

On the Computation of Planetary Boundary Layer Height using the Bulk Richardson Number Method

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Abstract

Experimental data from four field campaigns are used to explore the variability of the bulk Richardson number of the entire planetary boundary layer (PBL), Ri_{bc} , which is a key parameter for calculating the PBL height (PBLH) in numerical weather and climate models with the bulk Richardson number method. First, the PBLHs of three different thermally-stratified boundary layers (i.e., strongly stable boundary layers, weakly stable boundary layers, and unstable boundary layers) from the four field campaigns are determined using the turbulence method, the potential temperature gradient method, the low-level jet method, or the modified parcel method. Then for each type of boundary layers, an optimal Ri_{bc} is obtained through linear fitting and statistical error minimization methods so that the bulk Richardson method with this optimal Ri_{bc} yields similar estimates of PBLHs as the methods mentioned above. We find that the optimal Ri_{bc} increases as the PBL becomes more unstable: 0.24 for strongly stable boundary layers, 0.31 for weakly stable boundary layers, and 0.39 for unstable boundary layers. Compared with previous schemes that use a single value of Ri_{bc} in calculating the PBLH for all types of boundary layers, the new values of Ri_{bc} proposed by this study yield more accurate estimates of PBLHs.

Keywords: planetary boundary layer height, bulk Richardson number, sounding, numerical weather and climate model

1 Introduction

The planetary boundary layer (PBL), or the atmospheric boundary layer, is the lowest part of the atmosphere that is directly influenced by the Earth's surface and has significant impacts on weather, climate, and the hydrologic cycle (Stull, 1988; Garratt, 1992; Seidel et al., 2010). The

1 height of the PBL (PBLH) is typically on the order of 1~2 km but varies significantly during a
2 diurnal cycle in response to changes in the thermal stratification of the PBL. It is an important
3 parameter that is commonly used in modeling turbulent mixing, atmospheric dispersion,
4 convective transport, and cloud/aerosol entrainment (Deardorff, 1972; Holtslag and Nieuwstadt,
5 1986; Sugiyama and Nasstrom, 1999; Seibert et al., 2000; Medeiros et al., 2005; Konor et al.,
6 2009; Liu and Liang, 2010; Leventidou et al., 2013). As a result, accurate estimates of the PBLH
7 under different thermal stratifications are critically needed.

8 The PBL is characterized by the presence of continuous turbulence, while turbulence is
9 lacking or sporadic above the PBL. Therefore, the PBLH can be viewed as the level where
10 continuous turbulence stops (Wang et al, 1999; Seibert et al., 2000). Using high-frequency
11 turbulence measurements (e.g., collected from ultrasonic anemometers on aircrafts), the PBLH
12 can be readily determined. This is known as the turbulence (Tur) method. It is highly reliable, but
13 the instruments required by this method are costly. A more economic option is to determine the
14 PBLH through analyzing temperature and wind profiles measured from radio soundings. In this
15 method, the PBLs are broadly classified as strongly stable boundary layers (Type I SBLs), weakly
16 stable boundary layers (Type II SBLs), or unstable boundary layers (UBLs) (Holtslag and Boville,
17 1993; Vogelezang and Holtslag, 1996). They are defined using the surface heat flux and the
18 potential temperature profile, as shall be seen later.

19 For strongly stable boundary layers or Type I SBLs, there is a strong inversion in the potential
20 temperature profile and the PBLH is usually defined as the top of the inversion where the potential
21 temperature gradient (PTG) first becomes smaller than a certain threshold γ_s (Bradley et al., 1993),
22 which is chosen to be $6.5 \text{ K (100m)}^{-1}$ following Dai et al. (2011). This is called the PTG method
23 hereafter. For weakly stable boundary layers or Type II SBLs, turbulence is generated from wind
24 shear due to relatively high wind speed and the PBLH is defined as the height of the low-level jet
25 (LLJ) (Melgarejo and Deardorff, 1974). This is called the LLJ method hereafter. For unstable
26 boundary layers or UBLs, buoyancy is the dominant mechanism driving turbulence, and the
27 PBLH is defined as the height at which a thin layer of capping inversion occurs. The PBLH of
28 UBLs is determined first by identifying a height at which a parcel of dry air, released adiabatically
29 from the surface, reaches equilibrium with its environment (Holzworth, 1964). This height is then
30 corrected by another upward search for another height at which the potential temperature gradient
31 first exceeds a threshold γ_c (Liu and Liang, 2010), which is chosen to be $0.5 \text{ K (100 m)}^{-1}$ in this
32 study. This is called the modified parcel method hereafter.

33 For an atmosphere with discernible characteristics (i.e., a strongly stable potential temperature
34 profile for the Type I SBL, a strong LLJ for the Type II SBL, and a capping inversion layer for the
35 UBL), the three methods generally show good performances (e.g., Mahrt et al., 1979; Liu and
36 Liang, 2010; Dai et al., 2011). However, for an atmosphere without these discernible
37 characteristics, large errors can be introduced by these methods. As such, these methods are
38 usually used in experimental studies but not in numerical models since numerical models need to

1 determine the PBLH automatically. Instead, the bulk Richardson number (Ri_b) method is often
 2 used for numerical weather and climate models due to its reliability under a variety of atmospheric
 3 conditions (e.g., Holtslag and Boville, 1993; Jericevic and Grisogono, 2006; Richardson et al.,
 4 2013). The bulk Richardson number method assumes that the PBLH is the height at which the Ri_b
 5 reaches a threshold value (Ri_{bc} , which is called ‘the bulk Richardson number for the entire PBL’
 6 hereafter). The Ri_b at a certain height z is calculated with the potential temperature and wind speed
 7 at this level and those at the lower boundary of the PBL (generally the surface), as follows (Hanna,
 8 1969):

$$Ri_b = \frac{(g / \theta_{v0})(\theta_{vz} - \theta_{v0})z}{u_z^2 + v_z^2}, \quad (1)$$

11 where θ_{v0} and θ_{vz} are the virtual potential temperatures at the surface and at height z , respectively,
 12 g / θ_{v0} is the buoyancy parameter, u_z and v_z are the horizontal wind speed at height z . As can be
 13 seen from Eq. (1), the bulk Richardson number method is computationally cheap because it only
 14 requires low-frequency data. Nonetheless, the biggest challenge associated with the bulk
 15 Richardson number method is that the value of Ri_{bc} , has to be determined as a prior known. In
 16 previous studies, the value of Ri_{bc} varies from 0.15 to 1.0 (Zilitinkevich and Baklanov, 2002;
 17 Jericevic and Grisogono, 2006; Esau and Zilitinkevich, 2010), with values of 0.25 and 0.5 most
 18 widely used (e.g., Troen and Mahrt, 1986; Holtslag and Boville, 1993). One important cause of
 19 the large variability of Ri_{bc} is the thermal stratification in the PBL. For example, Vogelezang and
 20 Holtslag (1996) reported the Ri_{bc} value of 0.16-0.22 in a nocturnal strongly stable PBL and 0.23-
 21 0.32 in a weakly stable PBL. For unstable PBLs, a value larger than 0.25 is usually needed (Zhang
 22 et al., 2011). Esau and Zilitinkevich (2010) also showed that the Ri_{bc} for nocturnal SBLs was
 23 smaller than for neutral and long-lived stable PBLs based on a large-eddy simulation database.
 24 More recently, a linear relationship between the Ri_{bc} and the atmospheric stability parameter has
 25 been proposed and examined under stable conditions, which further suggests the impact of thermal
 26 stratification on the Ri_{bc} (Richardson et al., 2013; Basu et al., 2014).

28 The objective of this study is to examine the variation of Ri_{bc} with different thermal
 29 stratification conditions. To do so, a representative value of Ri_{bc} for each type of PBLs (i.e.,
 30 strongly stable boundary layers, weakly stable boundary layers, and unstable boundary layers)
 31 needs to be inferred. In our study, the Tur method, the PTG method, the LLJ method, and the
 32 modified parcel method are used to determine the PBLHs from observations made in four field
 33 campaigns, which are called ‘observed’ PBLHs. Using these ‘observed’ PBLHs as benchmarks,
 34 best choices of Ri_{bc} values under different stratification conditions are then inferred so that the
 35 estimates of PBLHs with the bulk Richardson number method matches the ‘observed’ PBLHs.
 36 These inferred values of Ri_{bc} are used to explore the impact of thermal stratification on the Ri_{bc} .

1 The study is organized in the following way: Sect. 2 describes the observational data used in
2 this study; Sect. 3 compares estimates of PBLH from different methods that are widely used to
3 determine the PBLH from measurements; Sect. 4 focuses on the bulk Richardson number method
4 and describes the search for a best choice of Ri_{bc} under different stratification conditions. Sect. 5
5 concludes the paper.

6 7 **2. Observational Data**

8 Observational data from four field campaigns that are conducted under different surface and
9 atmospheric conditions are used in this study. These field campaigns are the Litang experiment,
10 the Atmospheric Radiation Measurement (ARM) experiment, the Surface Heat Budget of the
11 Arctic Ocean (SHEBA) experiment, and the Cooperative Atmosphere-Surface Exchange Study
12 (CASES) in 1999 (CASES99). Each of these four field campaigns is briefly described, as follows:

13 The Litang site is located over a plateau meadow in the southeast of the Tibetan Plateau. The
14 campaign provides 105 effective radio soundings of wind and temperature in three observational
15 periods (7-16 March, 13-22 May, and 7-16 July, 2008), with a typical 6-h interval (about 00:30,
16 06:30, 12:30, and 18:30 LST). The 30-minute averaged wind and temperature at 3 m collected by
17 an eddy covariance system are also used for calculating the bulk Richardson number.

18 The ARM experiment was carried out over a plain farmland in Shouxian, China, from 14 May
19 to 28 December, 2008. During the campaign, soundings were collected every 6 h (about 01:30,
20 07:30, 13:30, and 19:30 LST). Due to instrument malfunction, some data are excluded and a total
21 of 842 radio soundings are retained. The 30-minute averaged wind and temperature measured at 4
22 m by an eddy covariance system are also used.

23 The SHEBA site is located around the Canadian icebreaker Dec Groseilliers in the Arctic
24 Ocean. The dataset provides radio soundings from mid-October, 1997 to early October, 1998.
25 During this period, rawinsondes were released 2 to 4 times a day (around 05:15, 11:15, 17:15, and
26 23:15 LST). Since the near-surface (2.5 m) data available from 29 October, 1997 to 1 October,
27 1998 at the SHEBA are hourly averages (Andreas et al., 1999; Persson et al., 2002), the surface
28 observations and soundings do not overlap well in time. To ensure accuracy, only soundings
29 released within 15 minutes around the hour were used in this study, yielding a total of 168 records.

