

[C1] This revised manuscript has been greatly improved. The target of this paper (EC and sulfate) is clearly defined. An additional simulation, NICAM-g6, helps to show advantages of the stretched grid system. Comparisons with WRF-CMAQ are useful to see effects of different model types. Statistical parameters shows the model performance quantitatively. All the figures have been much improved. A controversial part including health impacts has been eliminated. Now all the reviewer's comments have been addressed.

[A1] Thank you very much for editing and reviewing our manuscript again. We have shown the Point by Point Clarifications to the comments and suggestions.

[C2] I felt that the former manuscript exaggerated the model performance without any confidence. Now the description of this revised manuscript is based on concrete reasons. It also clearly and honestly indicates limitations of this model. I still have comments on this revised manuscript. I recommend that this paper is published after all the following comments have been addressed.

[A2] Thank you very much for reading our manuscript again. We have addressed your comments as follows;

Specific comments:

[C3] Line 39: What kind of the underestimation is caused by what kind of the underestimation in China?

[A3] Thank you for your question. Here, we would like to mention that the underestimation of the simulated sulfate and SO<sub>2</sub> concentrations over East Asia is strongly affected by the underestimation of the simulated sulfate and SO<sub>2</sub> concentrations in China. Therefore, we have modified the sentences as follows; "This model generally reproduces monthly mean distributions of the observed sulfate and SO<sub>2</sub> over East Asia, with the high correlations ( $R > 0.6$ ), but the underestimation of the simulated concentrations by 40% (sulfate) and 50% (SO<sub>2</sub>). Their underestimation of the simulated sulfate and SO<sub>2</sub> concentrations over East Asia are strongly affected by their underestimation in China and ...."

[C4] Line 46 and others: What kind of a variation is intended to be shown by the word "weekly variation"? It may imply a typical variation from Sunday to Saturday due to human activities. Please reconsider the word.

[A4] Thank you for your suggestions. In this simulation, we did not consider the weekly cycle of the anthropogenic emission inventories, so that we cannot capture the weekly variation due to human activities. We just represented a variation governed by the synoptic system. Therefore, we have changed the word 'weekly variation' into 'synoptic variation'.

[C5] Line 120 and Line 306: I think it is not necessary to mention the model inter-comparison here in the context of this paper.

[A5] Thank you for your suggestions. We have removed this part from the revised manuscript.

[C6] Line 189: What does “current study” mean here?

[A6] We have modified the word ‘current’ into ‘future’, because as a next step for our study, we aimed to extend the use of the stretched-grid system to the global uniform high-resolution NICAM-SPRINTARS,

[C7] Line 253: Takemura et al., 2002a -> Takemura et al., 2002

[A7] Thank you for your corrections.

[C8] Line 274: Is “one-hour” accurate?

[A8] Thank you for your suggestions. It is not one-hour, but ‘one-day’. We have corrected it.

[C9] Line 299: Are these cloud and precipitation schemes used in NICAM-g6str too? If so, these description should be included in the section 2.1.

[A9] Thank you for your comments. The answer is NO. The cloud and precipitation schemes used in NICAM-g6str and NICAM-g6 are different, because of different spatial resolution. The schemes used in NICAM-g6str were mentioned in Line 203-206. Therefore, we mentioned the schemes used in NICAM-g6 here. To clarify them, we have modified this part as follows; “Apart from the NICAM-g6str simulation, in the NICAM-g6 simulation the cloud physics apply both ....”.

[C10] Line 376: Arakane et al., 2013 -> Arakane et al., 2014

[A10] Thank you for your corrections.

[C11] Line 380: Is MSL Mean Sea Level? What is “for the model bottom of MSL”?

[A11] Thank you for your comments. This is totally our mistake. We have deleted the words ‘for the model bottom of MSL’ from the revised manuscript. We also have removed the words in the caption of Figure 3 in the revised manuscript.

[C12] Line 388: This sentence is confusing. Why is NICAM-g6str higher than NICAM-g6 because the spatial resolution in NICAM-g6str is finer than that in NCEP-FNL?

[A12] Thank you for your suggestions. This sentence is also our mistake. In the target region, at both the surface and the height of 5 km, the absolute biases in the temperature between NICAM-g6str and NCEP-FNL or between NICAM-g6 and NCEP-FNL are within 1.5 °C. The 3 °C difference mentioned in the revised manuscript is found around Chinese inner mountains, which is out of the main target region. Therefore, we have removed the statement from the new revised manuscript. The difference in the spatial resolution between NICAM-g6str and NCEP-FNL causes the difference in the temperature around the Japanese Alps. In the new revised manuscript, we have modified this part as follows; “The absolute biases in the temperature between NICAM-g6str and NCEP-FNL or between NICAM-g6 and NCEP-FNL are within 1.5 °C at the surface and the height of 5 km. Around the Japanese Alps, however, the NICAM-g6str-simulated temperature is lower than the NCEP-FNL-estimated one by at most 2.5 °C, because of the differences in the resolved topography due to the different spatial resolution between NICAM-g6str and NCEP-FNL.”

[C13] Line 389: larger -> higher?

[A13] Thank you for your corrections.

[C14] Line 394: Does it mean that the stretched grid system does not affect the general circulations and only affects fields around complex topography?

[A14] The first comment ‘the stretched grid system does not affect the general circulations’ is totally right. We found that the stretched grid system worked correctly without any artificial flows. In contrast, we cannot conclude the second comment ‘only affects fields around complex topography’ in this section, because we only compared the meteorological fields obtained by NICAM-g6str with those by NCEP-FNL (coarse resolution). Surely, we found differences between NICAM-g6str and NCEP-FNL, but this is mainly caused by the differences in the spatial resolution between NICAM-g6str and NCEP-FNL. We have modified the sentence as follows; “Therefore, it is concluded that the stretched-grid systems does not affect the general circulations under the nudging technique in this study”.

[C15] Line 397: Aerosol concentrations should not be six-hourly “instant” values.

[A15] Thank you for your correction. Yes, we actually used six-hourly ‘mean’ values. We have corrected it.

[C16] Line 406: NICAM-g6-simulated -> NICAM-g6str-simulated?

[C17] Line 469: NICAM-g6-str -> NICAM-g6str

[A16&A17] Thank you for your corrections.

[C18] Line 480: NICAM-g6str reproduces with a large uncertainty?? What does it mean?

[A18] Thank you for your comments. We have reconsidered it and removed the words ‘with a large uncertainty’ from the revised manuscript.

[C19] Line 561: NICAM-g6str at Tsukuba -> NICAM-g6 at Tsukuba

[A19] Thank you for your comments. We have checked Table 2, but we think our statement here is corrected. So we did not change the word ‘NICAM-g6str at Tsukuba’.

[C20] Line 567: I do not understand why reasons for August 12 and 14 can be assumed like this. How about a plume from volcanoes? A plume from volcanoes sometimes causes a high peak, and models sometimes fail to simulate its exact path. I think it can be checked in simulated results.

[A20] Thank you for your suggestion. Surely, the volcanoes like Miyakejima could affect the SO<sub>2</sub> concentrations over the Kanto region and in our simulation we considered the SO<sub>2</sub> emission from volcanoes, but in the target year the SO<sub>2</sub> emission from industrial sources, especially power station in Tokyo Bay, is the largest contributor to the SO<sub>2</sub> concentrations over the Kanto region, even though the volcanoes emit some of SO<sub>2</sub> plumes. When we checked the model results of SO<sub>2</sub> during these days, the strong SO<sub>2</sub> plumes from industrial sources over the Tokyo Bay arrived at the inner areas such as Kisai on August 12. In contrast, on August 14 the situation is different. On August 14, the observed wind speed and direction were not special and they are almost comparable to the NICAM-g6str-simulated ones (Figures 7 and 8). However, NICAM-g6str did not reproduce the peak of the observed SO<sub>2</sub>. Therefore, we assumed that some of local SO<sub>2</sub> emission was stronger. This SO<sub>2</sub> emission could include SO<sub>2</sub> from volcanoes, because the daily emission strength of SO<sub>2</sub> from volcano is unknown. Therefore, we have modified this part of the revised manuscript as follows; “On August 12, NICAM-g6str normally reproduced the peaks of the observed SO<sub>2</sub>, but with the blunter and slightly shifted peaks. In the NICAM-g6str simulation, the strong SO<sub>2</sub> plumes from industrial sources over the Tokyo Bay arrived at the inner areas such as Kisai. On August 14, although the NICAM-g6str-simulated winds were comparable to the observed ones (Figures 7 and 8), NICAM-g6str did not reproduce the sharp peaks of the observed SO<sub>2</sub>,

especially at Komae and Tsukuba. It may imply that special meteorological fields cause the observed peaks on August 12, whereas unaccounted SO<sub>2</sub> emission from local sources or sporadic volcanoes is stronger on August 14.”

[C21] Line 581: Japanese areas are not shown in Figures 14 and 15 for EC.

[A21] Thank you for your comment. Yes, the observation for EC is not available in Japan. Here, we just compared the results obtained by NICAM-g6str with those obtained by NICAM-g6. Therefore, we have modified it in the revised manuscript. “In China, the NICAM-g6str-simulated EC concentrations are comparable to the NICAM-g6-simulated ones with the R values of 0.71 (NICAM-g6str) and 0.68 (NICAM-g6), whereas in Japan (no available measurements) the NICAM-g6str-simulated EC concentrations are larger than NICAM-g6-simulated ones at the Japanese urban areas such as Nagoya (136.97°E, 35.17°N) and Osaka (135.54°E, 34.68°N).”

[C22] Line 598: An evaluation of the prescribed oxidants should be able to be done by sensitivity analyses described in the next section 3.2.

[A22] Thank you for your suggestion. We have inserted this point to our answer [A25]. As your suggested, we have moved this sentence to the next section.

[C23] Line 608: Only dry deposition? What about wet deposition?

[A23] Thank you for your comment. Here, we would like to mention that most EC is mainly scavenged through the wet deposition, whereas SO<sub>2</sub> is scavenged through both the dry and wet depositions as well as oxidations. Therefore, we have modified the last sentence as follows; “Although EC is also a primary product, the horizontal distributions of NICAM-g6str-simulated EC are larger than those of NICAM-g6str-simulated SO<sub>2</sub>, possibly because EC is less scavenged through the dry deposition and oxidation processes compared to SO<sub>2</sub>.”

