Southern Ocean by Gladstone et al. (2001), who used modelled oceanic forcing and modern reconstructed wind fields, as well as observed calving amounts to seed the iceberg module. Subsequently, the same iceberg module was implemented in an earth system model of intermediate complexity (EMIC) by Jongma et al. (2009) to investigate the impact of icebergs on the Southern Ocean under pre-industrial conditions. In the latter study, the icebergs were seeded based on a prescribed constant calving flux based on observational estimates, but moved according to the modeled winds and currents and interacted with the model atmosphere and ocean. Martin and Adcroft (2010) then implemented the iceberg model into a coupled global climate model (CGCM) using the model’s variable runoff as a calving flux though still lacking an ice sheet component. Most recently, Bügelmayer et al. (2014) took the next step by using an EMIC with both dynamically coupled ice sheet and iceberg model components. In their model setup, the climate – ice-sheet – iceberg system was fully interactive, with the icebergs’ calving positions and amounts being determined by the ice sheet model, and with the ice sheet responding to the icebergs' effect on climate.

Coupled climate-iceberg models have been used for several specific purposes, such as the investigation of drift patterns of icebergs under present-day (Venkatesh and El-Tahan, 1988; Bigg et al., 1996) and glacial climate conditions (Death et al., 2005). In addition, these models have been utilized to study the effect of icebergs on the climate during present (e.g. Gladstone et al, 2001; Martin and Adcroft, 2010), pre-industrial (Jongma et al., 2009 Bügelmayer et al., 2014) and past times (Levine and Bigg, 2008; Wiersma and Jongma, 2010; Green et al., 2011; Jongma et al., 2013) using both prescribed and interactively modelled forcing fields, and have shown that icebergs and their melt water have an impact on climate. The spatial distribution of the icebergs' freshwater flux is according to the atmospheric and oceanic forces acting on the icebergs as they determine the icebergs' movement.

Computing iceberg melting and tracks is linked to various types of uncertainties. First, the iceberg's drift and melting, as computed in the iceberg module, are based on empirical parameters and simplifications (e.g. Jongma et al., 2009) that would need further observations to be improved. Second, uncertainties in the reconstructed and modelled wind fields and ocean currents, used to force the icebergs, directly affect the distribution of the freshwater. Third, the initial size distribution of the icebergs is prescribed and based on present day observations (Dowdeswell et al., 1992). Yet, this chosen size distribution may not be a valid representation of calving events in past or future climate conditions.

We therefore propose in this study to extend the approach of Bügelmayer et al. (2014), evaluating in detail the impact of the modelled forcing fields and iceberg size distributions. We use the same earth system model of intermediate complexity (*i*LOVECLIM) coupled to an ice sheet/ice shelf model (GRISLI) and an iceberg module to answer the following research questions.

1. How do atmospheric and oceanic forcing fields affect the icebergs (their lifetime and movement) in the northern hemisphere under pre-industrial conditions?
2. How sensitive is the pre-industrial northern hemisphere climate and Greenland ice sheet to spatial variations in the iceberg melt flux?
3. Do the northern hemisphere climate and the Greenland ice sheet respond differently to icebergs of different initial size distributions?
4. Is the northern hemisphere climate and the Greenland ice sheet response to icebergs of different initial size distribution dependent on the prevailing climate conditions (pre-industrial (PI), warmer than PI and colder than PI?

We will address these questions by presenting results from 15 different sensitivity experiments (Table 1) that differ in the applied forcing (atmospheric, oceanic, pre-industrial, warmer, colder climate) and the initial size distribution (CTRL (standard sizes), BIG, SMALL, Table 2) of the icebergs.

We will first introduce the model and the experimental set-up, then present the results and the discussion, followed by a conclusion section.

2 *Methods*

We use the earth system model of intermediate complexity *i*LOVECLIM (version 1.0) which is a code fork of the LOVECLIM climate model version 1.2 (Goosse et al., 2010). *i*LOVECLIM differs in the ice sheet module included (Roche et al., 2014) and the further developed iceberg module (Bügelmayer et al., 2014), but shares some physical climate components (atmosphere, ocean and vegetation) with LOVECLIM.

2.1 *Atmosphere – Ocean – Vegetation Model*

The climate model *i*LOVECLIM consists of the atmospheric model ECBilt (Opsteegh et al., 1998), a quasi-geostrophic, spectral model with a horizontal resolution of T21 (5.6° in latitude/longitude) and three vertical pressure levels (800, 500, 200hPa). The atmospheric state (including e.g., temperature, humidity) is calculated every four hours. Precipitation depends on the available humidity in the lowermost atmospheric level and the total solid precipitation is given to the ice-sheet model at the end of one model year, as are the monthly surface temperatures.

*i*LOVECLIM includes the sea-ice and ocean model CLIO, which is a 3D ocean general circulation model (Deleersnijder and Campin, 1995; Deleersnijder et al. 1997; Campin and Goosse, 1999) consisting of a dynamic – thermodynamic sea-ice model (Fichefet and Morales Maqueda, 1997, 1999). Due to its free surface, the freshwater fluxes related to iceberg melting can be directly applied to the ocean's surface. The horizontal resolution is 3°x3° in longitude and latitude and the ocean is vertically divided into 20 unevenly spaced layers. CLIO consists of a realistic bathymetry. The oceanic variables (e.g., sea surface temperature and salinity) are computed once a day.

The vegetation (type and cover) is calculated by the vegetation model VECODE (Brovkin et al., 1997), which runs on the same grid as ECBilt. VECODE accounts for fractional use of one grid cell because of the small spatial changes in vegetation. It simulates the dynamics of two plant functional types (trees and gras) as well as bare soil, in response to the temperature and precipitation coming from ECBilt.

The Antarctic ice sheet is prescribed according to present-day conditions following the ETOPO1 topography (http://www.ngdc.noaa.gov/mgg/global/global.html). Icebergs are parameterized in the form of homogenous uptake of latent heat around Antarctica and ice shelf melting is computed according to the prevailing ocean conditions. The Greenland ice sheet is coupled actively using the GRISLI ice-sheet model.

2.2 *GRISLI – Ice Sheet Model*

The ice-sheet model included in *i*LOVECLIM is the Grenoble model for Ice Shelves and Land Ice (GRISLI), which is a three-dimensional thermomechanical model that was first developed for the Antarctic (Ritz et al., 1997, 2001) and was further developed for the northern hemisphere (Peyaud et al., 2007). GRISLI consists of a Lambert azimuthal grid with a 40x40km horizontal resolution. In the present study, it computes the evolution of the thickness and extension of the Greenland ice sheet (GrIS) only, as we do not consider the southern hemisphere grid. GRISLI distinguishes three types of ice flow: inland ice, ice streams and ice shelves. Calving takes place whenever the ice thickness at the border of the ice sheet is less than 150 metres and the points upstream do not provide enough inflow of ice to maintain this thickness. After one model year, the total yearly amount of calving is given to the iceberg module where icebergs are generated daily, as described in detail in Section 2.3. The runoff of GRISLI is calculated at the end of the year by computing the difference between the ice sheet thickness at the beginning of the model year and the end of the year, and taking into account the mass loss due to calving. The runoff is then given to ECBilt where it is re-computed to fit its time-step (4 hours) and incorporated into the land routing system. GRISLI is run for one model year and then provides the runoff and calving, as well as the updated albedo- and topography fields to the atmosphere – ocean – vegetation component. A more detailed explanation of the coupling between ECBilt, CLIO and the ice sheet model GRISLI is provided in Roche et al. (2014) and Bügelmayer et al. (2014).