30 The CASES99 is the second experiment of CASES that conducted in Kansas, USA. The
31 terrain is relatively flat (the average slope is about 0.5 degrees). In the campaign, the National
32 Oceanic and Atmospheric Administration (NOAA) Long-EZ and Wyoming King Air
33 accomplished the aircraft measurements at 50 and 25 Hz sample rates, respectively, during 6-27
34 October, 1999 when the PBL was primarily stable. Since the lowest flight level was restricted (e.g.,
35 for security reasons), only 35 effective aircraft soundings are used in our study. The 5-minute
36 averaged near-surface (3 m) wind and temperature data recorded at the NO.16 flux tower in
37 CASES99 (www.eol.ucar.edu/projects/cases99) are also used. The surface observations and
38 soundings in CASES99 overlap well in time, but their horizontal positions slightly differ due to

1 the movement of aircraft. Due to the fact that most of the sounding data from CASES99 were
 2 collected under strongly stable conditions and data under other conditions were too limited, in this
 3 study, only soundings under strongly stable conditions (i.e., in Type I SBLs) are used except in
 4 Fig. 1 where one weakly stable boundary layer case from CASES99 is presented in order to
 5 compare the LLJ method to the Tur method.

6 In the postprocessing, a 20 m moving-window average is used for all the soundings from all
 7 the sites (except the turbulence measurements by aircrafts in CASES99) to remove the
 8 measurement noise.

10 3. PBLHs Determined from Observational Data

11 As mentioned in the introduction, the PBLs during a typical diurnal cycle are categorized into
 12 three types: Type I SBLs (i.e., strongly stable boundary layers at night), Type II SBLs (i.e.,
 13 weakly stable boundary layers at early morning/night), and UBLs (i.e., unstable boundary layers
 14 during the daytime). The PTG method, the LLJ method, and the modified parcel method are
 15 usually used to determine the PBLH for Type I SBLs, Type II SBLs, and UBLs, respectively.
 16 Based on previous studies (e.g., Holtslag and Boville, 1993; Vogelezang and Holtslag, 1996), they
 17 are classified using the surface heat flux H and the potential temperature θ profile:

$$18 \quad \begin{cases} H \geq \delta & \text{for UBLs} \\ H < \delta \text{ and } d^2\theta/dz^2 < 0 & \text{for Type I SBLs} , \\ H < \delta \text{ and } d^2\theta/dz^2 \geq 0 & \text{for Type II SBLs} \end{cases} \quad (2)$$

19 where δ is the minimum H for unstable conditions, which, in practice, is specified as a small
 20 positive value instead of zero (Liu and Liang, 2010). Due to different thermodynamic properties of
 21 land and ice, the value of δ is specified as 1 W m^{-2} over land and 0.5 W m^{-2} over ice through trial
 22 and error. Under stable conditions (i.e., $H < \delta$), the PBLs are further classified into Type I SBLs
 23 and Type II SBLs according to $d^2\theta/dz^2$. For Type I SBLs, the PTG decreases with height and
 24 the inversion near the surface is relatively strong, so there is always a sudden decrease of PTG at
 25 the PBL top (e.g., see Fig. 1a1). As such, the derivative of PTG with respect to z should be
 26 negative, that is, $d^2\theta/dz^2 < 0$. For Type II SBLs, the PTG increases with height and the
 27 inversion is relatively weak. No sudden change of PTG is seen at the PBL top (e.g., see Fig. 1a2)
 28 and thus $d^2\theta/dz^2 \geq 0$. In this study, $d^2\theta/dz^2$ is calculated between 40 m and 200 m; the
 29 selection of 40 m as the lower boundary is to avoid near-surface variability caused by landscape
 30 heterogeneity.

31 Note that cases with $-\delta < H < \delta$ (i.e., under near-neutral conditions) are typically treated as
 32 Type II SBL cases according to our classification. This is because stable stratification usually
 33 prevails above the boundary layer and wind shear is the only source of turbulence under near-
 34 neutral conditions. Both these features are similar to those of a stable boundary layer, and as a

1 result, the near-neutral cases are treated as SBL cases (Serbert et al., 2000). It appears there might
2 be an abrupt change in the calculation of PBLH at $H \approx \delta$ if different values of Ri_{bc} are used for
3 SBLs and UBLs, which is the aim of this study. However, we note that changes of Ri_{bc} at $H \approx \delta$
4 from SBLs to UBLs have little effect on the PBL height determination, because the Ri_b increases
5 drastically with height at the PBL top under near neutral condition and using Ri_{bc} for either SBLs
6 or UBLs gives reasonable estimates of PBLH (Fig. S1).

7 For any of the three types of PBLs, the Tur method is the most direct and accurate approach
8 for the PBLH estimation because it measures the turbulence intensity directly. Figure 1 shows
9 vertical profiles of potential temperature, mean wind velocity, bulk Richardson number, and wind
10 velocity perturbations from CASES99 for a Type I SBL (a1 to d1) and a Type II SBL (a2 to d2).
11 The wind velocity perturbations (u' , v' , w'), or turbulence intensities, are obtained by removing
12 the slowly varying part of the corresponding winds (u , v , w) through a high-pass wavelet filter
13 (Wang et al., 1999; Wang and Wang, 2004). In the Tur method, continuous wavelet transform is
14 applied to the absolute magnitude of turbulent fluctuations of each velocity component. The
15 PBLH is automatically determined to be the level at which the absolute magnitude of these
16 velocity fluctuations shows the most rapid decrease with height (Dai et al., 2011; 2014). The
17 PBLHs determined by u' , v' , w' are then averaged using the absolute magnitude of the reciprocal
18 velocity fluctuations as weights. As can be seen from Fig. 1d1 and 1d2, the PBLH determined by
19 the Tur method are denoted by the red solid lines.

20 Figure 1 further shows the PBLHs determined by with the PTG method (see the red solid line
21 on a1) and the LLJ method (see the red solid line on b2) for Type I and Type II SBLs, respectively.
22 It is clear that the estimates of PBLHs with these two methods are comparable to the PBLHs
23 determined from the Tur method, suggesting that the PTG method and the LLJ method work well
24 for Type I and Type II SBLs, respectively.

25 Figure 2 shows the sounding profiles taken from Litang on July 9, 2008 and the PBLHs
26 estimated by the PTG, LLJ, and modified parcel methods for the three different PBLs,
27 respectively. At midnight (0035 LST), the PBL was very stable due to radiative cooling from the
28 surface and is classified as a Type I SBL. According to the PTG method, the PBLH was found at
29 the top of the strong inversion (125 m, see Fig. 2a). In early morning (0635 LST), the surface
30 temperature increased and thus the inversion near the surface became weak; the low-level wind
31 speed increased rapidly and formed a LLJ. The PBL is classified as a Type II SBL. With the LLJ
32 method, the PBLH was determined at the height of the maximum wind (260 m, see Fig. 2b). As
33 the surface heating continues, a super-adiabatic layer in which the potential temperature decreases
34 with the height formed near the surface and a UBL was developed by midday (1245 LST). With
35 the modified parcel method, the PBLH is estimated to be 1654 m (see Fig.2c). Consequently, it
36 can be concluded that the three methods mentioned above are useful for a PBL with discernible
37 characteristics (Fig. 1 and 2).

1 However, for a PBL without these discernible characteristics, these methods may introduce
2 large biases (see Fig. 3, and also see e.g., Russell et al., 1974; Martin et al., 1988; Balsley et al.,
3 2006; Meillier et al., 2008). For Type I SBLs, when the underlying inversion is not strong, it will
4 be difficult to determine the PBLH by the PTG method due to the fact that the maximum PTG can
5 be less than the threshold γ_s (Fig.3b1). For Type II SBLs, when there is no clear wind-speed
6 maximum or when multiple maxima exist, the LLJ method will have difficulties in determining
7 the PBLH. For example, there were two maxima in the wind profile at (at 160 m and 400 m, see
8 Fig. 3c2). If the PBLH is simply determined as the height where the first maximum occurs, the
9 PBLH would be 160 m. Combining information from the Ri_b profile (Fig. 3d2), a more reasonable
10 estimate of the PBLH should be 400 m instead of 160 m since the Ri_b profile undergoes a
11 significant transition at 400m. For UBLs, similar complex situations may occur. The results of the
12 modified parcel method with a specified PTG threshold may be subjective since the threshold may
13 depend on the vertical resolution and data precision (Beyrich, 1997; Joffre et al., 2001). For
14 example, there are two PTG maxima at 900 m and 2000 m (see Fig. 3b3) due to the sharp drop of
15 relative humidity at these two heights. A more accurate estimate of the PBLH should be 900 m
16 when combining the information from the Ri_b profile (Fig. 3d3), while it might have been
17 determined to be 2000 m by the modified parcel method if $\gamma_c = 0.5 \text{ K (100m)}^{-1}$ is used.

18 Although these special cases do not always exist, they limit the applications of the three
19 methods. The accuracy of the determined PBLH can be improved with additional information, as
20 have been demonstrated before. The following sums up the procedures that are used in this study
21 for estimating PBLH by using these four methods: First, whenever turbulence measurements are
22 available, the Tur method is used to determine the PBLH. Second, for Type I SBL cases with a
23 relatively weak inversion (the local PTG maximum is $< 6.5 \text{ K (100 m)}^{-1}$ between 40 m-200 m), if
24 there is a LLJ, the case is reclassified to a Type II SBL; if not, the case is removed. Third, Type II
25 SBL cases without clear wind-speed maximum are removed. Last, when there are multiple wind
26 maxima for a Type II SBL or multiple PTG maxima for a UBL, the information from the Ri_b
27 profile is combined to determine the PBLH. With these procedures, the obtained PBLHs by using
28 these methods are treated as ‘observed’ PBLH hereafter. The observed PBLH and the bulk
29 stability parameter ($PBLH/L$, where L is the surface Obukhov length) for these four field
30 experiments are provided in Table 1.

31 32 **4. The Bulk Richardson Number Method and the Ri_{bc}**

33 The PTG method, the LLJ method, and the modified parcel method are usually used to
34 determine the PBLH in observational data. However, they do not work well when the PBL has no
35 distinct features that are required by these methods. Instead, the bulk Richardson number method
36 with a prescribed Ri_{bc} is often used in numerical methods to automatically determine the PBLH.
37 For example, in the non-local PBL scheme of the Community Climate Model version 2 (CCM2),
38 Eq. (1) is applied to estimate the PBLH with $Ri_{bc} = 0.5$. The computation starts by calculating the

1 Ri_b between the surface and subsequent higher levels of the model. Once Ri_b exceeds Ri_{bc} , the
 2 PBLH is derived by a linear interpolation between the level with $Ri_b > Ri_{bc}$ and the level below.