[C24] Line 617: The doubled amount of SO<sub>2</sub> emissions can overcome the slight underestimation of the simulated sulfate compared with the observations. Therefore, the emission inventories of SO<sub>2</sub> should be improved for the better simulation of the sulfate. On the other hand, the results obtained by the sensitivity experiments of twice strength remain underestimated compared with the measurements. Then, what is a possible solution? Do following sentences are also indicating the emission inventories should be improved?

[A24] Thank you for your comments. The first comments ‘The doubled amount of SO<sub>2</sub> emissions can overcome ...’ is for SO<sub>2</sub> and sulfate in Line 633 of the revised manuscript, whereas the second comments ‘On the other hand, the results obtained by the sensitivity

experiments ...' is for EC in Line 617 of the revised manuscript. The inventories and sources of the EC and SO<sub>2</sub> are different, so we think each inventory should be improved by different ways, which was mentioned in the revised manuscript.

[C25] Line 644: How about effects of prescribed oxidants on hourly variations in this sensitivity analyses? The sentence in Line 537 has implied that prescribed oxidants cause the discrepancy of the hourly variations.

[A25] Thank you for your suggestions. As we mentioned in the revised manuscript, the relationship between the oxidants and the sulfate concentrations through the feedbacks is non-linear and complex. Therefore, the effects of the prescribed oxidants on hourly variations cannot be ignored. Of-course, we need to investigate the differences in the simulated sulfate concentrations with online-calculated and oxidants, but this investigation is beyond our present study (and will be the future study for us). In the revised manuscript, we have modified this sentence as follow; “These results and Figures 14 and 15 also suggest that the use of the prescribed oxidants for sulfate formation is not crucial for predicting monthly- and weekly-averaged sulfate mass concentrations at least by taking into account for diurnal and seasonal variations of the prescribed oxidants. At the same time, they also suggest that because the relationship between the oxidants and the sulfate concentrations through the feedbacks is non-linear and complex, the use of the prescribed oxidants for sulfate formation can affect the hourly variations of the sulfate concentrations, and thus the sensitivity of the oxidants to the simulated sulfate should be investigated.”

[C26] Line 654: An explanation of different Y-axes for observed and simulated values in Figure 17 should be added here, too.

[A26] Thank you for your suggestion. We have added the following comments to the revised manuscript; “using different Y-axes for the observed and simulated values”

[C27] Line 661: What is expected to show here by using the ratios of daytime and nighttime?

[A27] Thank you for your comment. We intended to compare the strength of the diurnal variation of the PM<sub>2.5</sub> using the simulations and observations. The ratios of daytime to nighttime could be an indicator of SOA contribution to the total PM<sub>2.5</sub>. We have modified the related sentences of the revised manuscript as follows; “As for the diurnal variation, the results show that the NICAM-g6str-simulated ratios (0.9-1.3) are larger than NICAM-g6-simulated ones (0.8-0.9), whereas the NICAM-g6str-simulated ones are smaller than the observed values (1.0-1.8). At Maebashi, where the ratio is higher than that at other sites, the issue of the poor model performance of the meteorological fields can be a major reason of the large underestimation, as mentioned in section 3.1. At all sites, especially Maebashi and Kisai, the possible underestimation of SOA may be a critical issue, as shown in the fact that the clear diurnal variation of PM<sub>2.5</sub> during

[August 4-9 and the high value of the ratios of daytime to nighttime](#) and suggested by previous studies (Matsui et al., 2009; Morino et al., 2010c). Morino et al. (2010c) ...”

[C28] References: Following references do not appear in the main text. “Carmichael et al., 2009; Chung et al., 2009; Koch et al., 2007; Lamarque et al., 2010; Moss et al., 2010; Ueda et al., 2009; Watanabe et al., 2010”

[A28] Thank you for your corrections. We have removed them from the references.

[C29] Figure 1: Where are 2 Sites (LIDAR measurements)? I could not find them in this figure.

[C30] Figure 14: It is better to insert R and Br within this figure.

[C31] Figure 15: A range of the color bar for SO<sub>2</sub> should be changed to see gradients more clearly.

[A29&A30&A31] Thank you for your suggestions and we have modified them in the revised manuscript.

1 **Application of a global nonhydrostatic model with a**  
2 **stretched-grid system to regional aerosol simulations**  
3 **around Japan**

4

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24 **Abstract**

25 An aerosol-coupled global nonhydrostatic model with a stretched-grid system has been  
26 developed. Circulations over the global and target domains are simulated with a single  
27 model, which includes fine meshes covering the target region to calculate meso-scale  
28 circulations. The stretched global model involves lower computational costs to simulate  
29 atmospheric aerosols with fine horizontal resolutions compared with a global uniform  
30 nonhydrostatic model, whereas it may require higher computational costs compared  
31 with the general regional models, because the stretched-grid system calculates inside  
32 and outside the target domain. As opposed to general regional models, the  
33 stretched-grid system does require neither a nesting technique nor lateral boundary  
34 conditions. In this study, we developed a new-type regional model for the simulation of  
35 aerosols over Japan, especially in the Kanto areas surrounding Tokyo, with a maximum  
36 horizontal resolution of approximately 10 km. This model usually reproduces temporal  
37 variations and their averages of the observed weather around Japan. This model  
38 generally reproduces monthly mean distributions of the observed sulfate and SO<sub>2</sub> over  
39 East Asia, with the high correlations ( $R>0.6$ ), but the underestimation of the simulated  
40 concentrations by 40% (sulfate) and 50% (SO<sub>2</sub>). Their underestimation of the simulated  
41 sulfate and SO<sub>2</sub> concentrations over East Asia are strongly affected by their  
42 underestimation in China and possibly by the uncertainty of the simulated precipitation  
43 around Japan. In the Kanto area, this model succeeds in simulating the wind patterns  
44 and the diurnal transitions around the center of the Kanto area, although it is inadequate  
45 to simulate the wind patterns and the diurnal transitions at some sites located at the edge

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50 of the Kanto area and surrounded on three sides by mountains, e.g., Maebashi, mainly  
51 due to the insufficient horizontal resolution. This model also generally reproduces both  
52 diurnal and synoptic variations of the observed and/or a regional aerosol-transport  
53 model, WRF-CMAQ, simulated EC, sulfate, and SO<sub>2</sub> concentrations in the Kanto area,  
54 especially with their high correlation ( $R>0.5$ ) at Komae/Tokyo. Although the aerosol  
55 module used in this study is relatively simplified compared to the general regional  
56 aerosol models, this study reveals that our proposed model with the stretched-grid  
57 system can be applicable for the regional aerosol simulation.

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## 59 **1 Introduction**

60 Aerosols can greatly affect regional air quality and contribute to global climate change  
61 (Forster et al., 2007). Recently, transboundary aerosol pollution, whereby regions  
62 beyond a given country's borders are affected by the aerosols generated in that country,  
63 has been of increasing concern (Ramanathan et al., 2008; Yu et al., 2012). The ongoing  
64 rapid economic growth in developing countries has the potential to exacerbate this issue  
65 (UNEP and WMO, 2011). Air pollution generated by aerosols is a critical public health  
66 issue due to the deleterious effects of these particles on human health (Dockery et al.,  
67 1993; Pope et al., 2009). Aerosols, which scatter and absorb solar radiation and act as  
68 cloud condensation nuclei, can directly and indirectly change the Earth's radiation  
69 budget. The majority of aerosols are emitted from localized areas, which are referred to  
70 as hotspots, such as megacities and biomass-burning regions, and are spread throughout  
71 the world via atmospheric transport (e.g., Ramanathan et al., 2008). Therefore, global  
72 aerosol-transport models should consider the important regional-scale characteristics of  
73 aerosol hotspots to reliably estimate their impacts on air quality and climate change.

74 Most existing global aerosol-transport models do not address the spatial variability of  
75 aerosols in the vicinity of hotspots due to their coarse horizontal resolution of 100–300  
76 km (Kinne et al., 2006; Textor et al., 2006). In addition, global aerosol-transport models  
77 with coarse resolutions frequently adopt a spectral transform method with a hydrostatic  
78 approximation to effectively calculate atmospheric dynamics. This spectral transform  
79 method is less effective than the grid-point method (Stuhne and Peltier, 1996; Taylor et

80 al., 1997; Randall et al., 2000) for high horizontal resolutions (Tomita et al., 2008).  
81 Models that employ the grid-point method flexibly define grid points to enable an  
82 adaptive focus on study regions. Thus, global models based on the grid-point method  
83 seem most appropriate for use in simulating aerosol transport from hotspots to outflow  
84 regions.

85 For this purpose, we utilized the global Nonhydrostatic Icosahedral Atmospheric Model  
86 (NICAM) developed by Tomita and Satoh (2004) and Satoh et al. (2008). NICAM has  
87 been employed for the global simulation of atmospheric processes with high-resolution  
88 grid spacing, whose size is comparable to the typical deep convective cloud scale.  
89 Miura et al. (2007) performed a one-week computation with a horizontal resolution of  
90 3.5 km using the Earth Simulator at the Japan Agency for Marine-Earth Science and  
91 Technology (JAMSTEC) to successfully simulate a Madden-Julian Oscillation (MJO)  
92 event. Suzuki et al. (2008) implemented an aerosol transport model named the Spectral  
93 Radiation-Transport Model for Aerosol Species (SPRINTARS; Takemura et al., 2005)  
94 in NICAM (we refer to this aerosol-coupled model as NICAM-SPRINTARS) and  
95 performed a one-week simulation with a horizontal resolution of 7 km using the Earth  
96 Simulator. Although these global, highly resolved calculations are promising with  
97 regard to long-term climate simulations for decades, their requirement of vast computer  
98 resources substantially limits their use in short-duration and/or case-specific simulations  
99 due to the current limitations of computational resources. To overcome this limitation,  
100 we adopt a compromise approach based on a new grid transformation named the

101 stretched grid system, which was developed and implemented in NICAM by Tomita  
102 (2008a) for computationally effective simulations in the target region (see, also, Satoh  
103 et al. 2010). We applied this approach to NICAM-SPRINTARS, which we named  
104 Stretch-NICAM-SPRINTARS, to calculate aerosol transport processes with high  
105 horizontal resolutions over aerosol source regions.