2.3 *Iceberg Module*

As discussed in detail in Bügelmayer et al. (2014), the dynamic – thermodynamic iceberg module (Jongma et al., 2009; Wiersma and Jongma, 2010) included in *i*LOVECLIM is based on the iceberg-drift model of Smith and co-workers (Smith and Banke, 1983; Smith, 1993; Loset, 1993) and on the developments done by Bigg et al. (1996, 1997) and Gladstone et al. (2001). According to the calving mass and locations calculated by GRISLI over one model year, icebergs of up to ten size classes are generated. The provided ice mass is re-computed to fit the daily time-step of the iceberg module, taking into account the seasonal calving cycle, with the maximum calving occurring from April to

June and the minimum occurring in late summer (Martin and Adcroft, 2010). The control size distribution of the icebergs is according to Bigg et al. (1996) and based on observations of Dowdeswell et al. (1992) that represent the Greenland present day distribution (Table 2). It does not take into account huge tabular icebergs as those calved from Antarctica, but is a valid representation for icebergs calving from the Greenland ice sheet. Therefore, the thickness and width of the calving front as defined in GRISLI affects the amount of ice mass available to generate icebergs, but not the icebergs' dimensions. Icebergs are moved by the Coriolis force, the air- , water- , and sea-ice drag, the horizontal pressure gradient force and the wave radiation force. The forcing fields are provided by ECBilt (winds) and CLIO (ocean currents) and are linearly interpolated from the surrounding grid corners to the icebergs' positions. The icebergs melt over time due to basal melt, lateral melt and wave erosion and may roll over as their length to height ratio changes. The heat needed to melt the bergs is taken from the ocean layers corresponding to the icebergs' depth and the freshwater fluxes are put into the ocean surface layer of the current grid cell. The refreezing of melted water and the break-up of icebergs is not included in the iceberg module.

2.4 *Experimental Set – Up*

We have performed 15 sensitivity experiments that differ in the initial size distribution (CTRL / SMALL / BIG, Table 2), in the applied CO_2 forcing (pre-industrial =280ppm, 4xCO_2 =1120ppm, ¼xCO_2 =70ppm) or in the forces that move the icebergs (atmosphere and ocean). A summary of the experiments performed is given in Table 1. All runs were started from an equilibrated climate and Greenland ice sheet under pre-industrial conditions that has already been used in the study of Bügelmayer et al. (2014). The fact that the initial ice sheet thickness is about $^1/_3$ bigger than the observed one does not impact our results because all the experiments are started from the same ice sheet and climate conditions and thus changes at the end of the model runs are only due to the different forcing fields or iceberg size distribution. The model runs were conducted for 200 model years (pre-industrial) and 1000 model years (4xCO2, ¼xCO_2), respectively. The last 100 years are presented in the results.

2.4.1 *Impact of Forcing Fields*

To differentiate between the impact of the ocean and the atmosphere, the equation of motion (Eq. 1) of an iceberg is used:

$$M\frac{dVi}{dt} = -MfkxVi + Fa + Fr + Fw + Fp + Fs \qquad (1)$$

with M being the Mass of the iceberg, **V** its velocity, the first term ($-Mfkx$ **V**i) on the right side corresponds to the Coriolis force, the second and third are the air drag (Fa and wave radiation force (Fr) and therefore depend on the atmospheric winds; the last three terms represent the oceanic forcing namely water drag (Fw), horizontal pressure gradient(Fp) and sea-ice drag (Fs).

In the so-called "COM" experiments, the icebergs are moved according to Equation 1, thus by the combined atmospheric and oceanic forcing. In the so-called "ATM" set-up (Table 1), all the forcing terms corresponding to ocean currents are set to zero, thereby ensuring that the icebergs are only moved by the Coriolis and the atmospheric forcing. In the "OCE" set-up on the contrary, the air drag and the wave radiation force are defined to be zero, thus only the Coriolis force and the ocean currents are acting on the bergs.

The differentiation between atmospheric and oceanic forces was only made in the equation of motion of an iceberg. The melting of icebergs, which depends on bottom- and lateral melt (oceanic forcing) and the wave erosion (atmospheric forcing), was not altered.

2.4.2 *Initial Size Distribution*

By comparing the CTRL, SMALL and BIG experiments, we are able to investigate the impact of the initial size distribution. In the CTRL experiments, depending on the available mass, icebergs of all 10 size classes can be generated (Bügelmayer et al., 2014). In the SMALL (BIG) experiments, the available mass is used to generate an equal amount of the three smallest (biggest) iceberg sizes (Table 2). The differences in the resulting atmosphere and ocean conditions as well as the ice-sheet allow us to identify the different impact of the BIG and the SMALL icebergs on the climate and the ice sheet. We conducted three sets of experiments using these three size distributions, the first set was done under pre-industrial equilibrium conditions for 200 years. In the second one, a "warm" experiment, we applied a CO_2 concentration four times as strong as the pre-industrial value (1120 vs 280ppm CO_2) and in the third, a "cold" experiment, only a quarter of the pre-industrial CO_2 concentration is used (70 vs 280ppm CO_2). The latter two sets of experiments were done to analyse the effect of the size (CTRL/SMALL/BIG) distribution during periods of a strongly changing ice-sheet under non – equilibrated conditions.

3 Results

3.1 Impact Of Forcing Fields And Initial Iceberg Size On The Transport And Lifetime Of Icebergs (Pre-Industrial)

3.1.1 The Control Experiments (CTRL-COM, CTRL-ATM, CTRL-OCE)

The iceberg distribution of the CTRL-COM experiment displays the general transport of icebergs of all size classes due to atmospheric and oceanic forces (Fig. 1a). We find that most icebergs are transported along the eastern and western coast of Greenland, following the oceanic currents. Further, they are moved southward along the North American coast and spread into the North Atlantic. In the Arctic, most bergs are found close to Ellesmere Island, due to the calving sites in this region (not shown) and are then widely distributed by the Beaufort Gyre and the prevailing winds.

By applying only atmospheric forcing, we find that CTRL-ATM icebergs are transported further into the North Atlantic and Arctic Ocean (Fig. 1d) than seen in CTRL-COM. After calving, they are quickly pushed away from the Greenland ice sheet (GrIS) margin. In CTRL-ATM less bergs than in CTRL-COM move along the coast of Greenland as can be seen in the number of bergs travelling along the coast (Fig. 1d, f), highlighting the lack of ocean currents. Overall, the amount of iceberg melt flux released in CTRL-ATM (mid- to high latitudes: 150 m^3/s) is of the same magnitude as in CTRL-COM and over the same area (Fig. 2a). Yet, the lifetime of CTRL-ATM icebergs, that is the time (in months) it takes to completely melt the bergs, is up to one year shorter than in CTRL-COM (Fig. 3) because they are transported faster away from the ice sheet and into warmer conditions.