3 To avoid overestimating the shear production in Eq. (1) for relatively high wind speeds (i.e., in
 4 Type II SBL) and to account for turbulence generated by surface friction under neutral conditions,
 5 Vogelezang and Holtslag (1996) proposed an updated formulation, which is employed in the
 6 Community Atmosphere Model version 4 (CAM4), written as:

$$7 \quad Ri_b = \frac{(g/\theta_{vs})(\theta_{vz} - \theta_{vs})(z - z_s)}{(u_z - u_s)^2 + (v_z - v_s)^2 + 100u_*^2}, \quad (3)$$

8
 9 where z_s is the height of the lower boundary for the PBL (generally the top of the atmospheric
 10 surface layer), θ_{vs} is the virtual potential temperature at the height z_s , u_s and v_s are the wind speed
 11 components at z_s . z_s is often taken as 20 m, 40 m, or 80 m for SBLs (Vogelezang and Holtslag,
 12 1996) and taken as 0.1PBLH (Troen and Mahrt, 1986) or the height of super-adiabatic layer (z_{SAL})
 13 where the potential temperature first reaches its local minimum for UBLs. In our study, $z_s = 40$ m
 14 or 80 m are used under stable conditions while $z_s = 0.1PBLH$ or z_{SAL} are used under unstable
 15 conditions. The term $100u_*^2$ makes Eq. (3) more applicable for the near-neutral condition, which
 16 is classified as a Type II SBL in our study (Seibert et al., 2000).

17 Under unstable conditions, the virtual potential temperature at the lower boundary θ_{vs} is
 18 replaced by θ'_{vs} (Troen and Mahrt, 1986; Holtslag et al., 1995):

$$19 \quad \theta'_{vs} = \theta_{vs} + b_s \frac{\overline{(w'\theta'_v)_0}}{w_m}, \quad (4)$$

20 where $b_s=8.5$, $\overline{(w'\theta'_v)_0}$ is the virtual heat flux at the surface, and w_m is a turbulent velocity scale:

$$21 \quad w_m = (u_*^3 + 0.6w_*^3)^{1/3}, \quad (5)$$

22 and

$$23 \quad w_* = \left[\frac{g}{\theta_{v0}} \overline{(w'\theta'_v)_0} h \right]^{1/3} \quad (6)$$

24 is the convective velocity scale. The second term on the right side of Eq. (4) represents a
 25 temperature excess, which is a measure of the strength of convective thermals.

26 In this study, the virtual potential temperature is estimated as the potential temperature in the
 27 calculation because the former can lead to significant fluctuations in the estimated PBLH due to
 28 inaccurate humidity measurements (Liu and Liang, 2010).

29 After Ri_b is computed from Eqs. (3-6), the PBLH can be determined as the height where
 30 the Ri_b exceeds Ri_{bc} . In our study, instead of calculating the PBLH using a prescribed Ri_{bc} , we
 31 infer a representative Ri_{bc} for each type of PBLs using the ‘observed’ PBLH (see Section 3)
 32 and examine the variation of the inferred Ri_{bc} with thermal stratification. It is pointed out here

1 that our methodology is different from that of Richardson et al. (2013), who proposed a
2 stability-dependent Ri_{bc} for SBLs:

$$3 \quad Ri_{bc} = \alpha \frac{PBLH}{L} \quad (7)$$

4 where $PBLH/L$ is a bulk stability parameter and L is the surface Obukhov length. α is a
5 proportionality constant, which depends on surface characteristics and/or atmospheric
6 conditions. It varies between 0.03 and 0.21 with suggested values of 0.045 and 0.07
7 (Richardson et al., 2013; Basu et al., 2014). As shown in Fig. 1c1-c2, in the Type I SBL case,
8 a relatively reliable PBLH (133 m) was calculated with $\alpha=0.045$, but an overestimation (184
9 m) occurs when $\alpha=0.07$. While in the Type II SBL case both α values (0.045 and 0.07) yield
10 too small estimates of PBLH, because the two values are determined by idealized stably
11 large-eddy simulation datasets (Richardson et al., 2013) and observational datasets under
12 weakly and moderately stable conditions (Basu et al., 2014), respectively. In addition, Eq. (7)
13 is only applicable for SBLs but not UBLs. As such, instead of adopting this equation, we
14 inferred a representative Ri_{bc} value for each type of PBLs in our study.

15 Because each profile provides a Ri_{bc} value, a representative Ri_{bc} at each experimental site is
16 determined by fitting a linear relationship between the numerator and the denominator of Eq. (3)
17 at the PBLH, as will be shown in Sect. 4.1, or using statistical error minimization methods, as will
18 be shown in Sect. 4.2.

19

20 4.1 Representative Ri_{bc} from the linear fitting method

21 The representative Ri_{bc} values for Type I SBLs are shown in Fig. 4. The soundings are taken
22 from Litang, CASES99, ARM, and SHEBA, with the height z_s of 40 m (left) and 80 m (right).
23 Note that with $z_s = 80$ m, only cases with a PBLH ≥ 80 m are used. Except for CASES99, the
24 fitted Ri_{bc} values at each site are about 0.25. The difference in Ri_{bc} when different z_s (40 or 80 m)
25 are used is small. However, the results are slightly more consistent with $z_s = 40$ m compared to z_s
26 = 80 m, as can be seen from the higher correlation coefficients at ARM and CASES99. The value
27 of Ri_{bc} for Type I SBLs from CASES99 aircraft measurements is 0.20-0.21, which is smaller than
28 the values determined from radio soundings at other experimental sites. This may be because the
29 depth of the nocturnal inversion is generally thicker than the depth of the turbulent layer (Mahrt et
30 al., 1979; Andre and Mahrt, 1982). Therefore, the PBLH determined by the Tur method is smaller
31 than that determined by the PTG method. Note the Tur method is always used to determine the
32 PBLH for CASES99 since turbulence measurements are available at this site.

33 The inferred Ri_{bc} for Type II SBLs are shown in Fig. 5. Compared to the results in Fig. 4, the
34 correlation coefficients in Fig. 5 are smaller, indicating that the PBLH is more difficult to
35 determine for weakly stable boundary layers, which is consistent with previous studies (e.g., Esau
36 and Zilitinkevich, 2010). The correlation coefficients indicate that the agreement with $z_s = 80$ m is

1 slightly better than with $z_s = 40$ m. In particular, the inferred Ri_{bc} is sensitive to the height z_s in the
 2 SHEBA data. It changes from 0.21 to 0.29 as the height z_s changes from 40 m to 80 m. The main
 3 cause of the large variation of Ri_{bc} is because the LLJs above the ice surface in SHEBA are
 4 considerably strong (up to 20 m s^{-1}) and the vertical wind speed gradients are large, so the
 5 denominator in Eq. (3) decreases more rapidly with the height z_s than the numerator, which leads
 6 to an increase in the Ri_{bc} value when z_s increases from 40 m to 80 m. On the other hand, the Ri_{bc}
 7 over land varies little with z_s (Fig. 5), which is consistent with the findings of Vogelezang and
 8 Holtslag (1996) using the Cabauw data.

9 For UBLs, the height z_s is chosen to be $0.1PBLH$ (left) and z_{SAL} (right) in Fig. 6. As can be
 10 seen, the correlation coefficients are smaller than 0.4 at all sites, implying large variability in the
 11 Ri_{bc} inferred from each sounding. The representative value of Ri_{bc} is larger than 0.25 and varies
 12 from 0.28 to 0.34. However, it appears that the PBLH estimated by the bulk Richardson number
 13 method seems to be less sensitive to Ri_{bc} under unstable conditions. The estimates of PBLH using
 14 the bulk Richardson number method with $Ri_{bc} = 0.25$ or 0.5 are both in good agreement with the
 15 ‘observed’ PBLH at the three sites (Fig. 7). This is also in agreement with some previous studies
 16 (Troen and Mahrt, 1986). Therefore, the bulk Richardson number method is still reliable in
 17 estimating the PBLH of UBLs, despite the inferred Ri_{bc} shows large variability.

18

19 4.2 Representative Ri_{bc} from the error minimization method

20 It is seen that the linear fitting method yields small correlation coefficients under unstable
 21 conditions. Under stable conditions, the linear fitting method also has some disadvantages. For
 22 example, the inferred value of Ri_{bc} and the correlation coefficients highly depend on the larger
 23 value points, while the impact of the smaller value points is reduced (see e.g., Fig. 4a2). Therefore,
 24 we apply error minimization methods in this section to determine the optimal Ri_{bc} . The values of
 25 Ri_{bc} between 0.1-0.4 in stable conditions and 0.2-0.5 in unstable conditions are first used to
 26 calculate the PBLH; then, three statistical measures are used to examine the accuracy of the
 27 estimated PBLH (Gao et al., 2004):

$$28 \quad \text{Bias} = \frac{\sum_{i=1}^n |h_{Ri_b} - h_{obs}|}{n}, \quad (8)$$

$$29 \quad \text{SEE} = \sqrt{\frac{\sum_{i=1}^n |h_{Ri_b} - h_{obs}|^2}{n-2}}, \quad (9)$$

$$30 \quad \text{NSEE} = \sqrt{\frac{\sum_{i=1}^n (h_{Ri_b} - h_{obs})^2}{\sum_{i=1}^n (h_{obs})^2}}, \quad (10)$$

1 where h_{Rib} is the estimated PBLH by the bulk Richardson number method, and h_{obs} represents the
2 observed PBLH (i.e., calculated using the Tur, PTG, LLJ, or modified parcel method). Bias, SEE,
3 and NSEE is the absolute bias, standard error, and normalized standard error of h_{Rib} against h_{obs} ,
4 respectively, and n is the sampling number. Optimal values of Ri_{bc} can be determined based on the
5 minimum Bias, SEE, and NSEE. Note that the optimal Ri_{bc} determined based on the minimum
6 Bias, or the minimum SEE/NSEE can be different, however, the minimum SEE and the minimum
7 NSEE always yield the same optimal Ri_{bc} . In this study, minimum SEE and NSEE are used as the
8 final criterion for the optimal Ri_{bc} . To compare the error minimization method with the linear
9 fitting method, the correlation coefficients between h_{obs} and h_{Rib} are also presented.

10 The correlation coefficient, Bias, SEE, and NSEE with different values of Ri_{bc} for Type I
11 SBLs are shown in Fig. 8 when $z_s = 40$ m (top panels) and $z_s = 80$ m (bottom panels). Quadratic
12 curves are fitted to these data and then the maximum or minimum of the fitted quadratic curves
13 are obtained, which are used to select the optimal Ri_{bc} for each site. The weighted averages based
14 on the sampling number at the four sites are treated as the representative optimal Ri_{bc} across the
15 four sites (see the black dashed lines in Fig. 8) and the error bars depict the range of the optimal
16 Ri_{bc} across the four sites (Fig. 8). The variability of the optimal Ri_{bc} values for different sites is
17 probably caused by the diversity of surface characteristics (e.g., surface roughness). Compared to
18 the results with $z_s = 80$ m, the error bars are smaller and thus the optimal Ri_{bc} across different sites
19 are more concentrated with $z_s = 40$ m. Furthermore, the maximum correlation coefficient is larger
20 and the minimum Bias, SEE, and NSEE are smaller with $z_s = 40$ m.