106 In this study, we focused on Japan, especially the Kanto region surrounding Tokyo  
107 (Figure 1), because the Kanto region living more than 30 million people is one of the  
108 largest megacities in the world. In Japan, a monitoring system for the air pollution, e.g.,  
109 PM<sub>2.5</sub> (aerosol particles with diameters less than 2.5  $\mu\text{m}$ ) and SO<sub>2</sub>, has been operated  
110 by the Japanese government. Inorganic ions, mainly sulfate, have been measured over  
111 Japan and other Asian countries under EANET (Acid Deposition Monitoring Network  
112 in East Asia; <http://www.eanet.asia/index.html>). Measurements of carbonaceous  
113 aerosols were limited, with the exception of intensive measurements (Fine Aerosol  
114 Measurement and Modeling in Kanto Area, FAMIKA) in the Kanto region during  
115 summer 2007 (Hasegawa et al., 2008; Fushimi et al., 2011). For the model evaluation  
116 using these measurements, we simulated aerosol spatial distributions during August  
117 2007 using Stretch-NICAM-SPRINTARS with a horizontal resolution of approximately  
118 10 km over the Kanto region. Because the model framework of  
119 Stretch-NICAM-SPRINTARS is identical to that of globally uniform grid simulation  
120 (we named it Global-NICAM-SPRINTARS), with the exception of the grid  
121 configuration, and involves lower computational costs than global simulations, the

122 investigation of the model performance of Stretch-NICAM-SPRINTARS can be simply  
123 and effectively extended to improve the original NICAM-SPRINTARS with globally  
124 uniform high resolution for near-future simulations. To evaluate aerosol simulations  
125 with the stretched-grid system, in this study we also conducted  
126 Global-NICAM-SPRINTARS, but with relatively low resolution (approximately 100  
127 km) due to the limited computational resources. The model intra-comparison approach,  
128 with the exception of the grid system and the spatial resolution, is very meaningful to  
129 investigate impacts of the stretched-grid system on the aerosol simulations. In addition,  
130 Stretch-NICAM-SPRINTARS can be a new-type model that is also applicable for a  
131 regional simulation of aerosols, because it focuses on a specific regional domain  
132 without require a nesting technique nor boundary conditions, unlike general regional  
133 models.

134 For the model evaluation in the target Japan, we mainly focused on a representative  
135 primary aerosol, i.e., elemental carbon (EC), and a representative secondary aerosol, i.e.,  
136 sulfate. EC is directly emitted from anthropogenic combustion processes, and is a good  
137 indicator to monitor the transport pattern. The global and regional modelings for sulfate,  
138 which is formed from SO<sub>2</sub> in the atmosphere, are more deeply understood compared to  
139 modelings for the other secondary aerosols such as nitrate and organic aerosols (e.g.,  
140 Barrie et al., 2001; Holloway et al., 2008; Hallquist et al., 2009; Morino et al., 2010a,  
141 2010b). In addition, sulfate is the largest contributor to the total secondary inorganic  
142 aerosols (e.g., Zhang et al., 2007), and the sulfate mass concentrations are larger than

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**削除:** Although the model inter-comparison using different modules coupled to different dynamic cores cannot clarify the reasons of the difference in the results among the models (e.g., Textor et al., 2006), t

147 that the nitrate ones in August 2007 over the Kanto area (Morino et al., 2010c).  
148 Originally, these basic components (EC and sulfate) are suitable for the evaluation in  
149 this study, primarily because the stretched-grid system was applied to the simulations of  
150 atmospheric pollutants over the land in the mid-latitude band for the first time and  
151 secondly because the original SPRINTARS is more simplified compared to  
152 conventional regional aerosol models.

153 This paper is organized as follows: the model framework of NICAM and SPRINTARS  
154 and the experimental design are described in Section 2. We show two model results; (1)  
155 Stretch-NICAM-SPRINTARS with glevel-6, in which “glevel” is the number of  
156 divisions of an icosahedron used to construct the horizontal grid, (hereafter referred to  
157 as the “NICAM-g6str” model) and (2) Global-NICAM-SPRINTARS with glevel-6  
158 (hereafter referred to as the “NICAM-g6” model). In Section 3, the model results are  
159 validated using in-situ measurements in terms of meteorological fields including  
160 precipitation and aerosol species, especially EC, sulfate and SO<sub>2</sub>. For the model  
161 evaluation of chemical species, we also made use of results in a regional aerosol model,  
162 the Community Multiscale Air Quality (CMAQ) driven by the Weather Research and  
163 Forecasting (WRF) model named WRF-CMAQ, shown by Shimadera et al. (2013). We  
164 also present the validation of total aerosol amounts, i.e., PM<sub>2.5</sub>, and aerosol optical  
165 product, i.e., extinction for spherical aerosols. Finally, the conclusions are summarized  
166 in Section 4.

167

168 **2 Model description**

169 **2.1 Nonhydrostatic Icosahedral Atmospheric Model (NICAM)**

170 NICAM, which employs an icosahedral grid-point method with a nonhydrostatic  
171 equation system (Tomita and Satoh, 2004; Satoh et al., 2008, 2014), is run with a  
172 maximum horizontal resolution of 3.5 km (Tomita et al. 2005; Miura et al., 2007) and  
173 can be applied to a transport model of aerosols and gases as a conventional atmospheric  
174 general circulation model (Suzuki et al., 2008; Niwa et al., 2011; Dai et al., 2014a,  
175 2014b; Goto, 2014). NICAM can also be employed for regional-scale simulations by  
176 adopting a stretched-grid system (Tomita, 2008a; Satoh et al., 2010). The stretched  
177 icosahedral grid was developed from a general grid transformation method, i.e., the  
178 Schmidt transformation method, for a horizontal grid system on a sphere. In the  
179 Schmidt transformation, the grid interval on a sphere lacks uniformity with a finer  
180 horizontal resolution close to the center of the target region. Tomita (2008a) showed  
181 that the Schmidt transformation minimizes potential errors involving the isotropy and  
182 homogeneity of the target region. The stretched-grid system can solve the main  
183 problems associated with commonly used regional models, which occur from artificial  
184 perturbations near boundary areas in cases where meteorological and aerosol fields are  
185 prescribed. In addition, the computational cost of the stretched-grid system is  
186 substantially lower than that of a global calculation under the same horizontal resolution  
187 in the target region. For example, when the globally uniform grid with a maximum  
188 horizontal resolution of 10 km is applied to the global simulation, the minimum

189 required theoretical computational cost is 64-256 times higher than the cost of the  
190 stretched-grid system in this study. Compared to conventional regional models, the  
191 computational cost may increase because the stretched-grid system requires the  
192 calculation outside the target domain. Furthermore, the model framework of the  
193 stretched global model is identical to that of the uniformed global model without special  
194 modifications, whereas the model framework of regional models is usually different  
195 from that of global models. These advantages can facilitate additional developments for  
196 global simulations by testing a new scheme with minimal computational cost.  
197 Compared with general regional models, the stretched-grid system is more suitable for  
198 the future study, which aimed to extend its use to the global uniform high-resolution  
199 NICAM-SPRINTARS.

200 In this study, we adopt the stretched-grid system to focus on the Kanto region, including  
201 Tokyo, using glevel-6 resolution and the stretched ratio of 100 (we call it  
202 NICAM-g6str), which is the ratio of the largest horizontal grid spacing located on the  
203 opposite side of the earth from Tokyo to the smallest horizontal grid spacing near  
204 Tokyo. As a result, a minimum horizontal resolution of 11 km around the center  
205 (140.00°E, 35.00°N) was used. NICAM implements comprehensive physical processes  
206 of radiation, boundary layer and cloud microphysics. The radiation transfer model is  
207 implemented in NICAM with the k-distribution radiation scheme MSTRN, which  
208 incorporates scattering, absorption and emissivity by aerosol and cloud particles as well  
209 as absorption by gaseous compounds (Nakajima et al., 2000; Sekiguchi and Nakajima,

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211 2008). The vertical turbulent scheme comprises the level 2 scheme of turbulence closure  
212 by Mellor and Yamada (1974), Nakanishi and Niino (2004, 2009) and Noda et al.  
213 (2009). The cloud microphysics consist of the six-class single-moment bulk scheme  
214 (water vapor, cloud water, rain, cloud ice, snowflakes and graupel) (Tomita, 2008b).  
215 Based on our experience in previous studies, we did not employ cumulus  
216 parameterization in this study (e.g., Tomita et al., 2005; Sato et al., 2009; Nasuno, 2013).  
217 The topography used in this study is based on GTOPO30 (the horizontal resolution is 30  
218 arc seconds, that is approximately 1 km) courtesy of the U.S. Geological Survey. The  
219 vertical coordinates system adopts Lorenz grid and  $z^*$  (terrain-following) coordinates  
220 with the 40 layers of z-levels and model top of 40 km height (Satoh et al., 2008).The  
221 timestep was set to 20 seconds.

222

## 223 **2.2 SPRINTARS**

224 Based on the approach of Suzuki et al. (2008), the three-dimensional aerosol-transport  
225 model—Spectral Radiation-Transport Model for Aerosol Species (SPRINTARS;  
226 Takemura et al., 2000, 2002, 2005; Goto et al. 2011a,b,c)—was coupled to NICAM in  
227 this study. The SPRINTARS model calculates the mass mixing ratios of the primary  
228 tropospheric aerosols, i.e., carbonaceous aerosol (EC and OC, organic carbon), sulfate,  
229 soil dust, sea salt and the precursor gases of sulfate, namely,  $\text{SO}_2$  and dimethylsulfide  
230 (DMS). The aerosol module considers the following processes; emission, advection,  
231 diffusion, sulfur chemistry, wet deposition and dry deposition, including gravitational  
232 settling. For carbonaceous aerosols, the 50% mass of EC from fossil fuel sources is

233 composed of externally mixed particles, whereas other carbonaceous particles are  
234 emitted and treated as internal mixtures of EC and OC (EC-OC internal mixture).  
235 Biogenic secondary organic aerosols (SOAs) from terpenes are treated but are greatly  
236 simplified by multiplying a conversion factor to the terpenes emission (Takemura,  
237 2012). In addition, anthropogenic SOAs from toluene and xylene are disregarded in this  
238 study. The bulk mass concentrations of EC, OC, and sulfate are calculated by  
239 single-modal approach, which means that the SPRINTARS model does not explicitly  
240 treat aerosol dynamic processes such as coagulation and condensation. The particle size  
241 distribution of the dry particles are prescribed in a logarithmic normal size distribution  
242 with dry mode radii of 18, 100, 80 and 69.5 nm, for pure EC, EC-OC internal mixture,  
243 biogenic SOA and externally mixed sulfate, respectively (Goto et al., 2011a). The  
244 hygroscopicities, densities and refractive indices for the aerosols are set to the same  
245 values used by Takemura et al. (2002) and Goto et al. (2011a). The combinations of the  
246 pre-calculated cross-sections of the extinction and simulated mixing ratios for each  
247 aerosol species provide the simulated aerosol extinction coefficient for each timestep of  
248 the model (Takemura et al., 2002). The sulfur chemistry in SPRINTARS considers only  
249 three chemical reactions to form sulfate through gas-phase oxidation of  $\text{SO}_2$  by  
250 hydroxyl radical (OH) and aqueous-phase oxidation by ozone and hydrogen peroxide.  
251 The large part of  $\text{SO}_2$  are emitted from fossil fuel combustion, biomass burning, and  
252 volcano eruption, whereas some of  $\text{SO}_2$  are formed from the oxidation of DMS, which  
253 is emitted naturally from marine phytoplanktons. The numerical solution in the  
254 oxidations adopts an approximation in a quasi first-order reaction using the same