The effect of the oceanic forcing is in strong contrast to the atmospheric one as it causes the CTRL-OCE icebergs to stay closer to the GrIS margin (Fig. 1g). The icebergs movement reflects the prevailing ocean currents, mainly the Beaufort Gyre, the East Greenland and the Labrador Current. Much less icebergs are moved from the ice sheet into the Greenland – Iceland – Norwegian (GIN) Seas and the North Atlantic in CTRL-OCE compared to CTRL-COM (Fig. 1a,g) due to the lack of wind forcing, which is also reflected in the area that they cover (Fig. 2c,d). Also In the Arctic Ocean the CTRL-OCE icebergs do not spread as much, but a slightly larger iceberg melt flux (IMF) is released because the bergs are not transported southwards by the wind, but stay and melt in there. Overall, the amount of freshwater flux is comparable to the CTRL-COM experiment, though over a much smaller area (CTRL-COM: 1.4×10^{13} m^2, CTRL-OCE: 0.8×10^{13} m^2, Fig. 2a) and over a longer time period.

The CTRL-OCE icebergs melt up to 4 months slower than CTRL-COM bergs because they stay close to the GrIS margin and thus in colder water (Fig. 3).

3.1.2 *The BIG Experiments (BIG-COM, BIG-ATM, BIG-OCE)*

The spatial distribution of the BIG-COM icebergs displays, first, the effect of the Coriolis force since there is an eastward movement in the North Atlantic (Fig. 1b). The Coriolis force depends on the size and velocity of the icebergs and thus, is acting stronger on big icebergs than on small ones. Second, the area covered by BIG-COM bergs is larger in the North Atlantic than in CTRL-COM (Fig. 2d). Over the mid-to high latitudes the area covered by more than 10 BIG-COM icebergs is only slightly bigger than the one of CTRL-COM (Fig.2a), even though their lifetime is up to three years longer (Fig. 3). But in total there are less BIG bergs generated than in the CTRL experiment because more mass is needed per berg (Table 2).

Applying only wind forcing on the BIG icebergs (BIG-ATM) transports less icebergs into the North Atlantic and especially the GIN Seas (Fig. 1e) where they cover about half the area of BIG-COM ($4x10^{12}m^2$ compared to $7x10^{12}m^2$), but release the same amount of freshwater (150 m^3/s, Fig. 2c). The BIG-ATM icebergs are not transported as far as the BIG-COM bergs in all the regions considered and especially in the GIN Seas (Fig. 2c). There, the BIG-ATM bergs follow the strong southward component of the wind without being distributed further into the GIN Seas. Similar to the CTRL experiment, the BIG-ATM icebergs melt up to two years faster than the ones of BIG-COM or BIG-OCE (Fig. 3).

The impact of oceanic forcing on the iceberg distribution is simulated in BIG-OCE. Since the big icebergs melt slowly, they are transported further south than CTRL-OCE bergs (Fig. 1h). In the GIN Seas the BIG-OCE bergs are spread from the coast and cover almost the same area as the BIG-ATM (Fig. 2c). In the Arctic Ocean the BIG-OCE icebergs release a higher averaged melt flux than BIG-COM and BIG-ATM ($125m^3/s$ compared to $75m^3/s$ and $95m^3/s$, respectively; Fig. 2b), but over a smaller area. This is because of the missing wind forcing which prevents the icebergs from being distributed out of the Arctic Ocean, instead the bergs are stuck close to their calving sites. The higher IMF in BIG-OCE does not strongly impact the Arctic climate because of the prevailing cold conditions. Thus, more IMF, which is released to the ocean surface layer at 0°C and consequently cools and freshens it, does not cause noticeable changes. The area covered by BIG bergs over the mid-to high latitudes is clearly bigger than SMALL-, or CTRL-OCE (Fig. 2a) because of their lifetime, which is about two years longer compared to CTRL-OCE (Fig, 3).

3.1.3 The SMALL Experiments (SMALL-COM, SMALL -ATM, SMALL -OCE)

Generating only SMALL-COM icebergs results in a similar iceberg distribution as in CTRL-COM (Fig. 1c), but less widespread. The amount of freshwater that is released by SMALL-COM bergs is almost the same over the mid- to high latitudes as CTRL-COM, but over a smaller area (Fig. 2a) because all the SMALL-COM icebergs are melted within two years, compared to three years in CTRL-COM (Fig. 3).

In the icebergs' distribution of the SMALL-ATM model runs (Fig. 1f), it is clearly visible that the light, small bergs are easily pushed away from their calving sites by the atmospheric forcing, but as in the COM experiments, over a smaller area because they melt faster. In the North Atlantic, the general pattern is directed westward, in contrast to BIG-ATM icebergs that are strongly influenced by the Coriolis force.

The wide-spread distribution of SMALL-ATM is in strong contrast to the one of SMALL-OCE (Fig. 1i). The oceanic forcing restricts the icebergs' transport to the shore and due to their smaller size SMALL-OCE bergs melt before being distributed as far as CTRL-OCE and especially BIG-OCE (Fig 2a).

In short, the impact of the forcing fields is clearly seen in the icebergs' distribution and especially lifetime since 90% of all the atmospheric forced icebergs (SMALL-, BIG-, and CTRL-ATM) melt up to two years faster compared to the oceanic forced bergs and compared to the icebergs of the SMALL-, BIG-, and CTRL-COM set-up.

3.2 *Impact Of Forcing Fields And Initial Iceberg Size On Pre-Industrial Climate*

The resulting sea surface and air temperatures (SST, TAIR) are comparable between the CTRL-COM,-ATM, and –OCE experiments (Fig. 4a,b), despite the different spatial distribution of the iceberg melt flux . The biggest spread in IMF is found in the Arctic Ocean (BIG-COM: 75m^3/s, CTRL-OCE: 150m^3/s, Fig. 2b), but these differences do not result in an altered climate state due to the prevailing cold conditions that are less sensitive to the freshening and cooling effect of icebergs (not shown). Also in the GIN Seas and North Atlantic the SST and TAIR do not significantly differ between the experiments, even though these are sensitive areas because of the located convection sites. This indicates that since the amount of freshwater released is comparable in the model runs, the exact location of the release does not have a strong impact on the prevailing climate conditions. Further, the shorter lifetime of the atmospheric driven icebergs does not cause differences in the resulting climate and the GrIS because the calving flux provided by GRISLI is almost constant over the years and comparable in all the pre-industrial experiments. Therefore, the same amount of freshwater is supplied to the ocean. Under pre-industrial equilibrium conditions the atmospheric and oceanic forcing do transport the icebergs differently, but the resulting spatial patterns of the iceberg melt flux cause only local differences in the Greenland ice sheet volume (Table 3), the oceanic and atmospheric conditions.

3.3 *Impact Of Initial Iceberg Size Under A Changing Climate*

To have more confidence in using the present day iceberg distribution also for simulations of past and future climates, we conducted two more sets of experiments with enhanced or reduced radiative forcing to obtain warmer and colder climate states. This change in radiative forcing was applied through adjustment of the atmospheric CO_2 concentration in two experiments, the so-called HIGH = 4xCO_2 (1120ppm) and LOW= ¼xCO_2 (70ppm), with a duration of 1000 years. For each of these settings, we performed experiments with CTRL, BIG and SMALL icebergs. The HIGH experiments resulted in an up to 4°C warmer global mean temperature and caused the Greenland ice sheet to lose 10% of its volume, whereas the LOW experiments caused the mean global temperatures to decrease about 4°C and an increase of the Greenland ice sheet volume of up to 4%, compared to the pre-industrial ice sheet volume (Table 3).