21 Compared to Type I SBLs, correlation coefficients are smaller and errors are larger for Type II
22 SBLs (Fig. 9), again indicating that the PBLH for weakly stable boundary layers is more difficult
23 to determine. However, the maximum correlation coefficient, minimum Bias, SEE and NSEE, and
24 the range of optimal Ri_{bc} show smaller differences between different values of z_s (40 m or 80 m).
25 Compared to the results of the linear fitting method, the values of Ri_{bc} are generally larger for each
26 site, which is understandable given that the scatter distribution is mostly above the fitted lines in
27 Fig. 5, especially at ARM and SHEBA. The optimal Ri_{bc} based on minimum SEE and NSEE for
28 Type II SBLs is 0.30-0.31. The result is consistent with the value (= 0.3) from Melgarejo and
29 Deardorff (1974).

30 For UBLs, Figure 10 shows that the maximum correlation coefficient is larger, the minimum
31 Bias, SEE, and NSEE are smaller, and the values of optimal Ri_{bc} for each site are more
32 concentrated with $z_s = z_{SAL}$ (bottom panels) compared to $z_s = 0.1PBLH$ (top panels). Therefore,
33 z_{SAL} is more appropriate as the lower boundary height in estimating the PBLH under unstable
34 conditions. The minimum SEE and NSEE indicate that the optimal $Ri_{bc} = 0.39$ under unstable
35 conditions. The results with $z_s = 40$ or 80 m are also examined but not shown here. The maximum
36 correlation coefficient and minimum Bias, SEE, and NSEE are close to those with $z_s = 0.1PBLH$,
37 but the values of optimal Ri_{bc} are more scattered across different sites.

1 Through the above statistical error minimization methods, the optimal Ri_{bc} for different
2 stratifications and sites with different choices of z_s are summarized in Table. 2. It appears that the
3 optimal Ri_{bc} value increases when the PBL stability decreases (i.e., as the PBL becomes more
4 unstable). The optimal Ri_{bc} value is 0.24 ($z_s = 40$ m) or 0.23 ($z_s = 80$ m) for Type I SBLs, 0.30 (z_s
5 = 40 m) or 0.31 ($z_s = 80$ m) for Type II SBLs. And for UBLs, the optimal Ri_{bc} value falls between
6 0.33 and 0.39, depending on the choice of z_s . To be exact, the best choices of Ri_{bc} suggested by
7 this study are 0.24 ($z_s = 40$ m), 0.31 ($z_s = 80$ m), and 0.39 ($z_s = z_{SAL}$) for Type I SBLs, Type II
8 SBLs, and UBLs, respectively. Note z_s is recommended to be 80 m for Type II SBLs, given that
9 the surface layer is usually thicker for Type II SBLs than for Type I SBLs.

11 4.3 Impacts of Thermal Stratification on Ri_{bc}

12 With the above analyses, the best choices of Ri_{bc} are inferred under different thermal
13 stratification conditions. Hence, the traditional way of determining the PBLH using a single value
14 of Ri_{bc} without considering the dependence of Ri_{bc} on thermal stratification (e.g., Troen and Mahrt,
15 1986) needs to be revised. For example, the Yonsei University (YSU) PBL scheme in the Weather
16 Research and Forecasting (WRF) model assumes $Ri_{bc} = 0.25$ over land (Hong, 2010), while $Ri_{bc} =$
17 0.5 is used in the Holtslag and Boville (HB) boundary-layer scheme in CCM2 (Holtslag and
18 Boville, 1993). To examine the impact of thermal stratification on Ri_{bc} , we obtained a single
19 representative Ri_{bc} for all stratification conditions with the same sounding data from Litang, ARM,
20 and SHEBA sites, assuming the lower boundary height z_s of 40 m, 80 m, and z_{SAL} for Type I SBLs,
21 Type II SBLs, and UBLs, respectively. According to the minimum SEE and NSEE, the optimal
22 choice of Ri_{bc} for all PBL types is 0.33 (Fig. 11), which is close to that used in CAM4 ($Ri_{bc} = 0.3$,
23 Neale et al., 2010). In Fig. 12, the errors when a single value of Ri_{bc} is used ($Ri_{bc} = 0.33$
24 determined by our study, $Ri_{bc} = 0.25$ in WRF-YSU, and $Ri_{bc} = 0.5$ in CCM2-HB) are presented, as
25 compared to the errors with a new scheme that uses $Ri_{bc} = 0.24$, 0.31, and 0.39 for Type I SBLs,
26 Type II SBLs, and UBLs, respectively. It is found that the new scheme with variable Ri_{bc} is more
27 reliable in estimating PBLH, suggesting that the impact of atmospheric stability or thermal
28 stratification on Ri_{bc} is significant and the variation of Ri_{bc} with atmospheric stability should be
29 taken into account when estimating the PBLH using the bulk Richardson number method.

30 To further investigate the improvements in estimating PBLHs with the new, variable Ri_{bc}
31 values, simulations using CAM4 are conducted at the ARM site, with the default (= 0.3) and the
32 new, variable Ri_{bc} values used to estimate PBLHs. Fig. 13 shows a comparison between the
33 observed and the CAM4-simulated PBLHs with the default and new Ri_{bc} values over a six-day
34 period. It can be seen that the simulated PBLHs with the new Ri_{bc} values have a more pronounced
35 diurnal cycle, which are also closer to the observations. Over the whole observational period,
36 results indicate that the Bias, SEE, NSEE are 270.1 m, 379.3 m, 0.75 with the new, variable Ri_{bc}
37 values, respectively, and are 306.2 m, 417.5 m, 0.83 with the default Ri_{bc} value, respectively.
38 Again, these results indicate that the impact of thermal stratification on Ri_{bc} should be considered

1 in calculating PBLH with the bulk Richardson number method and the new Ri_{bc} values determined
2 in this study improves model results in real applications. It is pointed out here that there are still
3 large biases in the CAM4-simulated PBLH even with the new Ri_{bc} values, which are probably
4 related to the biases in the model physics and parameterizations (e.g., parameterizations of land-
5 atmospheric interactions and boundary layer turbulence). Unraveling how biases in these model
6 physics and parameterizations affect the PBLH is nevertheless out of the scope of this study.

8 **5. Conclusions**

9 The PBLH is an important parameter in boundary layer research and accurate estimates of the
10 PBLH are vital for many environmental applications. In this study, we investigated several
11 methods for computing the PBLH under different stratification conditions. The Tur method is
12 considered as the most accurate approach for any atmospheric stratification due to its direct
13 measurement of turbulence intensity. However, such a method is expensive and thus cannot be
14 widely applied. On the other hand, determining of the PBLH with radio soundings through the
15 PTG, LLJ, and modified parcel methods is more affordable. These methods usually work well
16 when the PBL has certain unique features but may fail under special conditions (e.g., a weak
17 underlying inversion for strongly stable boundary layers, multiple wind maxima for weakly stable
18 boundary layers, and no clear maximum of vertical gradient of potential temperature for unstable
19 boundary layers). With corrections made for these special cases, we used the Tur, PTG, LLJ, and
20 modified parcel methods to determine PBLHs from Litang, ARM Shouxian, SHEBA, and
21 CASES99 field experiments and the estimated PBLHs using these methods are treated as observed
22 PBLHs.

23 The bulk Richardson number method is more commonly used in numerical models due to its
24 reliability for all atmospheric stratification conditions, which requires a specified value of the bulk
25 Richardson number for the entire PBL, or Ri_{bc} . In many numerical models, the Ri_{bc} is specified as
26 one single value (e.g., 0.25 for WRF-YSU, 0.5 for CCM2-HB, 0.3 for CAM4) and hence its
27 dependence on the thermal stratification is ignored. This study infers a representative Ri_{bc} for each
28 stratification condition from observed PBLHs using linear fitting and statistical error minimization
29 approaches. Results indicate that the best choices for Ri_{bc} are 0.24, 0.31, and 0.39 for strongly
30 stable boundary layers (Type I SBLs), weakly stable boundary layers (Type II SBLs), and unstable
31 boundary layers (UBLs), respectively. Both offline and online evaluation shows the new and
32 variable Ri_{bc} values proposed in this study yield more reliable estimates of the PBLH, suggesting
33 that the variation of Ri_{bc} should be considered in the computing the PBLH with the bulk
34 Richardson number method.

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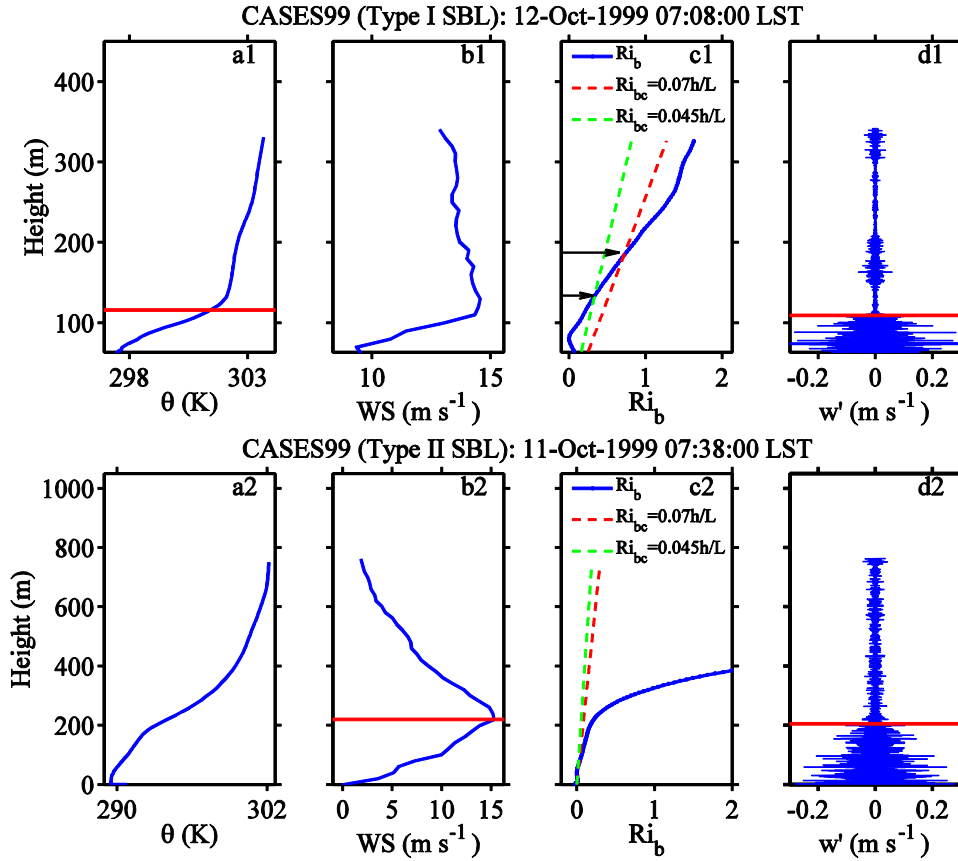
1 Table 1. The ‘observed’ PBLH and the stability parameter at four observational sites.