255 integrated time resolution as that of the dynamic core. The pH value in the  
256 aqueous-phase is fixed at 5.6, because the SPRINTARS model treats limited ions in the  
257 aqueous-phase (e.g., Takemura et al., 2000). The oxidant distributions (OH, ozone and  
258 hydrogen peroxide) were offline provided by a chemical transport model. The  
259 atmospheric removal of aerosols in SPRINTARS includes wet (due to rainout and  
260 washout) and dry (due to turbulence and gravity) deposition processes, whereas those of  
261 SO<sub>2</sub> only include rainout and dry deposition by turbulence. In the cloudy grid, the mass  
262 fractions of sulfate out of the cloud droplets to the mass of sulfate in the grid were fixed  
263 at 0.5, whereas the fractions for SO<sub>2</sub> were determined by Henry's law (Takemura et al.,  
264 2002). As for pure EC, EC-OC internal mixture, and biogenic SOA, the mass fractions  
265 were fixed at 0.1, 0.3, and 0.3, respectively. Because the SPRINTARS model does not  
266 predict the mass mixing ratio of the chemical tracers inside the clouds, it assumes that  
267 the tracers inside the clouds are evaporated from the clouds at one timestep. In this  
268 study, the particle mass concentrations for diameters less than 2.5 μm (defined as  
269 PM<sub>2.5</sub>) are calculated by summing EC, organic matter by multiplying OC by 1.6  
270 (Turpin and Lim, 2001), sulfate and ammonium aerosols. Because this model cannot  
271 directly predict ammonium compounds, it is assumed that all sulfate is the form of  
272 ammonium sulfate, so that their concentration was estimated by multiplying the mass  
273 concentration of sulfate by 0.27, which is the molar ratio of ammonium ion to  
274 ammonium sulfate. The nitrate in this study is disregarded, primarily because the main  
275 objective in this study is modeling of sulfate as a representative secondary aerosols and  
276 secondly because the nitrate mass concentrations are lower than the sulfate ones with

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278 the target of August 2007 in Japan (Morino et al., 2010c).

279

### 280 **2.3 Design of the experiments**

281 The target period comprises one month in August 2007, in which an intensive  
282 measurement of aerosol chemical species was conducted under Project FAMIKA  
283 (Hasegawa et al., 2008; Fushimi et al., 2011). The six-hour meteorological fields (wind  
284 and temperature) were nudged above a height of 2 km using NCEP-FNL reanalysis data  
285 (<http://rda.ucar.edu/datasets/ds083.2/>). The one-day sea surface temperature was also  
286 nudged using the NCEP-FNL data. The initial conditions were prescribed by the  
287 NCEP-FNL data for the meteorological fields and the one and a half months spinup  
288 results of the Stretch-NICAM-SPRINTARS model for the aerosol fields, respectively.

289 The emission inventories of anthropogenic EC, OC and SO<sub>2</sub> in this experiment were  
290 prepared by EAGrid2000 with a horizontal resolution of 1 km over Japan (Kannari et al.,  
291 2007), REAS version 2 with a horizontal resolution of 0.25° over Asia (Kurokawa et al.,  
292 2013) and the AeroCom inventory with a horizontal resolution of 1° over other areas of  
293 the world (Diehl et al., 2012). Because EAGrid2000 does not explicitly estimate EC and  
294 OC inventories, we estimated the inventories to be consistent with those from previous  
295 studies (Morino et al., 2010a,b; Chatani et al., 2011) by modifying the PM<sub>2.5</sub> inventory  
296 of EAGrid2000 using scaling factors of EC/PM<sub>2.5</sub> and OC/PM<sub>2.5</sub> based on sources.  
297 These inventories of anthropogenic EC and SO<sub>2</sub> in 2007 are described in Figure 2. The  
298 emissions of SO<sub>2</sub> from volcanoes in Japan, such as Miyakejima and Sakura-jima, were

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300 obtained from statistical reports (<http://www.seisvol.kishou.go.jp/tokyo/volcano.html>)  
301 by the Japan Meteorological Agency (JMA). In this study, the distributions of three  
302 hourly averaged monthly oxidants (OH, ozone and hydrogen peroxide) were derived  
303 from a global chemical transport model (CHASER) coupled to the Model for  
304 Interdisciplinary Research on Climate (MIROC), named MIROC-CHASER, with the  
305 spatial resolution of 2.8° by 2.8° (Sudo et al., 2002).

306 To evaluate model performances in the stretched-grid system, we also simulated  
307 NICAM-SPRINTARS with the globally uniformed grid simulation in glevel-6  
308 resolution (the horizontal resolution is set to 110 km and we call it NICAM-g6).

309 Global-NICAM-SPRINTARS with relatively low resolution has been applied to aerosol  
310 simulations and well compared with in-situ measurements and satellite remote sensing

311 (Dai et al., 2014a; Goto, 2014). Apart from the NICAM-g6str simulation, in the

312 NICAM-g6 simulation, the cloud physics apply both the prognostic  
313 Arakawa-Schubert-type cumulus convection scheme (Arakawa and Schubert, 1974) and  
314 the diagnostic large-scale clouds described by Le Treut and Li (1991). The large-scale  
315 cloud module is based on single moment bulk scheme for cloud mixing ratio. The  
316 precipitation rate is parameterized by Berry (1967). Except for the grid system and the  
317 horizontal resolution (which determines the module of the cloud physics),  
318 Global-NICAM-SPRINTARS was identical to Stretch-NICAM-SPRINTARS.

319 Therefore, the comparison between NICAM-g6str and NICAM-g6 led to clarify  
320 impacts of the horizontal resolution on the aerosol distribution.

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325

## 326 **2.4 Observation**

327 In this study, we focused on the aerosol chemical component of EC as the primary  
328 particle and sulfate as the secondary particle. To evaluate the model results over the  
329 Kanto region, we used observations of the surface mass concentrations of EC and  
330 sulfate in four cities under Project FAMIKA: Maebashi/Gunma (139.10°E, 36.40°N),  
331 Kisai/Saitama (139.56°E, 36.09°N), Komae/Tokyo (139.58°E, 35.64°N) and  
332 Tsukuba/Ibaraki (140.12°E, 36.05°N). The EC particles in PM<sub>2.5</sub> were collected every  
333 six hours with quartz fiber filters and analyzed with the thermal/optical method  
334 according to the IMPROVE protocol (Chow et al., 2001). The sulfate particles in PM<sub>2.5</sub>  
335 were also collected every six hours with Teflon filters and analyzed by ion  
336 chromatography. In addition to the limited FAMIKA dataset, we utilized measurements  
337 taken by the EANET (Acid Deposition Monitoring Network in East Asia;  
338 <http://www.eanet.asia/index.html>) and the 4th national survey report of acid rain over  
339 Japan in fiscal year 2007  
340 ([http://tenbou.nies.go.jp/science/institute/region/journal/JELA\\_3403041\\_2009.pdf](http://tenbou.nies.go.jp/science/institute/region/journal/JELA_3403041_2009.pdf)) to  
341 assess the monthly mean concentrations of sulfate and SO<sub>2</sub> at Japanese and Korean sites.  
342 We also obtained Chinese measurements by *Zhang et al.* [2012], as part of the Chinese  
343 Meteorological Administration Atmosphere Watch Network (CAWNET). To validate  
344 the concentration of SO<sub>2</sub> for the Kanto region, we accessed monitoring stations operated  
345 by Japanese and local governments.

346 In the validation of the meteorological fields simulated by NICAM-g6str and  
347 NICAM-g6, we used meteorological fields (wind and temperature) reanalyzed by  
348 NCEP-FNL over East Asia. In the Kanto region, we obtained measurements for the  
349 meteorological parameters (temperature, relative humidity (RH) and wind) at or near  
350 the 7 sites of Project FAMIKA and additional cities: Tsuchiura/Ibaraki (140.20°E,  
351 36.07°N), which is the city nearest to Tsukuba; Yokohama/Kanagawa (139.64°E,  
352 35.45°N); Chiba/Chiba (140.12°E, 35.62°N); Adachi/Tokyo (139.82°E, 35.77°N); and  
353 Machida/Tokyo (139.43°E, 35.53°N), which is the city nearest to Komae, as shown in  
354 Figure 1(b). For precipitation, we used a measurement taken by the Automated  
355 Meteorological Data Acquisition System (AMeDAS) at 21 sites over Japan including  
356 the following 10 Kanto's sites: Yokohama; Chiba; Tsukuba; Tokyo, which is near  
357 Adachi; Maebashi; Huchu, which is near Machida; Konosu, which is near Kisai; Abiko  
358 (140.11°E, 35.60°N); Saitama (139.59°E, 35.88°N); and Nerima (139.59°E, 35.74°N)  
359 (Figure 1). To evaluate the spatial patterns of the precipitation obtained by  
360 NICAM-g6str and NICAM-g6, we used the quantities of the monthly mean  
361 precipitation around Japan that were derived from the Global Satellite Mapping of  
362 Precipitation (GSMaP; Okamoto et al., 2005; Kubota et al., 2007; Aonashi et al., 2009;  
363 Ushio et al., 2009) and the Meso Scale Model (MSM) developed by the JMA for rain  
364 forecast (Saito et al., 2006). The results by MSM are generally higher accurate than  
365 those in GSMaP, although the covering area in MSM is limited around Japan.

366 To evaluate the quantities of the total aerosol amounts, such as PM<sub>2.5</sub>, we compared the

367 simulated PM<sub>2.5</sub> concentrations with the observations at the 18 sites including the  
368 FAMIKA sites and other monitoring stations operated by the Japanese and local  
369 governments (Figure 1). The PM<sub>2.5</sub> concentrations were continuously observed using  
370 tapered element oscillating microbalance (TEOM) with Series 1400a Ambient  
371 Particulate Monitor. The instruments are controlled under the temperature of 50 °C, to  
372 minimize the influence of change in the ambient temperature and RH. However, it  
373 includes large uncertain due to the difficulty in completely eliminate the water content  
374 attached to aerosols and lacks of the calibration of the instrument in some of sites.  
375 Nevertheless, the observed PM<sub>2.5</sub> concentrations with hourly time resolution were still  
376 useful to validate the model results.