3.3.1 *Experiments With High Radiative Forcing*

The impact of the enhanced radiative forcing on the Greenland ice sheet is displayed in Fig. 5, where the resulting CTRL-HIGH ice sheet extensions and thickness are shown compared to the equilibrated CTRL-COM ice sheet (Fig. 5a,b).

As the ice sheet is shrinking and retreating from the coast (Fig. 5b), the amount of calving flux from the GrIS is decaying (0.003 Sv vs 0.02 Sv in the CTRL-COM), especially in South Greenland, and so is the icebergs melt flux. The released iceberg melt flux in the GIN Seas is in the range of 20 (SMALL-, CTRL-HIGH) to 50m^3/s (BIG-HIGH, Fig. 2c), compared to 150m^3/s in the CTRL-COM. Moreover, there are hardly any icebergs entering the North Atlantic, independent of the used size distribution (Fig. 2d). In the Arctic Ocean the HIGH experiments result in a bigger spread between the CTRL, BIG and SMALL runs than any other performed set-up. The BIG-HIGH bergs cover the smallest area because of the decreased calving flux much less BIG bergs are generated. Further, there are still SMALL bergs, but due to their size and the warmer conditions they melt faster than seen in the SMALL experiments performed under pre-industrial conditions. The CTRL-HIGH experiment covers a slightly smaller area than the CTRL-COM,-OCE or –ATM, but much bigger than BIG-, and SMALL-HIGH. This is because the different iceberg sizes allow for the production of a higher number of icebergs than in BIG and the existence of icebergs bigger than size 3 (as in SMALL) allows for a longer lifetime.

Although the size of the icebergs generated varies from the beginning, the resulting climate conditions, such as sea surface or air temperatures do not vary at the end of the 1000 year period between the SMALL-, BIG-, and CTRL-HIGH experiments (Fig. 4a, b), nor does the GrIS volume (Table 3). During periods of strong background changes, different iceberg distributions do not result in different climate states. This indicates that the applied forcing has a stronger impact than local differences due to the chosen iceberg size.

3.3.2 *Experiments With Low Radiative Forcing*

In contrast to the experiments with high radiative forcing, the low radiative forcing causes up to 4°C lower global mean temperatures and consequently the ice sheet’s volume is thickening and extending further down to the coast line (Fig. 5c). Similar to the other experiments performed, the impact of different initial size distributions of the icebergs is negligible on the resulting climate and ice sheet volume (Table 3).

Due to the increased ice sheet thickness, more calving flux is released (0.05 Sv in CTRL-LOW compared to 0.02 Sv in CTRl-COM) and so the iceberg melt flux increases to 300m^3/s in the mid-to high latitudes, compared to 150m^3/s in the pre-industrial experiments. The increase is seen almost everywhere around Greenland (Fig. 2a,c,d), except in the Arctic Ocean (Fig. 2b). In the Arctic Ocean the released IMF is in the same range as in the experiments performed under pre-industrial conditions because the ice sheet’s thickness and consequently the calving sites in North Greenland are not strongly altered by the colder climate (Fig. 5c). In the North Atlantic the released iceberg melt flux displays a big spread between the experiments with the BIG-LOW bergs being spread the furthest and releasing the most IMF (80m^3/s in BIG-LOW vs 45m^3/s in CTRL-LOW; Fig. 2b). Since the cold conditions prevent the BIG-LOW icebergs from melting quickly, almost all of them are transported into the North Atlantic where they finally melt. This is also partly the case for the CTRL-LOW bergs thereby resulting in a higher iceberg melt flux than the SMALL-LOW (Fig. 2b). Independent of the chosen size distribution, the resulting temperatures are about 5°C lower than during pre-industrial conditions in the North Atlantic and the GIN Seas (Fig. 4), displaying the strong CO_2 forcing.

These results show that the used initial size distributions do not alter the response of the climate and the GrIS to the applied forcing. This thus indicates that the extreme boundary conditions have a stronger impact on the results than the used iceberg sizes.

4 *Discussion*

By testing the impact of the atmospheric versus the oceanic forcing on icebergs' lifetime and movement, we find that the atmospheric forcing causes the bergs to travel further away from their calving sites and into the North Atlantic, whereas the ocean currents lead to iceberg tracks closer to shore. It is difficult to compare our results to previous studies, since the studies that investigated the impact of the background forcing (Smith 1993; Keghouche et al., 2002) focused on observations of single icebergs and the ability of models reproducing their specific tracks. Bigg et al. (1997) noted that the modelling of specific iceberg tracks is very unlikely to be successful and it is important to notice that we do not expect our model to resolve single tracks due to its coarse resolution, but to reflect the wide spread effect of icebergs on climate.

In our model, the impact of icebergs on climate does not strongly depend on the two types of forcing (atmospheric and oceanic), yet their lifetime is shortened by up to two years when they are transported by atmospheric forces only. Bigg et al. (1997) showed that about 80% of the small bergs (size class 1 to 3, Table 2) melt within the first year, which is higher than in our SMALL-COM set-up where about 60% are melted. Also Venkatesh and El-Tahan (1988) conducted a study to investigate the impact of modelling complete deterioration of icebergs on the prediction of their tracks. In their study they showed that most of the icebergs corresponding to size class 1 to 3 used in this study, disappear within 3 to 22 months, reassuring our results. The maximum lifetime of the BIG bergs is found to be almost seven years, which is slightly longer than modelled by Bigg et al. (1997). This discrepancy can be due to the pre-industrial climate conditions used in our study that are slightly colder than the present day conditions applied by Bigg et al. (1997).

To better understand the response of the modelled climate to the initial size distribution, we performed different sensitivity experiments. First, using pre-industrial conditions we find that independent of the forcing, SMALL icebergs release less freshwater and spread over a smaller area than BIG and CTRL icebergs. In the North Atlantic the impact of the Coriolis force is especially pronounced in the BIG-ATM and BIG-COM runs, confirming the findings of Roberts et al. (2014). In their study they noted that BIG icebergs travel further south than small icebergs due to the stronger impact of the Coriolis force. Even though the SMALL icebergs cause locally different ocean and atmospheric conditions than the BIG bergs, the overall effect on climate and especially on the Greenland ice sheet is negligible.

Second, we repeated the experiments under a strongly increased and decreased radiative forcing for 1000 years. During this time scale changes in the Southern Ocean can impact the Northern Hemisphere. Jongma et al. (2009) showed that including active icebergs increases the net production of Antarctic Bottom Water by 10% under pre-industrial conditions. We do neglect this direct effect of icebergs here since icebergs and Antarctic ice-sheet runoff are computed using parameterizations that depend on the prevailing climate conditions. But we do not expect that the size of the icebergs released from Greenland have an impact on the Southern Hemisphere, thus, the uncertainty introduced by not actively coupling the Antarctic ice sheet is present and comparable in all the radiative forcing experiments.

There might be different reasons why the climate conditions and the GrIS are not strongly affected by the initial size distribution during strong radiative background conditions. One reason could be that the ice sheet and the climate model are too insensitive to the experienced changes as they have a relatively coarse resolution. Therefore, it would be interesting to repeat this study with a finer model grid. Another reason might be that in the experiments where really strong forcing was applied (HIGH=1120ppm CO_2, LOW= 70ppm CO_2), the feedbacks related to calving have a smaller signal than the forcing and are therefore overruled.