PBL Types	PBLH (m)				<i>PBLH/L</i>			
	Litang	CASES99	ARM	SHEBA	Litang	CASES99	ARM	SHEBA
Type I SBL	45~265	25~157	54~593	42~414	0.12~323.0	1.5~94.2	0.22~327.2	0.4~38.3
Type II SBL (H<0)	68~543	\	131~670	97~312	0.64~74.8	\	0.36~113.1	0.1~21.3
Type II SBL (H>0)	357~678	\	152~879	138~414	-33.4~-0.32	\	-34.1~-0.2	-55.1~-0.01
UBL	315~2594	\	293~1693	121~981	-866.4~-4.3	\	-350.9~-1.3	-342~-0.03

1 Table 2. Inferred bulk Richardson number of the entire PBL, Ri_{bc} , for different types of PBLs and sites, with
 2 different values of z_s , n refers to the sample number. ‘*’ indicates the best choice.

Regime	z_s (m)	Litang		CASES99		ARM		SHEBA		Total	
		Ri_{bc}	n	Ri_{bc}	n	Ri_{bc}	n	Ri_{bc}	n	Ri_{bc}	n
Type I	40	0.25	19	0.22	29	0.24	373	0.25	27	0.24*	448
	80	0.26	11	0.21	21	0.23	211	0.24	17	0.23	261
Type II	40	0.27	53	\		0.32	194	0.24	49	0.30	296
	80	0.24	53	\		0.33	194	0.31	49	0.31*	296
UBL	40	0.41	23	\		0.36	182	0.20	75	0.33	280
	80	0.41	23	\		0.38	182	0.32	75	0.39	280
	0.1PBLH	0.42	23	\		0.39	182	0.34	62	0.38	267
	z_{SAL}	0.39	23	\		0.41	182	0.36	75	0.39*	280

3

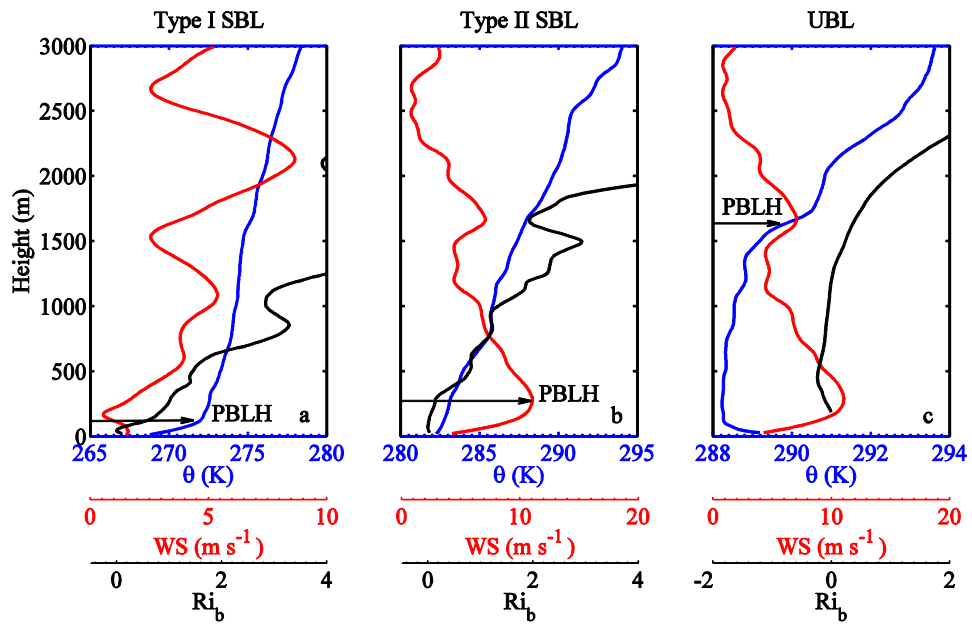


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2 Figure 1. Examples of vertical profiles of the Type I SBL (upper panels) and the Type II SBL (lower panels) from
 3 CASES99 aircraft measurements: (a) potential temperature (K); (b) horizontal wind speed (m s^{-1}); (c) bulk
 4 Richardson number Ri_b and Ri_{bc} ; (d) w' perturbation (m s^{-1}). The red solid lines on (a1) and (b2) denote the PBLH
 5 calculated by the PTG and LLJ methods, respectively, and those on (d) denote the PBLH determined by the Tur
 6 method. The black arrows on (c1) denote the PBLHs determined by the bulk Richardson number method with Ri_{bc}
 7 from Eq. (7).

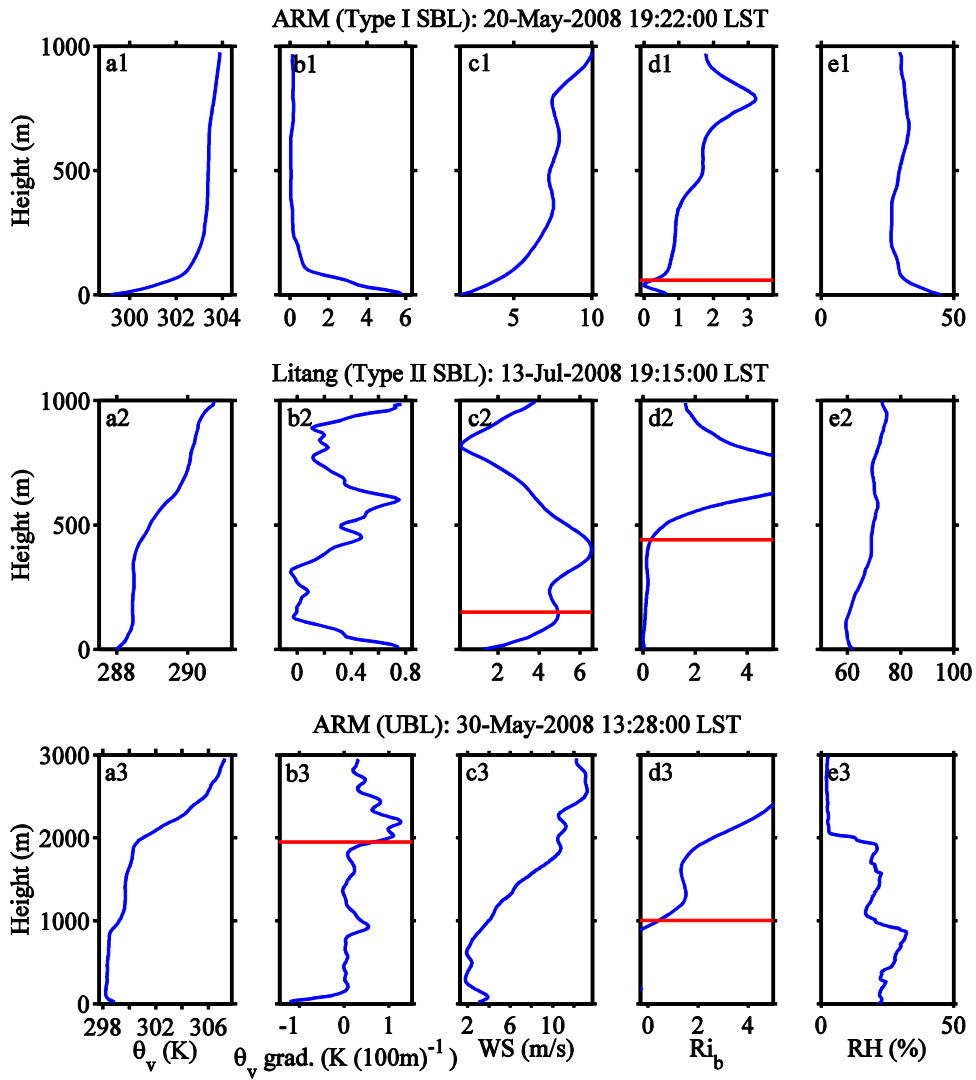
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2 Figure 2. Typical profiles of potential temperature (blue), wind speed (red), and Ri_b (black) for different types of
 3 boundary-layers: (a) Type I SBL, (b) Type II SBL, and (c) UBL. The indicated PBLHs in (a)-(c) are calculated by
 4 the PTG, LLJ, and modified parcel methods, respectively. The observations in (a)-(c) are from Litang on 08 July,
 5 2008 16:35 UTC (0035 LST), 08 July, 2008 2245 UTC (0645 LST), and 09 July, 2008 0445 UTC (1245 LST),
 6 respectively.



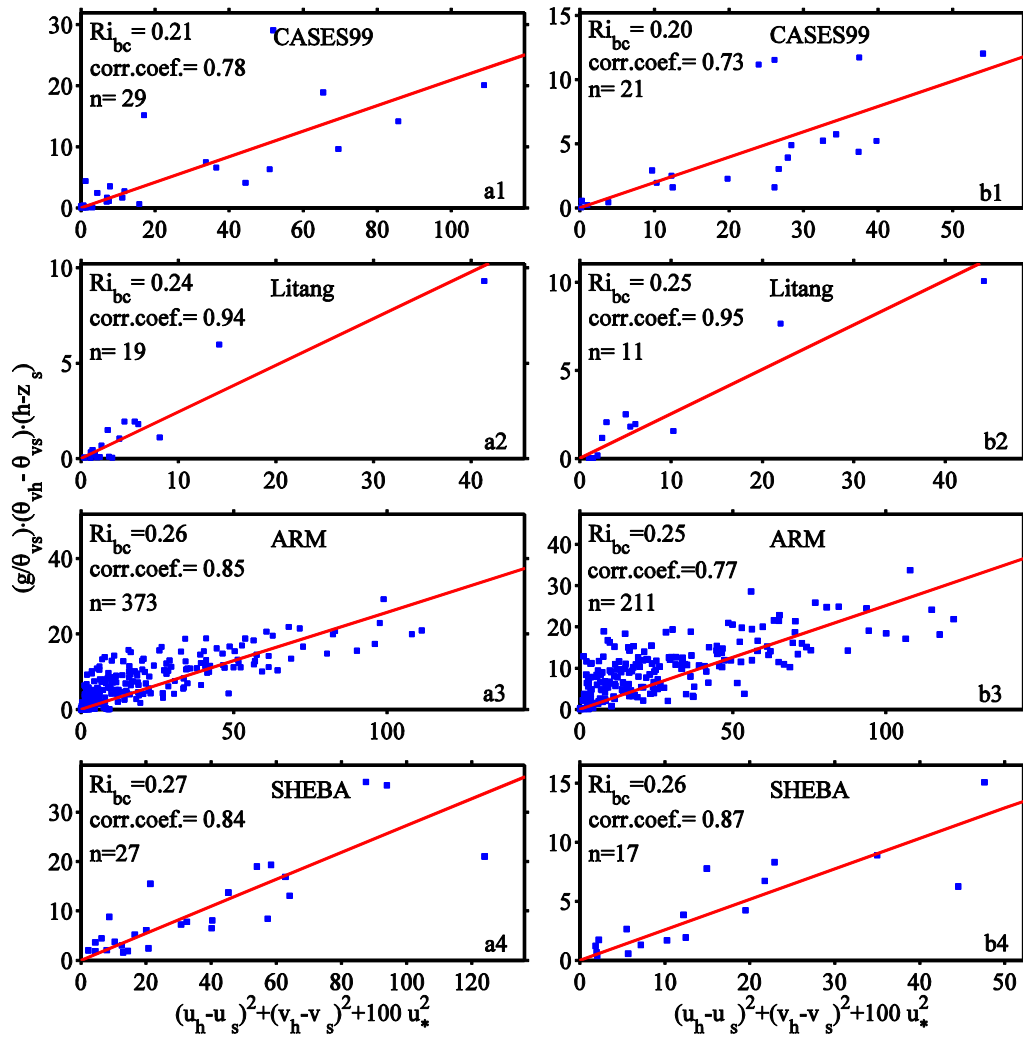
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Figure 3. Examples of vertical profiles in Type I SBLs (upper panels), Type II SBLs (middle panels), and UBLs (lower panels): (a) potential temperature (K); (b) potential temperature gradient (K (100m)⁻¹); (c) horizontal wind speed (m s⁻¹); (d) bulk Richardson number Ri_b ; (e) relative humidity (%). The red solid lines on (b3), (c2), and (d1-d3) denote the PBLH determined by the modified parcel, LLJ, and bulk Richardson number methods, respectively.