377 In Tsukuba and Chiba, light detection and ranging (LIDAR) measurements operated by  
378 the National Institute for Environmental Studies (NIES) of Japan were also available  
379 (Sugimoto et al., 2003; Shimizu et al., 2004). The LIDAR unit measured vertical  
380 profiles of the backscattering intensity at 532 and 1064 nm and the depolarization ratio  
381 at 532 nm. The backscattering intensity was converted to the extinction coefficient, and  
382 the depolarization ratio distinguished the extinction between spherical and non-spherical  
383 particles. In this study, we only used vertical profiles of the extinction for spherical  
384 particles. A detailed algorithm was provided by Sugimoto et al. (2003) and Shimizu et  
385 al. (2004).

386

### 387 **3 Validation of Stretch-NICAM-SPRINTARS**

388 **3.1 Meteorological fields**

389 So far, the stretched-grid system was mainly applied to the simulations of tropical  
390 cyclones or tropical convective clouds with small domain over oceans for the short-term  
391 period (less than several days) (e.g., Satoh et al., 2010; Arakane et al., 2014). In this  
392 study, we focused on the air pollution around Japan (for the longer period). Therefore,  
393 we first focused on the general circulation of the basic meteorological fields over the  
394 large domain, which can affect the air pollution over Japan. Figure 3 shows temperature  
395 and winds near the surface and the model height of approximately 5 km over Asia  
396 region (100°E-170°E, 10°N-50°N). In August, North Pacific High (or Ogasawasa High)  
397 mainly brings clear weather around Japan. A frequency of the precipitation is usually  
398 limited, but a total amount of the monthly mean precipitation is not small, because of  
399 typhoons and shower rain. In the focusing region, the general meteorological fields  
400 simulated by NICAM-g6str and NICAM-g6 are comparable to those obtained by  
401 NCEP-FNL. The absolute biases in the temperature between NICAM-g6str and  
402 NCEP-FNL or between NICAM-g6 and NCEP-FNL are within 1.5 °C at the surface  
403 and the height of 5 km. Around the Japanese Alps, however, the  
404 NICAM-g6str-simulated temperature is lower than the NCEP-FNL-estimated one by at  
405 most 2.5 °C, because of the differences in the resolved topography due to the different  
406 spatial resolution between NICAM-g6str and NCEP-FNL. As for wind, western winds  
407 over the northeastern part of Japan in both NICAM-g6str and NICAM-g6 are stronger  
408 compared to those in NCEP-FNL. With the exception of this bias, the performances of

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NICAM-g6str-simulated temperature tends to be  
larger than NICAM-g6-simulated one by at most  
3 °C, probably because the spatial resolution in  
NICAM-g6str is finer than that in NCEP-FNL.  
These positive biases between NICAM-g6str and  
NCEP-FNL can be seen around Japan.

418 both NICAM-g6str and NICAM-g6 are good. Therefore, it is concluded that the  
419 stretched-grid systems does not affect the general circulations under the nudging  
420 technique in this study.

421 To evaluate the model performances of the six-hourly mean concentrations of aerosol  
422 chemical species and SO<sub>2</sub> over the main target region, i.e., Kanto area, we used the  
423 six-hourly instant observations of temperature, RH, wind and precipitation at each  
424 station over the Kanto area shown in Figure 1. The results and summary are shown in  
425 Figures 4 to 7 and Table 1. The NICAM-g6 results, especially in terms of diurnal  
426 variations, tend to be far from the observations compared to the NICAM-g6str results,  
427 because NICAM-g6, with the horizontal resolution of approximately 100 km, does not  
428 fully resolve the topology over the Kanto area. Figure 4 illustrates the temporal  
429 variations of temperature at a height of 2 m. The temporal variations in the  
430 NICAM-g6str-simulated temperature are generally comparable to those in the observed  
431 temperatures with root-mean-square-error (RMSE) values of less than 3°C, with the  
432 exception of the results obtained for Maebashi and Machida. At these two sites, the  
433 mean values of the NICAM-g6str-simulated temperatures are lower than those of the  
434 observed temperatures by a maximum of 3.6°C. The correlation coefficients (R)  
435 between NICAM-g6str and the observation range from 0.7–0.9, whereas the R between  
436 NICAM-g6 and the observation range from 0.7–0.8, as shown in Table 1. Figure 5  
437 shows the temporal variations in RH at a height of 2 m. The temporal variations in the  
438 NICAM-g6str-simulated RH are similar to the observations, with the RMSEs in the

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uniformed grid systems are well reproduced

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442 range of 10–15%. In contrast, the NICAM-g6-simulated RH is overestimated compared  
443 to the observations, with the RMSEs in the range of 16–26%. The R values of RH  
444 between the simulation (both NICAM-g6str and NICAM-g6) and observations are  
445 approximately 0.6–0.8 (Table 1).

446 The temporal variations in the wind direction and speed simulated by NICAM-g6str are  
447 compared with the observations in Figures 6 and 7. Near the southern part of the Kanto  
448 area (Yokohama, Tsuchiura, Adachi and Machida), with the exception of Chiba, the  
449 NICAM-g6str-simulated wind direction is generally comparable to the observations,  
450 with a slight overestimation of the both NICAM-g6str and NICAM-g6 simulated wind  
451 speed compared with the observations. At these four sites, the R and RMSE values in  
452 NICAM-g6str range from approximately 0.5–0.7 and approximately 1.7–2.3 m/s,  
453 respectively. In Chiba located near the ocean, the R value of wind speed between  
454 NICAM-g6str and the observation is 0.41, whereas the NICAM-g6str-simulated wind  
455 directions generally agree with the observations. Conversely, at Maebashi and Kisai, the  
456 daily variations in the both NICAM-g6str and NICAM-g6 simulated wind directions  
457 differ significantly from those in the observations, in which the southern winds and  
458 northern winds frequently occur during the day and night, respectively, for example,  
459 during August 5–12. At these two sites, the NICAM-g6-simulated wind direction and  
460 speed is not closer to the observations compared to those obtained by NICAM-g6str.  
461 The R value for wind speed between the NICAM-g6str and the observations at these  
462 sites is estimated to be approximately 0.2. The observed southeasterly wind is long sea

463 breeze toward Maebashi Plateau surrounded on three sides by mountains around  
464 Maebashi. The observed winds are caused by daytime meso-scale thermal lows  
465 developed over the central Japan covering the Japanese Alps (Kuwagata and Sumioka,  
466 1991). The Japanese Alps with the highest terrain in Japan can affect the local  
467 meteorological fields even around 100-200 km away (Kitada et al., 1998). Therefore, it  
468 suggests that the horizontal resolution in this study using NICAM-g6str (10 km over the  
469 Kanto area) does not fully resolve the complex terrains of the Japanese Alps and the  
470 Maebashi plateau. Therefore, it suggests that it is inadequate to simulate the wind  
471 patterns and the diurnal transitions near high mountains around the Kanto area, whereas  
472 it is adequate to simulate them around the center of the Kanto area.

473 Figures 8-10 show comparisons of NICAM-g6str and NICAM-g6 simulated  
474 precipitation with the observations. Figure 8 compares the simulated precipitation with  
475 the MSM and GSMaP derived results. During the early August 2007, mainly due to  
476 passing of a typhoon over the western Japan, Okinawa, and Korea, the August mean  
477 precipitation in the western Japan is larger than that in the eastern Japan, especially the  
478 Kanto area. The monthly mean precipitation is estimated to be more than 200  
479 mm/month over the western Japan, whereas that is estimated to be less than 50  
480 mm/month over the eastern Japan. The horizontal patterns of the precipitation obtained  
481 by NICAM-g6str in East China Sea, Sea of Japan near the Japan coast, and Korea are  
482 closer to those derived from MSM and GSMaP than those obtained by NICAM-g6. In  
483 the Kanto area, however, the NICAM-g6str-simulated precipitation with the range of

484 50-200 mm/month is overestimated compared to the MSM and GSMaP results. The  
485 NICAM-g6-simulated precipitation over the Kanto area with the range of 100-200  
486 mm/month is also much overestimated. In Figure 9 showing the temporal variations in  
487 the amount of precipitation per day at 21 Japanese sites, the observed precipitation is  
488 extremely limited during August 7-19 in the Kanto area. In other regions, the magnitude  
489 of the precipitation is strong, although the precipitation is sporadic. In terms of the  
490 frequency of the precipitation, the NICAM-g6str performance is better than the  
491 NICAM-g6 one. Figure 10 illustrates the predictive value of daily precipitation, defined  
492 as the ratio of the number of days where the model correctly predicts the weather (less  
493 than 1 mm/day or more than 1 mm/day) to the number of the whole days. In the  
494 NICAM-g6str results, the predictive values at most of sites over the Kanto area and four  
495 sites over the non-Kanto area such as Nagoya and Osaka are calculated to be more than  
496 85%. The predictive values obtained by NICAM-g6str are mostly higher than those  
497 obtained by NICAM-g6. During the rainy days such as August 20, 22 and 23 over the  
498 Kanto area, both NICAM-g6str and NICAM-g6 capture the precipitation, whereas  
499 NICAM-g6str reproduces greater amounts of the precipitation and NICAM-g6  
500 reproduces longer periods and larger areas compared to the observations. NICAM-g6str  
501 does not always capture a sudden shower, as general meteorological models have  
502 difficulties in predicting this type of precipitation system (e.g., Kawabata et al., 2011).  
503 To increase the accuracy of such precipitation, more sophisticated cloud-microphysics  
504 model, e.g., NICAM-NDW6 model proposed by Seiki and Nakajima (2014) based on  
505 the double-moment bulk scheme with six water categories, may be required. In the

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507 western Japan, during the rainy days, e.g., August 22-23, both NICAM-g6str and  
508 NICAM-g6 usually capture large-scaled precipitation (Figure 9). Overall, NICAM-g6str  
509 usually reproduces the observed weather in the target regions and periods, whereas  
510 NICAM-g6 does not capture general feature such as the sporadic precipitation.