5 *Conclusions*

Within a fully coupled climate – ice sheet – iceberg model set up, we have performed sensitivity experiments to investigate the effect of the forcing fields such as winds and ocean currents, as well as the prescribed initial size distribution on the icebergs and the climate.

We find that, under pre-industrial conditions, the wind forcing pushes the icebergs further away from their calving sites and further into the North Atlantic, whereas the ocean currents transport the bergs close to Greenland and southward along the North American coast. The combined effect of the forces (control set-up) displays a lesser spread iceberg distribution in the Arctic Ocean and into the North Atlantic than the purely atmospheric driven bergs due to the restrictive effects of the oceanic forcing. The icebergs' spread depends on both the forcing fields and the icebergs size with the CTRL bergs being transported the furthest, followed by the BIG bergs (size class 8 to 10). The amount of released iceberg melt flux is comparable in all the experiments, though locally different. In our model set-up, the biggest impact of the applied forcing (atmospheric or oceanic) is on the icebergs' lifetime which is up to two years shorter if the icebergs are only transported by winds.

In the presented model framework, the implementation of icebergs of different size classes under equilibrated pre-industrial conditions reveals that there are local differences in the released freshwater flux. However, these differences do not cause significant changes in the resulting Greenland ice sheet volume and climate conditions.

When repeating the experiments with different size distributions with strong radiative cooling or warming (1120 ppm CO_2 or 70 ppm CO_2, 1000 model years), the response of the climate and the ice sheet volume are almost identical in all the performed experiments.

Even though the iceberg and freshwater distribution differ between the conducted experiments (all size classes, only SMALL and only BIG bergs, respectively), their impact on the northern hemispheric climate does not differ strongly. We can therefore conclude that for the resulting climate and ice sheet small spatial differences between the runs do not have a strong impact as long as there is a wide spread impact of icebergs (cooling and freshening) around Greenland. Furthermore, our results show that the response of the climate to the applied radiative forcing is much stronger than its response to the used initial size distribution of the icebergs.

The presented results make us confident in applying the prescribed present day iceberg sizes under different climates without introducing a strong bias.

Code availability

The iLOVECLIM source code is based on the LOVECLIM model version 1.2 whose code is accessible at http://www.elic.ucl.ac.be/modx/elic/index.php?id=289. The developments on the iLOVECLIM source code are hosted at https://forge.ipsl.jussieu.fr/ludus, but are not publicly available due to copyright restrictions. Access can be granted on demand by request to D. M. Roche (didier.roche@lsce.ipsl.fr). The specific experimental set-up used for this study is available at https://forge.ipsl.jussieu.fr/ludus.

Acknowledgements. M. Bügelmayer is supported by NWO through the VIDI/AC2ME project no 864.09.013. D. M. Roche is supported by NWO through the VIDI/AC2ME project no 864.09.013 and by CNRS-INSU. The authors wish to thank Catherine Ritz for the use of the GRISLI ice sheet model. Institut Pierre Simon Laplace is gratefully acknowledged for hosting the *i*LOVECLIM model code under the LUDUS framework project (https://forge.ipsl.jussieu.fr/ludus). This is NWO/AC2ME contribution number 08.

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LIST OF TABLES

	PRE-INDUSTRIAL (ATM & OCE FORCING) **= 280 ppm**	**ONLY ATMOSPHERIC FORCING**	**ONLY OCEANIC FORCING**	**4xCO$_2$** (ATM & OCE FORCING) **= 1120ppm**	**¼xCO$_2$** (ATM & OCE FORCING) **= 70ppm**
ALL SIZES	CTRL-COM	CTRL-ATM	CTRL-OCE	CTRL-HIGH	CTRL-LOW
BIG BERGS	BIG-COM	BIG-ATM	BIG-OCE	BIG-HIGH	BIG-LOW
SMALL BERGS	SMALL-COM	SMALL-ATM	SMALL-OCE	SMALL-HIGH	SMALL-LOW

Table 1: performed experiments

CLASS	**HEIGHT (m)**	**WIDTH (m)**	**VOLUME (m^3)**	**PERCENTAGE** of total available Volume	**EXPERIMENT**
1	67	67	5.16E05	0.15 / 0.33	CTRL / SMALL
2	133	133	4.07E06	0.15 / 0.33	CTRL / SMALL
3	200	200	1.38E07	0.2 / 0.33	CTRL / SMALL
4	267	267	3.28E07	0.15	CTRL
5	300	333	5.74E07	0.08	CTRL
6	300	400	8.28E07	0.07	CTRL
7	300	500	1.29E04	0.05	CTRL
8	300	600	1.86E08	0.05 / 0.33	CTRL / BIG
9	300	800	3.31E08	0.05 / 0.33	CTRL / BIG
10	300	1000	5.18E08	0.05 / 0.33	CTRL / BIG

Table 2: used initial iceberg classes

Experiment	**Mean**	**STDEV**	**% diff**
CTRL-COM	3,90E+15	2,53E+12	-
BIG-COM	3,91E+15	2,61E+12	-0,09
SMALL-COM	3,91E+15	1,96E+12	-0,08
CTRL-ATM	3,91E+15	1,90E+12	-
BIG-ATM	3,91E+15	2,14E+12	0,02

SMALL-ATM	3,91E+15	1,99E+12	-0,06
CTRL-OCE	3,91E+15	2,11E+12	-
BIG-OCE	3,91E+15	1,29E+12	-0,03
SMALL-OC	3,91E+15	2,20E+12	-0,14
CTRL-HIGH	3,50E+15	5,03E+12	-
BIG-HIGH	3,49E+15	4,40E+12	0,32
SMALL-HIGH	3,49E+15	5,69E+12	0,14
CTRL-LOW	4,04E+15	1,90E+12	-
BIG-LOW	4,06E+15	2,74E+12	-0,41
SMALL-LOW	4,04E+15	3,20E+12	-0,05

Table 3: Ice-sheet Volume (m^3): Mean and Standard deviation of last 100 years, % diff = difference between the ice sheet volume of the CTRL experiment and the BIG/SMALL experiments in percent