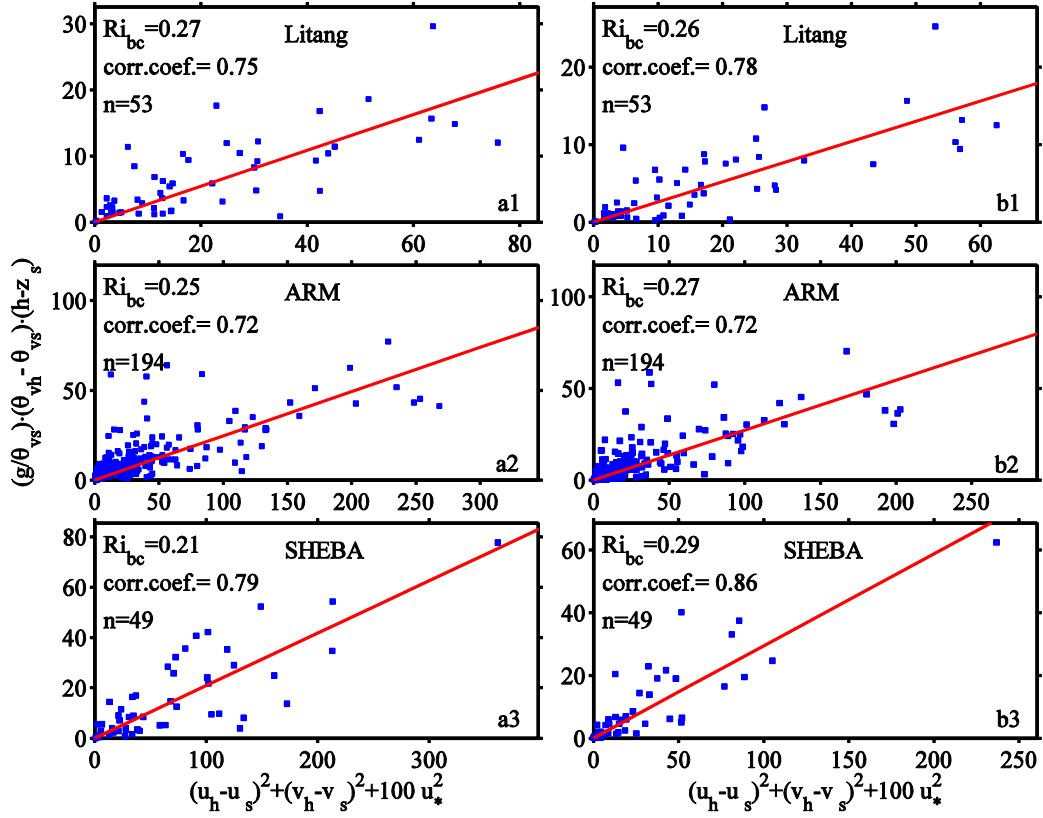
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2 Figure 4. Linear fitting method inferred Ri_{bc} for Type I SBLs, with $z_s = 40$ m (left) and $z_s = 80$ m (right). The red
 3 solid lines are the best linear fittings and their slopes represent the values of Ri_{bc} .

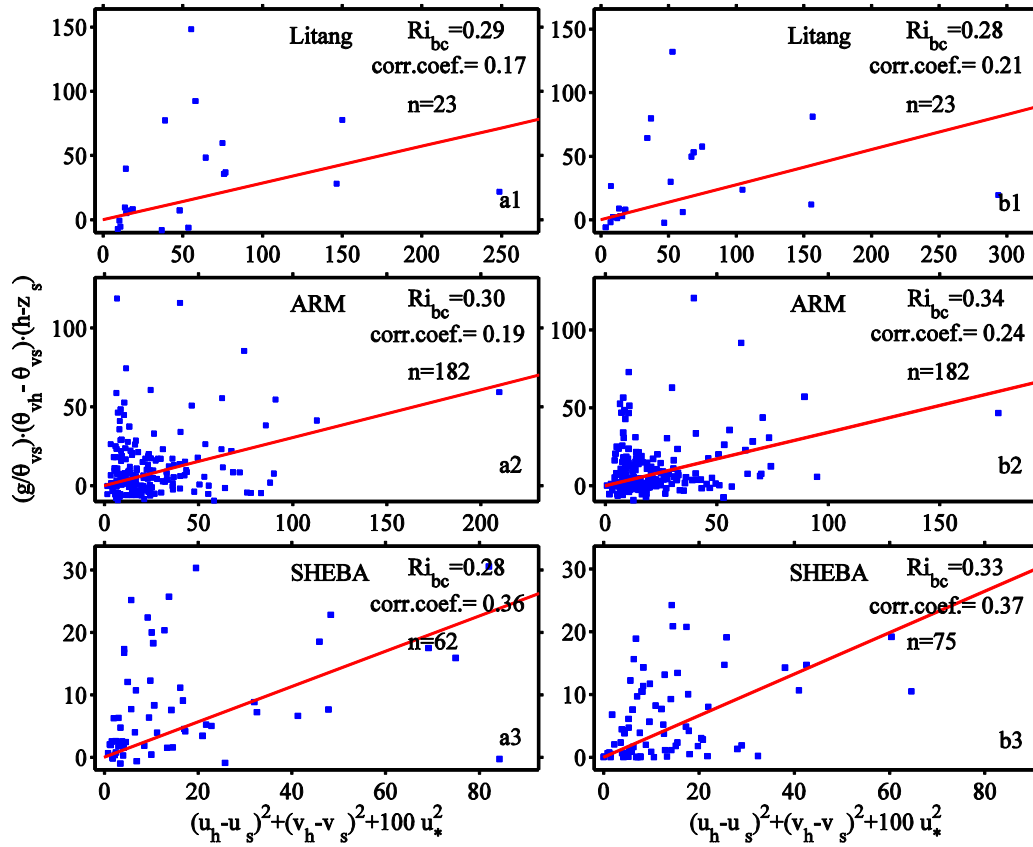


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2 Figure 5. Linear fitting method inferred Ri_{bc} for Type II SBLs, with $z_s = 40\text{m}$ (left) and $z_s = 80\text{ m}$ (right). The red
 3 solid lines are the best linear fittings and their slopes represent the values of Ri_{bc} .

4

5

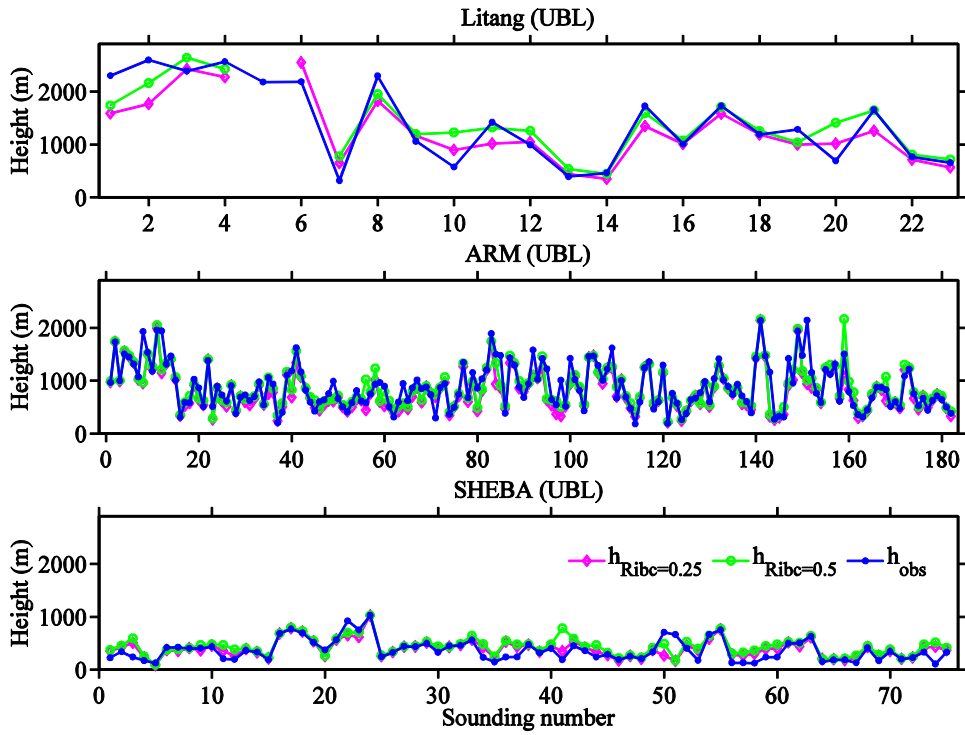


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Figure 6. Linear fitting method inferred Ri_{bc} for UBLs, with $z_s = 0.1$ PBLH (left) and $z_s = z_{SAL}$ (right). The red solid lines are the best linear fittings and their slopes represent the values of Ri_{bc} .