511

## 512 **3.2 Aerosol fields**

### 513 **3.2.1 Evaluation of chemical species**

514 Figures 11, 12, and 13 illustrates the temporal variations in the surface EC, sulfate, and  
515 SO<sub>2</sub> concentrations at the four stations (Maebashi, Kisai, Komae and Tsukuba) in the  
516 Kanto area using the simulations and the measurements. The simulations include  
517 NICAM-g6str, NICAM-g6, and the Community Multiscale Air Quality (CMAQ) driven  
518 by the Weather Research and Forecasting (WRF) model named WRF-CMAQ shown by  
519 their Figures 5 and 6 of Shimadera et al. (2013). Shimadera et al. (2013) calculated the  
520 WRF-CMAQ with a horizontal resolution of 5 km and an emission inventory that is  
521 similar to that in the present study. Table 2 summarizes the statistical parameters for the  
522 concentrations of EC, sulfate, and SO<sub>2</sub>. The temporal variation and the average of EC  
523 simulated by NICAM-g6str are better agreement with the observations obtained for  
524 Komae than those simulated by NICAM-g6 (Figure 11(c)). However, the averages of  
525 both NICAM-g6str and NICAM-g6 simulated EC concentrations at the other sites are  
526 much underestimated compared to the observations (Table 2). For Tsukuba shown in  
527 Figure 11(d), both the NICAM-g6str and NICAM-g6 simulated EC concentrations tend

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529 to be underestimated compared with the observed concentrations, especially during the  
530 daytime, even though the temporal variation of EC obtained by NICAM-g6str is closer  
531 to the observed one compared to those obtained by NICAM-g6. At Maebashi and Kisai,  
532 the temporal variation and the averages of EC obtained by NICAM-g6 are also  
533 underestimated compared with the observations by a factor of three to five.  
534 NICAM-g6str tends to have daily maximums of EC concentrations during the morning  
535 time, whereas NICAM-g6 tends to have daily maximums during the nighttime. The  
536 temporal variations of NICAM-g6str-simulated EC concentrations are generally  
537 comparable to those by WRF-CMAQ shown in Figure 11 and their Figure 3 of Chatani  
538 et al. (2014), with the exception of the results at Maebashi and Kisai where the EC  
539 concentrations obtained by NICAM-g6str are smaller than those obtained by  
540 WRF-CMAQ. At these sites, the difference in the EC concentrations between  
541 NICAM-g6str and WRF-CMAQ is probably caused by the difference in the horizontal  
542 resolution, which is most likely critical for properly simulating the air pollution  
543 delivered by the meteorological wind fields from the center of the Kanto region (Kusaka  
544 and Hayami, 2006). Table 2 also shows that the R obtained by NICAM-g6str at all sites  
545 are high or moderate, with the exception of Maebashi, whereas those obtained by  
546 NICAM-g6 and CMAQ are low. At most sites, the EC concentrations obtained by  
547 WRF-CMAQ shown in Figure 11, and WRF-CMAQ illustrated by Morino et al.  
548 (2010a,b) and Chatani et al. (2014), NICAM-g6str, and NICAM-g6 are also  
549 underestimated compared to the observations with the larger values of RSME. The  
550 underestimation of EC concentrations is investigated by sensitivity tests of EC emission

551 inventory in section 3.2.2.

552 At the same four sites, simulated sulfur components (sulfate and SO<sub>2</sub>) are compared  
553 with the observations in Figures 12 and 13. The observed SO<sub>2</sub> represents the ensemble  
554 results of monitoring stations operated by Japanese and local governments around each  
555 FAMIKA site. The mean differences in the sulfate mass concentrations between  
556 NICAM-g6str and the observations are within approximately 10% at Maebashi and  
557 Tsukuba, approximately -30% at Komae, and approximately +40% at Kisai. At all sites,  
558 the temporal variations of the NICAM-g6str-simulated sulfate concentrations are  
559 generally comparable to those obtained by the observations and WRF-CMAQ shown in  
560 Figure 12 (i.e., their Figure 6 of Shimadera et al., 2013) and illustrated in their Figure 3  
561 of Morino et al. (2010a), whereas the differences in the sulfate concentrations between  
562 NICAM-g6str and the observations are somewhat greater on August 7 and 8 at  
563 Maebashi where the performance of NICAM-g6str is relatively poor, mainly due to the  
564 inadequate horizontal resolution to reproduce the observed meteorological fields, as  
565 shown in section 3.1. The use of the prescribed distributions of three hourly averaged  
566 monthly oxidants may partly cause the discrepancy of the hourly variations of the  
567 sulfate between NICAM-g6str and the observations. The R obtained by all the models  
568 (NICAM-g6str, NICAM-g6, and WRF-CMAQ) is acceptable at most sites, with the  
569 exception of NICAM-g6str at Maebashi and WRF-CMAQ at Kisai. The RMSEs  
570 obtained by all the models are smaller at Komae and Tsukuba than those at Maebashi  
571 and Kisai. The six-hourly variations of the sulfate obtained by WRF-CMAQ are

572 sometimes missed by NICAM-g6str, partly due to the use of the prescribed oxidants.  
573 Even though NICAM-g6 reproduces the synoptic cycle of the observed sulfate, it has  
574 difficulties in simulating the diurnal cycle of the observed and NICAM-g6str-simulated  
575 sulfate, as shown in the results of EC by Figure 11. The averages of the sulfate  
576 concentrations obtained by NICAM-g6 tend to be smaller than those by NICAM-g6str  
577 and the observations. The possible impacts of the prescribed oxidant on the sulfate  
578 concentrations are investigated in section 3.2.2.

579 In Figure 13, NICAM-g6str and NICAM-g6 simulated SO<sub>2</sub> concentrations are  
580 compared by the observations. In the previous studies, the comparison in SO<sub>2</sub>  
581 concentrations between the simulation and observation was very limited, with the  
582 exception of their Figure 4 of Morino et al. (2010b), which showed large differences in  
583 the SO<sub>2</sub> concentrations between WRF-CMAQ and the observations by more than a  
584 factor of two. The R between NICAM-g6str and the observations are low, with the  
585 exception of Komae (R=0.62), but are approximately within the range obtained by  
586 WRF-CMAQ in Morino et al. (2010b). The differences in the mean SO<sub>2</sub> concentrations  
587 between NICAM-g6str and the observations and between NICAM-g6 and the  
588 observations are within approximately 20% at all sites, with the exception of  
589 NICAM-g6str at Maebashi and NICAM-g6str at Tsukuba (Table 2). The temporal  
590 variations in the simulated SO<sub>2</sub> concur with those in the observations. The observations  
591 sometimes show high SO<sub>2</sub> concentrations at all sites, e.g., up to 20 ppbv at Komae, in  
592 the afternoon on August 12 and 14. On August 12, NICAM-g6str normally reproduced

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594 the peaks of the observed SO<sub>2</sub>, but with the blunter and slightly shifted peaks. In the  
595 NICAM-g6str simulation, the strong SO<sub>2</sub> plumes from industrial sources over the Tokyo  
596 Bay arrived at the inner areas such as Kisai. On August 14, although the  
597 NICAM-g6str-simulated winds were comparable to the observed ones (Figures 7 and 8),  
598 NICAM-g6str did not reproduce the sharp peaks of the observed SO<sub>2</sub>, especially at  
599 Komae and Tsukuba. It may imply that special meteorological fields cause the observed  
600 peaks on August 12, whereas unaccounted SO<sub>2</sub> emission from local sources or sporadic  
601 volcanoes is stronger on August 14. The latter issue is improved by processing  
602 time-highly-resolved emission inventories of SO<sub>2</sub>, which can be estimated through a  
603 top-down approach using a data assimilation (Schutgens et al., 2012; Xu et al., 2013).

604 To assess the performance of both NICAM-g6str and NICAM-g6 in simulating the  
605 distributions of the air pollutants over Japan, we compared the August averages of the  
606 simulated EC, sulfate and SO<sub>2</sub> concentrations with the available measurements (Figures  
607 14 and 15). Although the EC observatories are limited, both the NICAM-g6str and  
608 NICAM-g6 simulated EC concentrations are much underestimated compared to the  
609 observations, with the relative bias (*Br*), defined as a ratio of the simulated value to the  
610 observed one, to be 0.15 (NICAM-g6str) and 0.16 (NICAM-g6). In China, the  
611 NICAM-g6str-simulated EC concentrations are comparable to the  
612 NICAM-g6-simulated ones with the R values of 0.71 (NICAM-g6str) and 0.68  
613 (NICAM-g6), whereas, in Japan (no available measurements) the  
614 NICAM-g6str-simulated EC concentrations are larger than NICAM-g6-simulated ones

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削除: at the Japanese urban areas such as Nagoya  
(136.97°E, 35.17°N) and Osaka (135.54°E, 34.68°N),

620 | at the Japanese urban areas such as Nagoya (136.97°E, 35.17°N) and Osaka (135.54°E,  
621 | 34.68°N).

622 The NICAM-g6str-simulated sulfate concentrations are larger and more comparable to  
623 the observations over China compared to NICAM-g6-simulated ones. In Japan, the hot  
624 spots with greater concentrations of more than  $5 \mu\text{g}/\text{m}^3$  are found only in the  
625 NICAM-g6str results. The *Br* values are estimated to be 0.59 (NICAM-g6str) and 0.53  
626 (NICAM-g6), whereas the *R* values are estimated to be 0.78 (NICAM-g6str) and 0.88  
627 (NICAM-g6), respectively. The results indicate that the sulfate concentrations obtained  
628 by both NICAM-g6str and NICAM-g6 tend to be underestimated by approximately  
629 | 40-50% compared with the observed sulfate concentrations. The underestimation over  
630 | East Asia is mainly caused by the underestimation in China and possibly by the  
631 uncertainty of the simulated precipitation around Japan. At Hedo located at Okinawa  
632 islands, for example, the underestimation of both NICAM-g6str and NICAM-g6  
633 simulated sulfate concentrations is caused by a possible underestimation of  
634 transboundary sulfate from the continent, which is attributed to a large uncertainty of  
635 the precipitation fields modulated by typhoon in the early August. However, the  
636 | correlations of sulfate between the simulations (both NICAM-g6str and NICAM-g6)  
637 | and observations are adequately acceptable. The simulated and observed  $\text{SO}_2$   
638 concentrations also correlate, with the *R* value of 0.63 (NICAM-g6str) and 0.48  
639 (NICAM-g6). The *Br* values are calculated to be 0.48 (NICAM-g6str) and 0.67  
640 (NICAM-g6). Figure 15 shows that the  $\text{SO}_2$ , which is a primary product, is localized

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**削除:** It suggests that the use of the prescribed oxidants for sulfate formation is not crucial for predicting monthly averaged sulfate mass concentrations at least if the diurnal and seasonal variations of the prescribed oxidants are considered.