FIGURE CAPTIONS

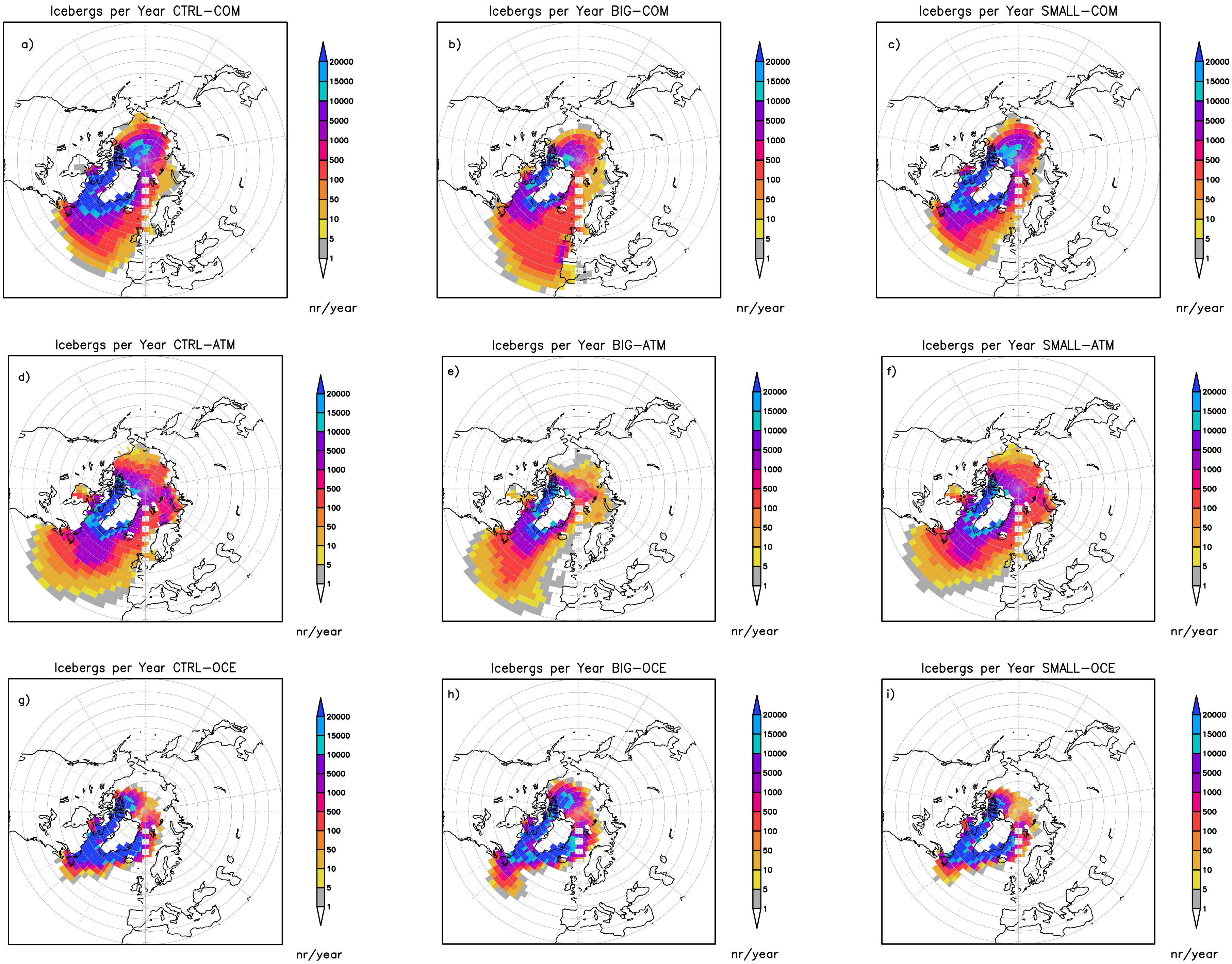
Figure 1: Number of icebergs passing by a grid cell per year (icebergs that are grounded are only counted once); first row: atmospheric forcing only (CTRL-, BIG-, SMALL-ATM); second row: oceanic forcing only (CTRL-, BIG-, SMALL-OCE), third row: the default set-up (icebergs are moved by both, atmospheric and oceanic forcing; CTRL-, BIG-, SMALL-COM)

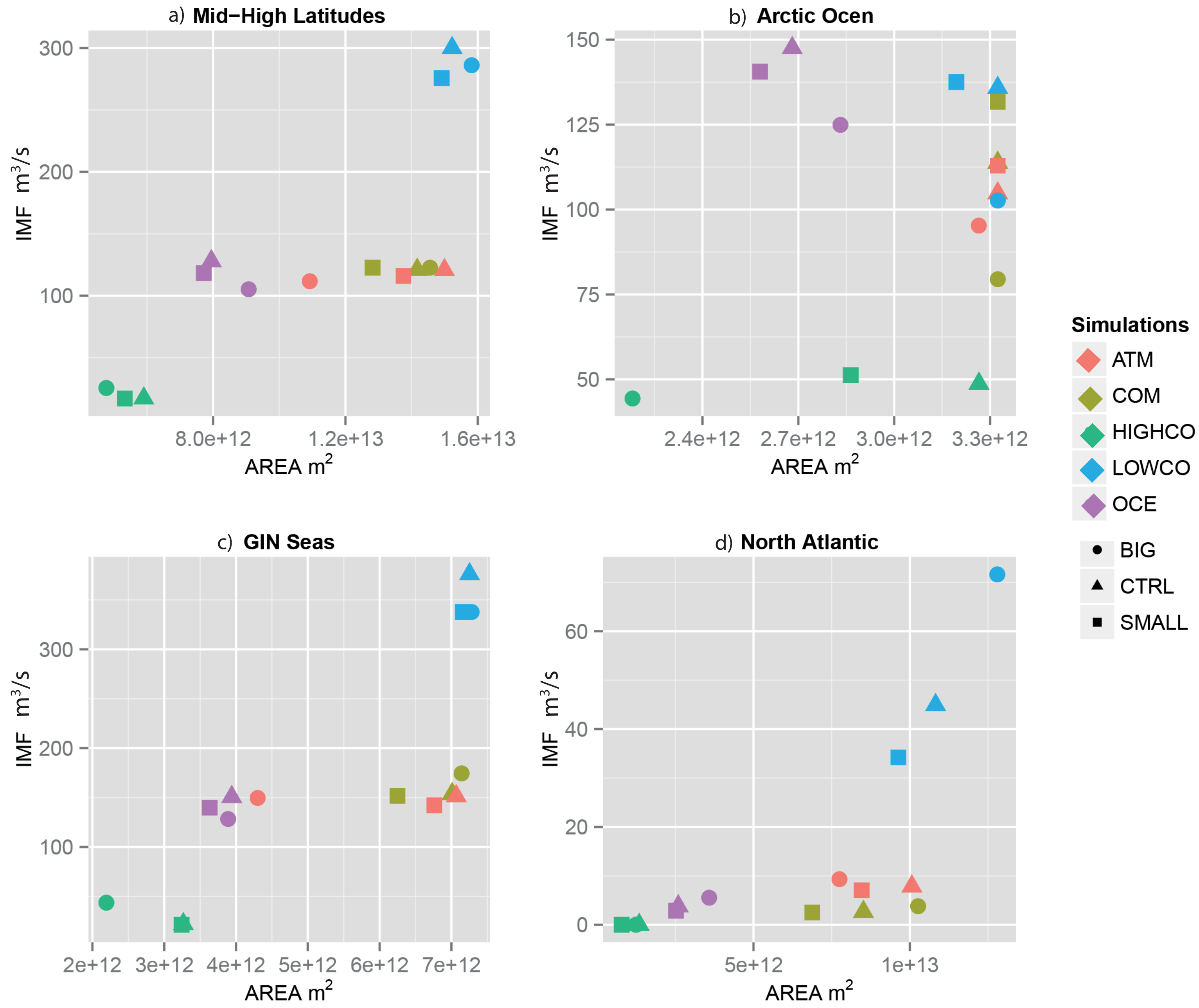
Figure 2: area (m^2) vs iceberg melt flux (m^3/s); the area is computed by taking into account all the gridcells that are passed by more than 10 icebergs (be aware that the area is 10^{13}m^2 in a), 10^{12} m^2 otherwise); a: Mid- to High Latitudes: mean computed over 40-90°N and 80°W-15°E, values of IMF: 30-180 m3/s; b: Arctic Ocean: 80-90°N and 180°W-180°E, values of IMF: 60-140 m3/s; c: Greenland – Iceland – Norwegian (GIN) Seas: 50-85°N and 45°W-15°E, values of IMF: 40-240 m3/s; d: North Atlantic: 45-60°N and 60-20°W, values of IMF: 0-50 m3/s;