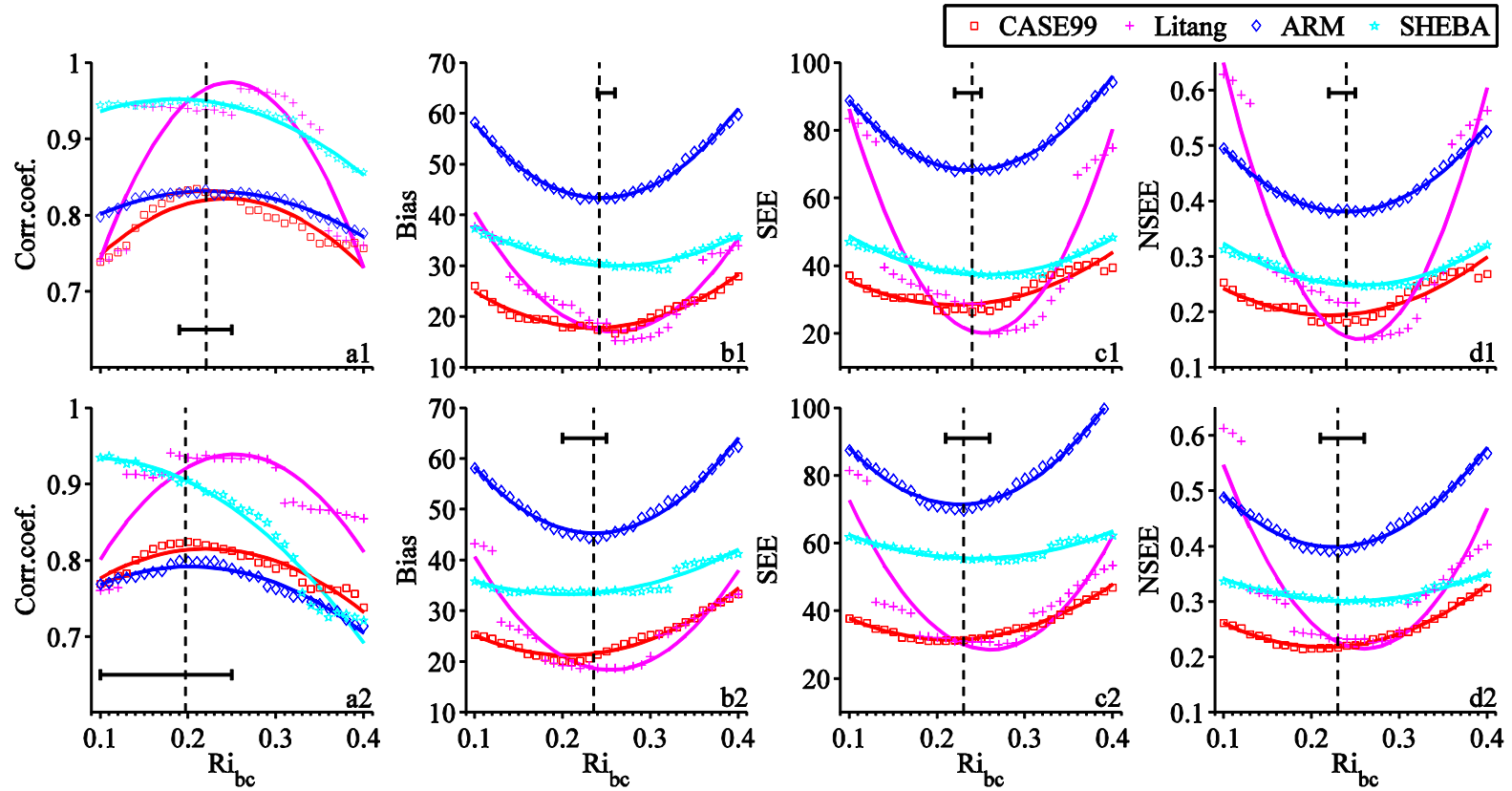
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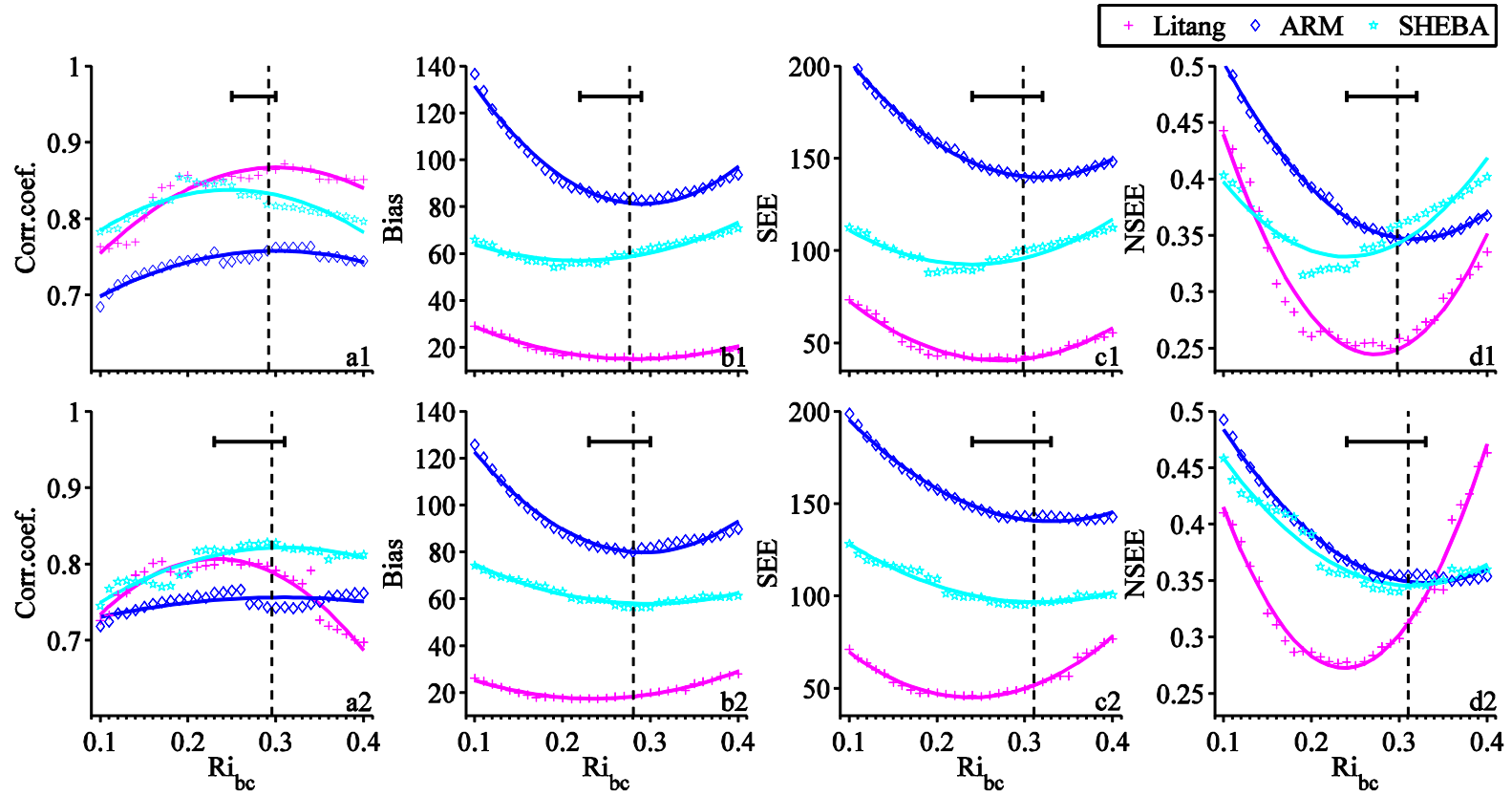
2 Figure 7. Comparisons of the heights of UBL at different sites determined by the bulk Richardson number method
 3 with $Ri_{bc} = 0.25$ (diamond) and 0.5 (circle) and the observed PBLHs (point).

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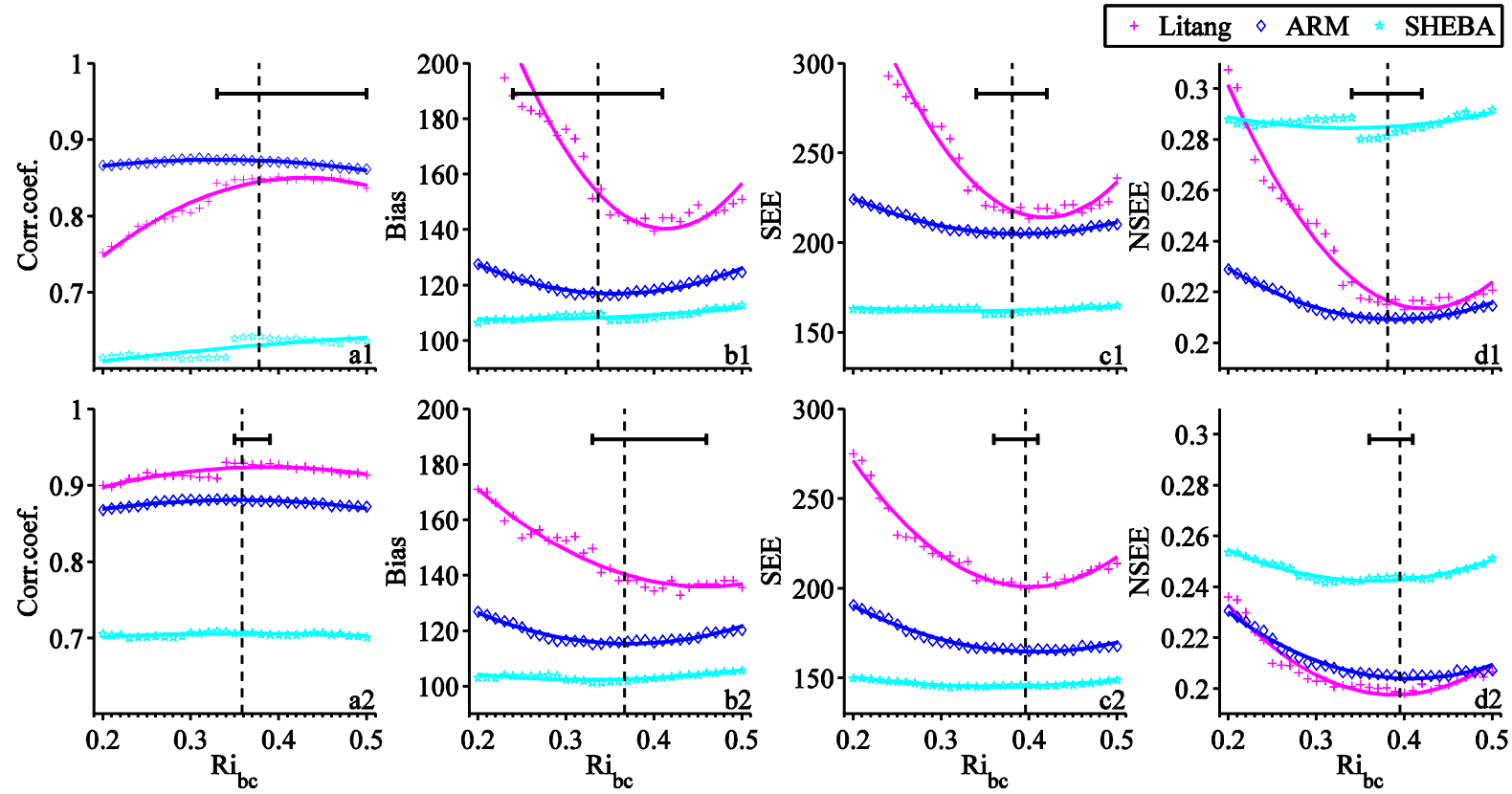
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3 Figure 8. Comparison between estimated PBLH using the bulk Richardson number method with $z_s = 40\text{m}$ (upper panels) and $z_s = 80\text{m}$ (lower panels) and observed PBLHs for Type I SBLs. The
 4 correlation coefficient (a), bias (b), standard error (c), and normalized standard error (d) are shown. The sounding data are taken from Litang (plus sign), CASES99 (square), ARM Shouxian
 5 (diamond), and SHEBA (pentacle). The curved lines are obtained by quadratic curve-fitting, the black vertical dashed lines indicate a representative Ri_{bc} for all four sites, and the error bars
 6 indicate the range of Ri_{bc} across the four sites.