646 near the source areas, whereas sulfate, which is as a secondary product, is distributed  
647 from the source to the outflow areas. Although EC is also a primary product, the  
648 horizontal distributions of NICAM-g6str-simulated EC are larger than those of  
649 NICAM-g6str-simulated SO<sub>2</sub>, possibly because EC is less scavenged through the dry  
650 deposition and oxidation processes compared to SO<sub>2</sub>.

651

### 652 3.2.2 Uncertainty in the simulation

653 Sensitivity tests were conducted to examine potential uncertainties derived from  
654 prescribed datasets related to EC and sulfate for the NICAM-g6str simulations. For the  
655 EC sensitivity tests, the emission quantities were set to half and twice of those used in  
656 the standard run in this study. The results for the FAMIKA sites are shown in Figure  
657 16(a) in which the bars show the simulated EC concentrations for both sensitivity tests.  
658 For the majority of the sites, with the exception of Komae, the results obtained by the  
659 sensitivity experiments of twice strength remain underestimated compared with the  
660 measurements. The large underestimation of the EC mass concentrations at Maebashi  
661 and Kisai was also shown by WRF-CMAQ of Shimadera et al. (2013) as well as the  
662 previous studies of WRF-CMAQ in Morino et al. (2010a,b) and Chatani et al. (2014).  
663 However, Fushimi et al. (2011) and Chatani et al. (2014) suggested that the difference  
664 in the EC concentrations between WRF-CMAQ and the measurements is largely  
665 attributed to an underestimation of the EC emission inventory, especially open biomass  
666 burning from domestic sources. The local EC emission can be estimated by a

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672 combination of the data assimilation and intensive measurements (Schutgens et al.,  
673 2012; Wang et al., 2012; Yumimoto and Takemura, 2013).

674 Sensitivity experiments of the SO<sub>2</sub> emissions and the prescribed OH radical used in  
675 sulfur chemistry were executed under half and twice the amounts used in the standard  
676 experiment. The results for the FAMIKA sites are shown in Figure 16(b) in which the  
677 bars show the simulated sulfate concentrations for both sensitivity tests under the  
678 different experiments. Compared with the SO<sub>2</sub> emissions used in the standard  
679 experiment, the doubled amount of SO<sub>2</sub> emissions can overcome the slight  
680 underestimation of the simulated sulfate compared with the observations. Therefore, the  
681 emission inventories of SO<sub>2</sub> should be improved for the better simulation of the sulfate.

682 In this sensitivity tests for oxidants, the SO<sub>2</sub> oxidation by OH radical strongly depends  
683 on the OH concentrations as well as the cloud cover area, whereas the SO<sub>2</sub> oxidation by  
684 ozone and hydrogen peroxide mainly depends on their concentrations, the cloud cover  
685 area, and the cloud water content. The cloud distributions are modulated by some  
686 feedbacks of the sulfate formation through the aerosol direct and indirect effects. As a  
687 result, the sensitivity of the OH radical concentrations to the simulated sulfate  
688 concentration is smaller than that we expected and that to the SO<sub>2</sub> emissions. We also  
689 determined that the sensitivities of the other oxidants to the simulated sulfate  
690 concentrations were small (not shown). These results [and Figures 14 and 15](#) also  
691 suggest that the use of the prescribed oxidants for sulfate formation is not crucial for  
692 predicting [monthly- and](#) weekly-averaged sulfate mass concentrations at least by taking

693 into account for diurnal and seasonal variations of the prescribed oxidants. At the same  
694 time, they also suggest that because the relationship between the oxidants and the  
695 sulfate concentrations through the feedbacks is non-linear and complex, the use of the  
696 prescribed oxidants for sulfate formation can affect the hourly variations of the sulfate  
697 concentrations, and thus the sensitivity of the oxidants to the simulated sulfate should be  
698 investigated.

699

### 700 3.2.3 PM2.5

701 Figure 17 shows the temporal variation in the surface PM2.5 mass concentration at the  
702 18 Japanese sites including 10 sites in the Kanto area using different Y-axes for the  
703 observed and simulated values. At most of the sites, both NICAM-g6str and NICAM-g6  
704 usually captures the synoptic variation of the observed PM2.5, whereas only  
705 NICAM-g6str reproduces the diurnal variation of the observed PM2.5. Table 3 shows  
706 the PM2.5 concentrations in daily, daytime (from 9 am to 4 pm), and nighttime (from 9  
707 pm to 4 am) averages and ratios of daytime to nighttime. The results show that the  
708 simulated PM2.5 concentrations are underestimated compared with the observations by  
709 more than a factor of two and by up to four at Maebashi. As for the diurnal variation,  
710 the results show that the NICAM-g6str-simulated ratios (0.9-1.3) are larger than  
711 NICAM-g6-simulated ones (0.8-0.9), whereas the NICAM-g6str-simulated ones are  
712 smaller than the observed values (1.0-1.8). At Maebashi, where the ratio is higher than  
713 that at other sites, the issue of the poor model performance of the meteorological fields

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716 can be a major reason of the large underestimation, as mentioned in section 3.1. At all  
717 sites, especially Maebashi and Kisai, the possible underestimation of SOA may be a  
718 critical issue, as shown in the fact that the clear diurnal variation of PM<sub>2.5</sub> during  
719 August 4-9 and the high value of the ratios o daytime to nighttime and suggested by  
720 previous studies (Matsui et al., 2009; Morino et al., 2010c). Morino et al. (2010c)  
721 implied that over the Kanto area SOA from anthropogenic sources, which were  
722 disregarded in this study, are large portion of total carbonaceous aerosols, even though  
723 WRF-CMAQ does not correctly reproduce such carbonaceous aerosols. More  
724 sophisticated SOA module, e.g., volatility basis-set approach proposed by Donahue et al.  
725 (2006) based on the categorization of organic vapors with similar volatility, is required  
726 for to produce SOA with higher accuracy. Originally, the underestimation of PM<sub>2.5</sub> is  
727 common among previous studies that employed regional aerosol-transport models  
728 (Morino et al., 2010b, Chatani et al., 2011), primarily because the concentrations of the  
729 observed PM<sub>2.5</sub> include undefined chemical species with mean fractions ranging from  
730 approximately 30–50% in the total PM<sub>2.5</sub> in the summer of Japan (datasets from the  
731 Tokyo Environment Agency and the Kawasaki Municipal Research Institute for  
732 Environmental Protection). Another possible reason is that the PM<sub>2.5</sub> mass  
733 concentration includes water attached to aerosols, depending on the ambient RH  
734 conditions. Therefore, these undefined chemical compounds in this study may account  
735 for a large portion of the difference between the simulated and the observed values.

736 To evaluate the vertical profiles of the PM<sub>2.5</sub> mass concentrations, we used the LIDAR

737 observation operated by the NIES-Japan network. Figure 18 shows the average results  
738 for the simulated and observed extinction coefficient of the spherical particles at  
739 Tsukuba and Chiba in August. At both sites, the vertical profiles and the magnitudes  
740 below 3 km height of the simulated extinction by both NICAM-g6str and NICAM-g6  
741 are comparable to the observed results, whereas the simulated extinction values tend to  
742 be smaller than the observed extinction values near the surface. These results near the  
743 surface are consistent with those obtained by the surface PM<sub>2.5</sub> comparison shown in  
744 Figure 17. In contrast, the extinction values observed by LIDAR include large  
745 variabilities, primarily because they are retrieved from the surface to the cloud base,  
746 which highly varies hour-by-hour and is basically difficult to detect with the high  
747 accuracy, and secondly because they depend not only on the PM<sub>2.5</sub> mass concentrations  
748 but also on the ambient RH and the water amount attached to aerosols. At both sites, the  
749 differences in the extinction between NICAM-g6str and NICAM-g6 are small below 1  
750 km height, whereas those are relatively large above 1 km height. The differences are  
751 attributed to the differences in the primary particles, mainly carbonaceous aerosols,  
752 between NICAM-g6str and NICAM-g6 (not shown). It means that it is attributed to the  
753 difference in the vertical transport between different spatial resolutions. Therefore,  
754 impacts of the difference in the spatial resolution on the distributions of both aerosols  
755 and their precursors should be addressed in the future work.

756

#### 757 **4 Summary**

758 An aerosol-coupled global nonhydrostatic model, which is based on the aerosol module  
759 of Spectral Radiation-Transport Model for Aerosol Species (SPRINTARS) and the  
760 global cloud-resolving model of Nonhydrostatic Icosahedral Atmospheric Model  
761 (NICAM), with a horizontal resolution of approximately 10 km or less in the target  
762 region, is proposed in the present study. Circulations over both the global and target  
763 domains are solved with a single model, whose mesh size varies with fine meshes  
764 covering the target region, to calculate meso-scale circulations in the study region. The  
765 stretched global model requires lower computational costs to simulate atmospheric  
766 aerosols with fine horizontal resolutions compared with the global uniform  
767 nonhydrostatic model, whereas it may require higher computational costs compared  
768 with the general regional models, because the stretched-grid system calculates inside  
769 and outside the target domain. As opposed to the general regional models, the  
770 stretched-grid system does require neither nesting techniques nor boundary conditions.

771 In this study, we developed the new-type regional model with a horizontal resolution of  
772 approximately 10 km to simulate aerosols over Japan, especially in the megacities of the  
773 Kanto area, including Tokyo. To evaluate the model performances in the stretched-grid  
774 system (hereafter referred to as the “NICAM-g6str”), we also simulated  
775 NICAM-SPRINTARS with the globally uniformed grid simulation in glevel-6  
776 resolution (the horizontal resolution is set to 110 km and we call it “NICAM-g6”). Both  
777 NICAM-g6str and NICAM-g6 well reproduce general circulations obtained by  
778 reanalysis of NCEP-FNL under the nudging technique over Asia including the target

779 | region. Only NICAM-g6str usually reproduces both diurnal and synoptic variations of  
780 | the observed weather (temperature, wind, and precipitation) around Japan. Both  
781 | NICAM-g6str and NICAM-g6 generally reproduce monthly mean distributions of the  
782 | observed sulfate and SO<sub>2</sub> over East Asia, with the high correlations of more than 0.5,  
783 | but the underestimation of the simulated concentrations by 40% (NICAM-g6str) and  
784 | 50% (NICAM-g6). The underestimation is mainly caused by the underestimation in  
785 | China and possibly by the uncertainty of the simulated precipitation around Japan. In  
786 | the Kanto area, the results obtained by NICAM-g6str are much closer to the  
787 | observations compared to those obtained by NICAM-g6. Only NICAM-g6str succeeds  
788 | in simulating the wind patterns and the diurnal transitions around the center of the  
789 | Kanto area, although it is inadequate to simulate the wind patterns and the diurnal  
790 | transitions at some sites located at the edge of the Kanto area and surrounded on three  
791 | sides by mountains, e.g., Maebashi, mainly due to the insufficient horizontal resolution.  
792 | NICAM-g6str also generally reproduces both diurnal and synoptic variations of the  
793 | observed and/or a regional aerosol-transport model (WRF-CMAQ) simulated EC,  
794 | sulfate, and SO<sub>2</sub> concentrations, especially with their high correlation (R>0.5) at  
795 | Komae/Tokyo. The standard and sensitivity experiments suggest that (1) emission  
796 | inventories of EC and SO<sub>2</sub> should be improved for the better simulation and (2) the use  
797 | of the prescribed oxidants for the sulfate formation is not crucial for predicting weekly  
798 | and monthly averaged sulfate mass concentrations at least if the diurnal and seasonal  
799 | variations of the prescribed oxidants are considered. As for PM2.5 simulations, only  
800 | NICAM-g6str captures both synoptic and diurnal cycles of PM2.5, with the exception

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804 of the underestimation of the simulated PM<sub>2.5</sub> by at least twice, probably due to the  
805 underestimation of secondary organic aerosol (SOA) from anthropogenic sources and  
806 the high uncertainties of the measurements.