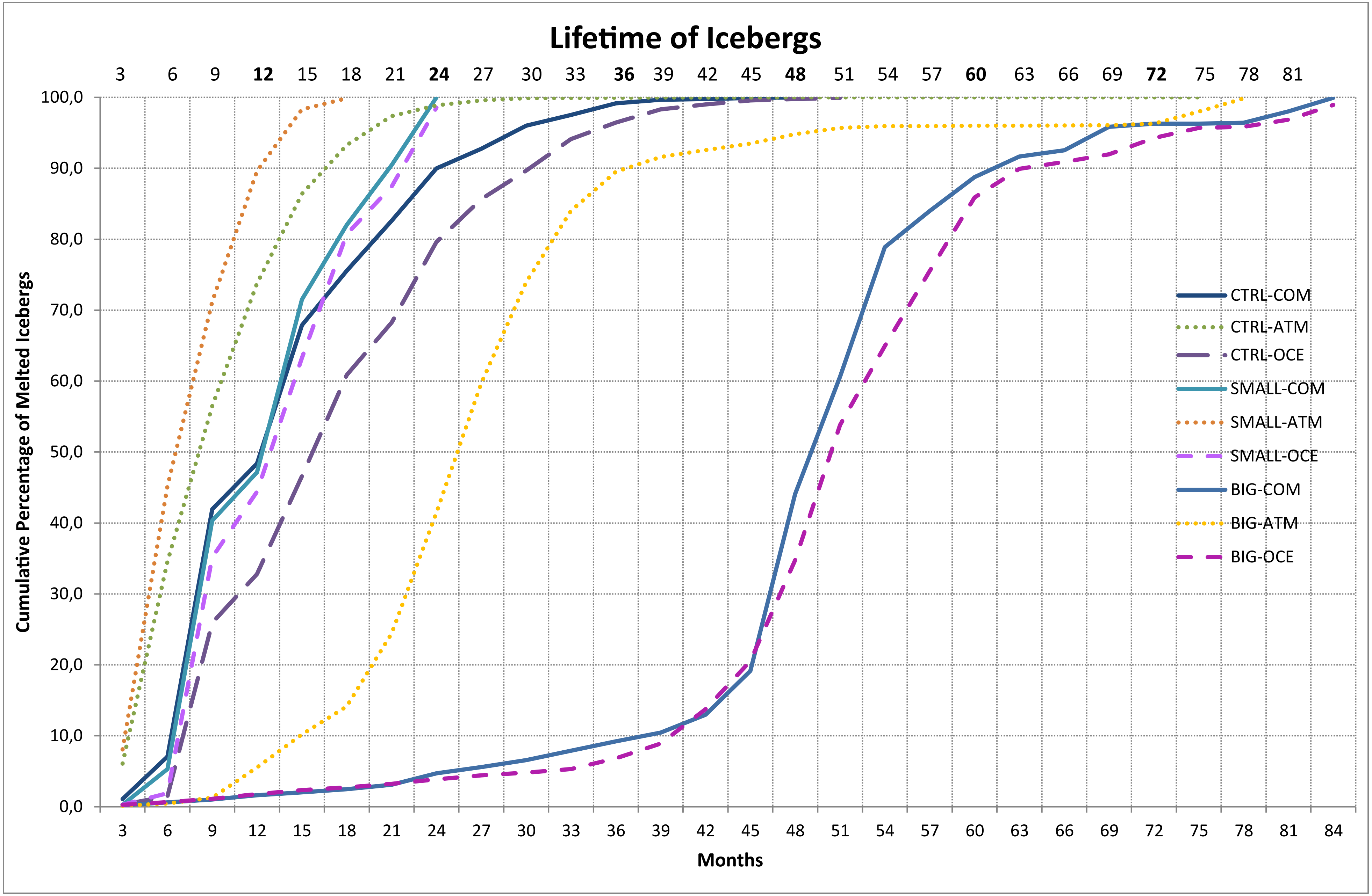
Figure 3: cumulative percentage of icebergs melted within a certain time; x – Axis corresponds to months, y-axis to cumulative percentage

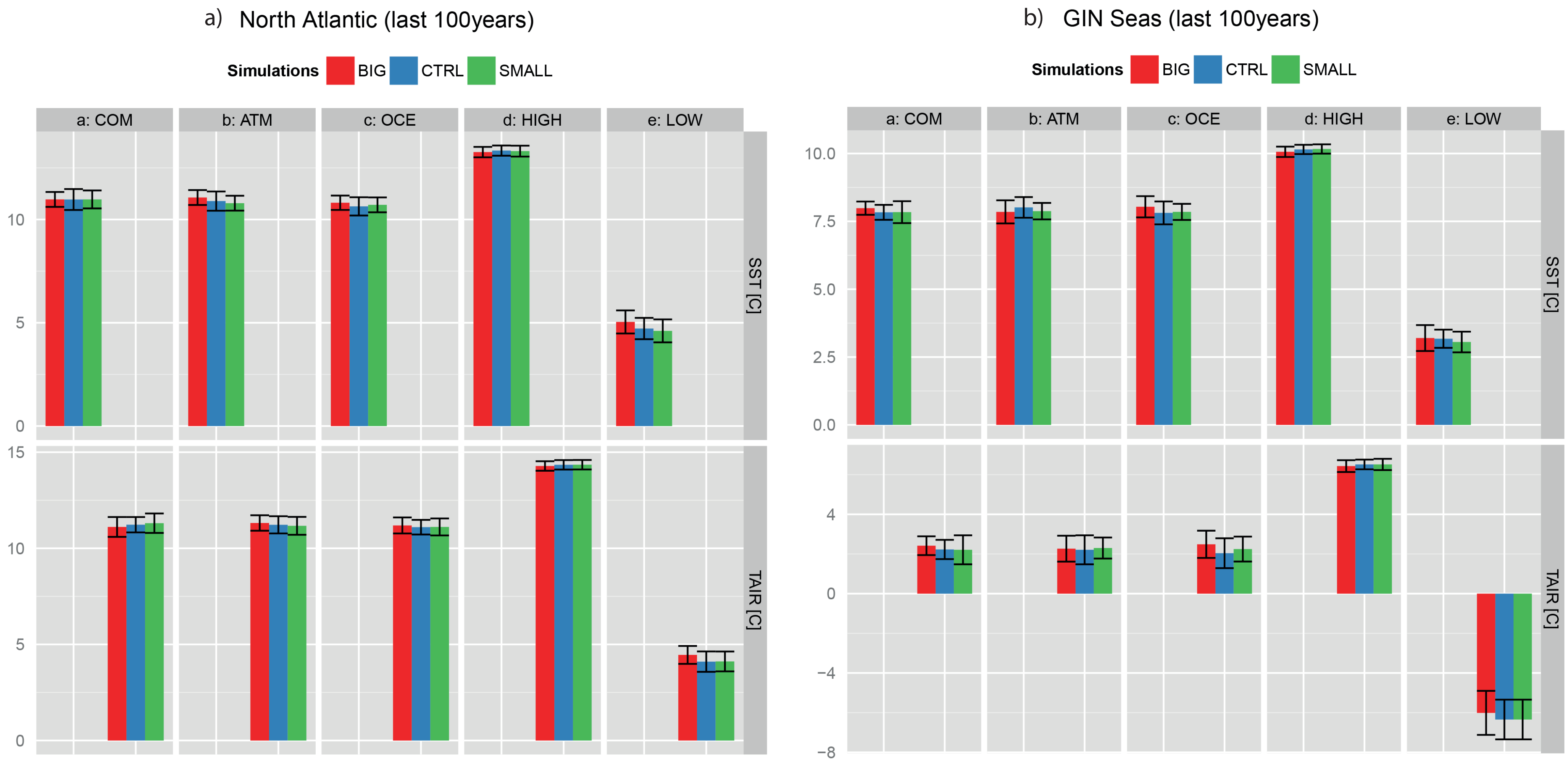
Figure 4: Mean + Standard deviation of last 100 years of the performed experiments: Sea Surface Temperature (SST, °C) and air temperature (TAIR, °C); red = BIG bergs, blue = CTRL, green = SMALL bergs; a: North Atlantic: mean computed over: 45-60°N and 60-20°W; b: Greenland – Iceland – Norwegian (GIN) Seas: 50-85°N and 45°W-15°E