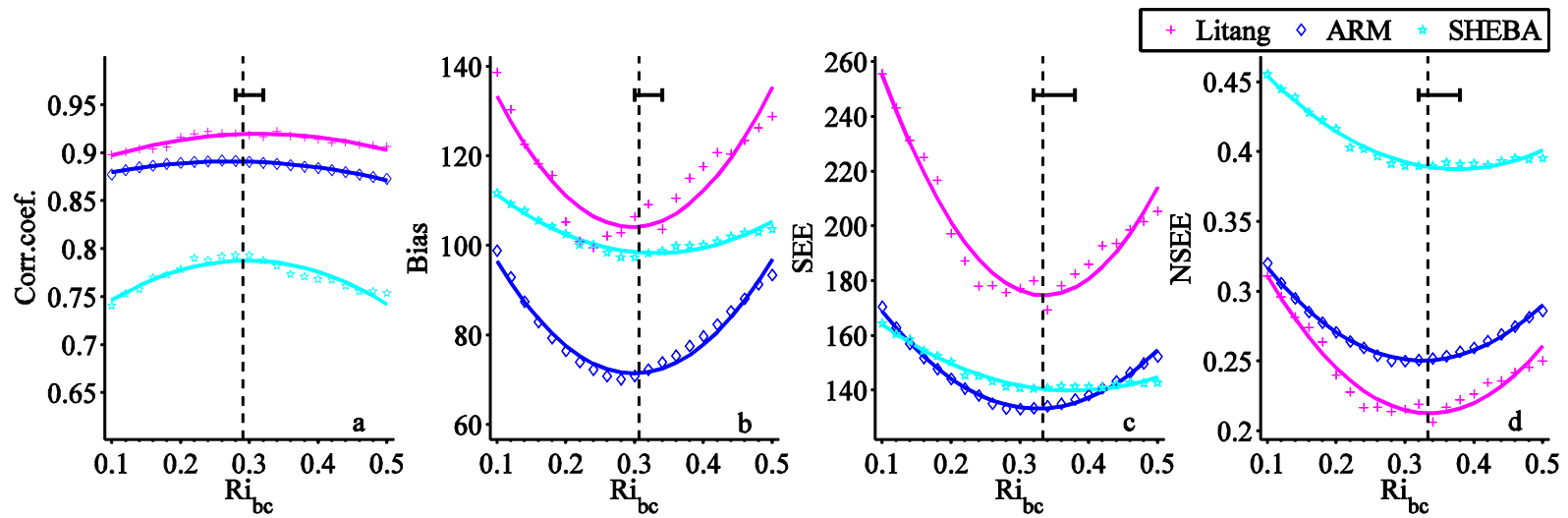
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2 Figure 9. Comparison between estimated PBLHs using the bulk Richardson number method with $z_s = 40\text{m}$ (upper panels) and $z_s = 80\text{m}$ (lower panels) and observed PBLHs for Type II SBLs.
 3 The correlation coefficient (a), bias (b), standard error (c), and normalized standard error (d) are shown. The sounding data are taken from Litang (plus sign), ARM Shouxian (diamond), and
 4 SHEBA (pentacle). The curved lines are obtained by quadratic curve-fitting, the black vertical dashed lines indicate a representative Ri_{bc} for all three sites, and the error bars indicate the range of
 5 Ri_{bc} across the three sites.



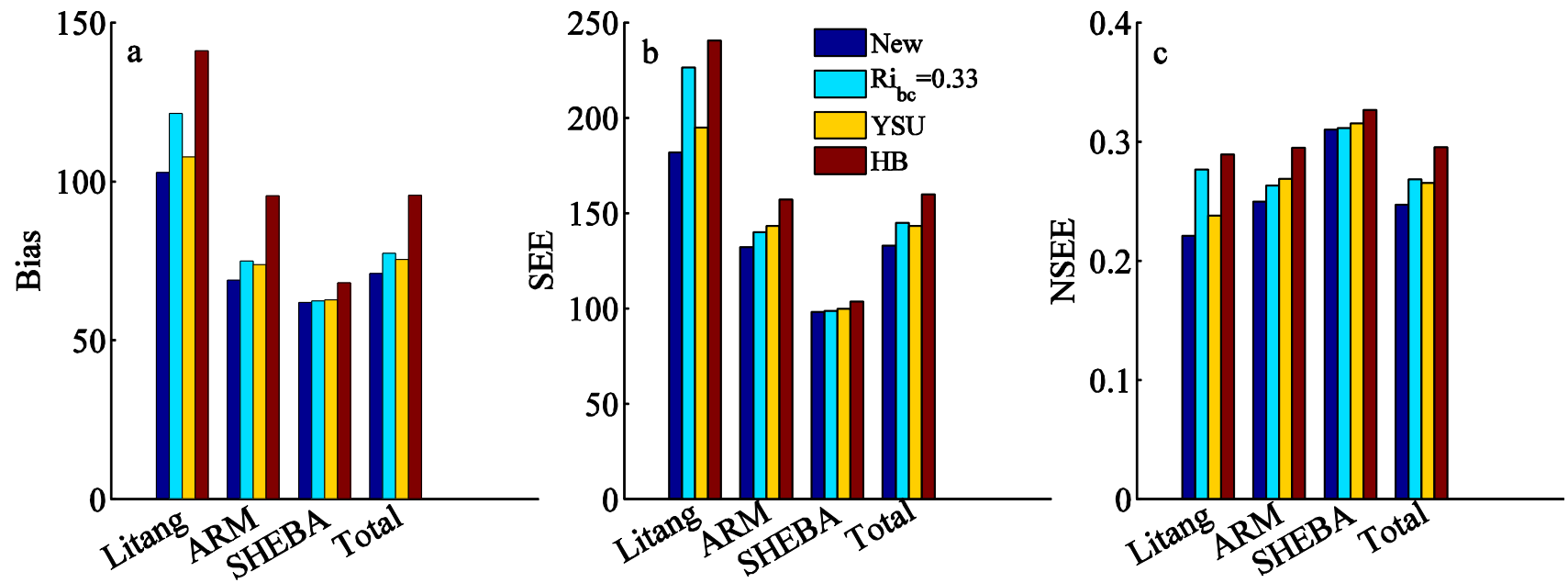
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2 Figure 10. Comparison between estimated PBLHs using the bulk Richardson number method with $z_s=0.1$ PBLH (upper panels) and $z_s = z_{SAL}$ (lower panels) and observed PBLHs for UBLs. The
 3 correlation coefficient (a), bias (b), standard error (c), and normalized standard error (d) are shown. The sounding data are taken from Litang (plus sign), ARM Shouxian (diamond), and SHEBA
 4 (pentacle). The curved lines are obtained by quadratic curve-fitting, the black vertical dashed lines indicate a representative Ri_{bc} for all three sites, and the error bars indicate the range of Ri_{bc}
 5 across the three sites.



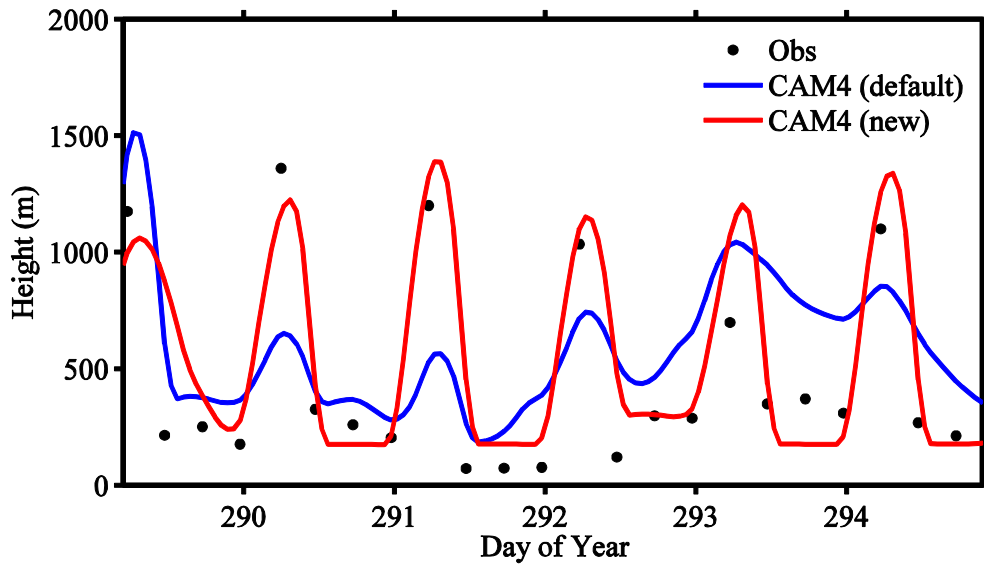
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2 Figure 11. Comparison between estimated PBLH using the bulk Richardson number method and observed PBLHs for all types of PBLs. The correlation coefficient (a), bias (b), standard error
 3 (c), and normalized standard error (d) are shown. The sounding data are taken from Litang (plus sign), ARM Shouxian (diamond), and SHEBA (pentacle). The curved lines are obtained by
 4 quadratic curve-fitting, the black vertical dashed lines indicate a representative Ri_{bc} for all three sites, and the error bars indicate the range of Ri_{bc} across the three sites.



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2 Figure 12. Comparisons between observed and estimated PBLHs with a single $Ri_{bc}=0.33$ for all PBL conditions, with $Ri_{bc}=0.25$ as in the YSU scheme, with $Ri_{bc}=0.5$ as in the HB scheme, and
 3 with the new, variable values ($Ri_{bc}=0.24, 0.31,$ and 0.39 for Type I SBLs, Type II SBLs, and UBLs, respectively): (a) bias; (b) standard error; (c) normalized standard error.



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Figure 13. Comparison of observed and simulated PBLHs using CAM4 with the default and new Ri_{bc} values during 16-21 Oct, 2008 at the ARM site.