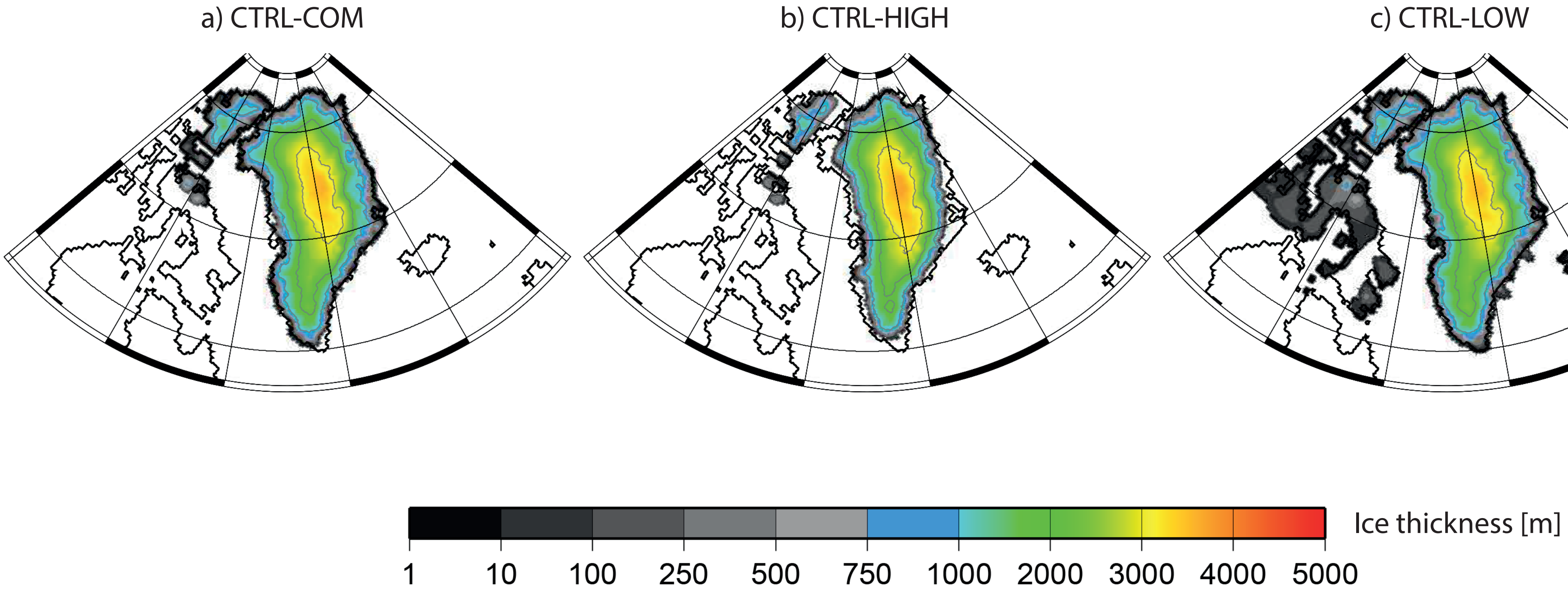
Figure 5: ice sheet thickness at the end of the experiments (m); a: CTRL-COM; b: CTRL-HIGH; c: CTRL-LOW


Icebergs per Year CTRL–COM
a)
Icebergs per Year BIG–COM
b)
Icebergs per Year SMALL–COM
c)
Icebergs per Year CTRL–ATM
d)
Icebergs per Year BIG–ATM
e)
Icebergs per Year SMALL–ATM
f)
Icebergs per Year CTRL–OCE
g)
Icebergs per Year BIG–OCE
h)
Icebergs per Year SMALL–OCE
i)
20000
15000
10000
5000
1000
500
100
50
10
5
1
nr/year


a) Mid−High Latitudes
IMF m³/s
300
200
100
8.0e+12
1.2e+13
1.6e+13
AREA m²
b) Arctic Ocen
IMF m³/s
150
125
100
75
50
2.4e+12
2.7e+12
3.0e+12
3.3e+12
AREA m²
c) GIN Seas
IMF m³/s
300
200
100
2e+12
3e+12
4e+12
5e+12
6e+12
7e+12
AREA m²
d) North Atlantic
IMF m³/s
60
40
20
0
5e+12
1e+13
AREA m²
Simulations
ATM
COM
HIGHCO
LOWCO
OCE
BIG
CTRL
SMALL


Lifetime of Icebergs
3
6
9
12
15
18
21
24
27
30
33
36
39
42
45
48
51
54
57
60
63
66
69
72
75
78
81
100,0
90,0
80,0
70,0
60,0
50,0
40,0
30,0
20,0
10,0
0,0
Cumulative Percentage of Melted Icebergs
3
6
9
12
15
18
21
24
27
30
33
36
39
42
45
48
51
54
57
60
63
66
69
72
75
78
81
84
Months
CTRL-COM
CTRL-ATM
CTRL-OCE
SMALL-COM
SMALL-ATM
SMALL-OCE
BIG-COM
BIG-ATM
BIG-OCE


a) North Atlantic (last 100years)
Simulations
BIG
CTRL
SMALL
a: COM
b: ATM
c: OCE
d: HIGH
e: LOW
SST [C]
TAIR [C]
0
5
10
15
b) GIN Seas (last 100years)
Simulations
BIG
CTRL
SMALL
a: COM
b: ATM
c: OCE
d: HIGH
e: LOW
SST [C]
TAIR [C]
0.0
2.5
5.0
7.5
10.0
−8
−4
0
4


a) CTRL-COM
b) CTRL-HIGH
c) CTRL-LOW
1
10
100
250
500
750
1000
2000
3000
4000
5000
Ice thickness [m]