807 Therefore, this new seamless aerosol-transport model, which covers global to regional  
808 scales, can be applied to regional simulations. It suggests that even the simplified  
809 aerosol module (e.g., prescribed oxidants for sulfur chemistry) is applicable for the  
810 regional simulation if the module is coupled to a dynamic core with high horizontal  
811 resolution. To more accurately simulate areas around Japan and develop the simplified  
812 aerosol module, we need to address the following objectives: (1) to increase the  
813 horizontal resolution (less than 10 km) to properly resolve wind fields, which can  
814 greatly influence the delivery of air pollution from Tokyo to subcities such as  
815 Maebashi; (2) to accurately reproduce the cloud and precipitation fields caused by  
816 thermal lows, for example, by applying the finer horizontal resolution and/or more  
817 sophisticated schemes of cloud microphysics such as the double-moment bulk scheme  
818 proposed by Seiki and Nakajima (2014); (3) to use better emission inventories by  
819 developing a data assimilation such as the Kalman smoother proposed by Schutgens et  
820 al. (2012) with intensive measurements in many sites; (4) to simulate strong peaks of  
821 PM<sub>2.5</sub> in the daytime in the Kanto region by implementing more sophisticated module  
822 of SOA formed from both anthropogenic and biogenic sources, such as the volatility  
823 basis-set approach proposed by Donahue et al. (2006), in this model; and (5) to treat  
824 nitrate aerosol through a thermodynamic equilibrium in the simulation of wintertime

825 and/or future scenarios where the relative contribution of nitrate will be larger than that  
826 of sulfate under the changes in emission of  $\text{NO}_x$  and  $\text{SO}_2$  (e.g., Ohara et al., 2007).  
827 These issues are directly connected to the further development of NICAM-SPRINTARS  
828 in both regional and global simulations. Near the future, we will present scenario  
829 experiments at regional scales of 10 km grids and/or address the issue of regional air  
830 quality and its health impacts in densely populated megacities.

831

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850

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1221 chemical signature, regional haze distribution and comparisons with global aerosols,  
1222 *Atmos. Chem. Phys.*, 12, 779-799, doi: 10.5194/acp-12-779-2012, 2012.

1223 Table 1. Statistical values (averages of the observation and simulations, correlation  
 1224 coefficient  $R$  and root-mean-square-error  $RMSE$ ) for meteorological fields using the  
 1225 simulations (NICAM-g6str and NICAM-g6) and observations at seven sites during the  
 1226 same period, as shown in Figures 4 to 7.

		Yokohama	Chiba	Tsuchiura	Adachi	Maebashi	Machida	Kisai
		Temperature						
Average	Observation	27.9	30.1	28.1	29.7	29.1	29.1	27.9
[°C] and difference	NICAM-g6str	26.9 (-1.1)	28.3 (-1.8)	28.3 (0.2)	27.3 (-2.3)	25.5 (-3.6)	25.9 (-3.2)	25.8 (-2.2)
	NICAM-g6	25.5 (-2.4)	26.2 (-3.9)	25.7 (-2.4)	25.5 (-4.1)	23.9 (-5.2)	25.5 (-3.6)	23.9 (-4.0)
R	NICAM-g6str	0.74	0.85	0.84	0.81	0.79	0.74	0.80
	NICAM-g6	0.76	0.67	0.79	0.78	0.71	0.77	0.75
RMSE [°C]	NICAM-g6str	1.9	2.3	1.9	3.0	4.3	3.9	3.0
	NICAM-g6	2.8	4.4	3.1	4.6	5.8	4.0	4.6
		RH						
Average	Observation	73.5	79.0	73.3	75.4	73.7	75.9	71.4
[%] and difference	NICAM-g6str	83.6 (10.0)	77.5 (-1.5)	76.4 (3.0)	77.9 (2.5)	82.7 (9.0)	82.5 (6.6)	81.6 (10.1)
	NICAM-g6	92.2 (18.6)	92.4 (13.4)	93.4 (20.0)	92.2 (16.8)	95.5 (21.9)	92.2 (16.3)	95.5 (24.1)
R	NICAM-g6str	0.64	0.68	0.69	0.72	0.72	0.72	0.81
	NICAM-g6	0.73	0.59	0.79	0.82	0.71	0.74	0.76
RMSE [%]	NICAM-g6str	12.7	8.9	11.0	10.1	14.6	12.9	13.3
	NICAM-g6	19.5	16.2	22.4	19.8	25.5	20.1	26.3
		Wind speed						
Average	Observation	2.9	2.6	1.6	2.6	1.2	2.7	1.9
[m/s] and	NICAM-g6str	4.2	3.8	3.1	3.4	3.1	3.0	2.7

difference		(1.3)	(1.1)	(1.4)	(0.9)	(1.9)	(0.3)	(0.8)
[m/s] (vs. observation) in bracket	NICAM-g6	3.7 (0.7)	5.0 (2.4)	1.0 (-0.7)	3.7 (1.1)	0.9 (-0.4)	3.7 (1.0)	0.9 (-1.0)
R	NICAM-g6str	0.72	0.41	0.65	0.51	0.19	0.59	0.16
	NICAM-g6	0.64	0.43	0.38	0.47	0.12	0.53	0.04
RMSE [m/s]	NICAM-g6str	1.9	2.0	1.8	1.7	2.3	1.3	1.7
	NICAM-g6	1.4	3.0	1.2	1.7	0.7	1.7	1.4

1227

1228 Table 2. Statistical values (averages of the observation and simulations, correlation  
1229 coefficient  $R$  and root-mean-square-error  $RMSE$ ) for EC, sulfate, and  $SO_2$   
1230 concentrations by the simulations (NICAM-g6str, NICAM-g6, and WRF-CMAQ) and  
1231 the observations at four FAMIKA sites during the period from August 6 to 11. The  
1232 WRF-CMAQ results are given by Shimadera et al. (2013).

		Maebashi	Kisai	Komae	Tsukuba
		EC			
Average [ $\mu\text{g}/\text{m}^3$ ] and difference [%] (vs. observation) in bracket	Observation	2.85	2.75	1.23	2.20
	NICAM-g6str	0.39 (-86)	0.60 (-78)	1.10 (-10)	0.73 (-67)
	NICAM-g6	0.52 (-82)	0.52 (-81)	0.49 (-60)	0.58 (-74)
	WRF-CMAQ	0.87 (-69)	1.17 (-58)	0.92 (-25)	0.77 (-65)
R	NICAM-g6str	-0.02	0.41	0.55	0.59
	NICAM-g6	-0.49	-0.28	-0.05	0.16
	WRF-CMAQ	0.08	0.33	0.37	-0.23
RMSE [ $\mu\text{g}/\text{m}^3$ ]	NICAM-g6str	2.62	2.33	0.72	1.85
	NICAM-g6	2.52	2.45	1.10	2.06
	WRF-CMAQ	2.18	1.83	0.88	1.98
		Sulfate			
Average [ $\mu\text{g}/\text{m}^3$ ] and difference [%] (vs. observation) in bracket	Observation	4.79 (-6)	2.86 (44)	4.18 (-32)	4.85 (-12)
	NICAM-g6str	4.51 (-34)	4.14 (11)	2.84 (-46)	4.25 (-26)
	NICAM-g6	3.17 (-21%)	3.17 (42%)	2.25 (-21%)	3.58 (-22%)
	WRF-CMAQ	3.77	4.08	3.30	3.80

R	NICAM-g6str	0.01	0.50	0.51	0.73
	NICAM-g6	0.05	0.56	0.86	0.75
	WRF-CMAQ	0.41	0.02	0.87	0.78
RMSE [ $\mu\text{g}/\text{m}^3$ ]	NICAM-g6str	3.61	2.81	2.71	2.49
	NICAM-g6	3.01	2.30	2.49	2.77
	WRF-CMAQ	2.30	3.37	1.62	2.56
		SO <sub>2</sub>			
Average [ppbv]	Observation	2.74	2.28	2.35	3.79
and difference [%]	NICAM-g6str	1.25 (-54)	1.90 (-17)	2.34 (-1)	2.34 (-38)
(vs. observation) in bracket	NICAM-g6	2.42 (-12)	2.45 (7)	2.52 (7)	3.21 (-15)
R	NICAM-g6str	0.02	-0.04	0.62	0.21
	NICAM-g6	-0.64	-0.52	0.22	-0.04
RMSE [ppbv]	NICAM-g6str	1.82	0.93	0.97	2.08
	NICAM-g6	1.29	0.94	0.85	1.29

1233

1234 Table 3. PM<sub>2.5</sub> concentrations in daily, daytime (from 9 am to 4 pm), and nighttime  
1235 (from 9 pm to 4 am) averages and mean ratios of daytime to nighttime using the  
1236 simulations (NICAM-g6str and NICAM-g6) and the observation at selected seven sites  
1237 in August.

	Maebashi	Kawasaki	Toride	Hasuda	Sapporo	Nagoya	Fukuoka
	Daily mean PM <sub>2.5</sub> [ $\mu\text{g}/\text{m}^3$ ] and standard deviation [ $\mu\text{g}/\text{m}^3$ ]						
Observation	24.9±12.8	23.2±12.9	17.6±9.7	20.6±11.5	12.7±6.3	17.3±10.1	14.3±7.5
NICAM-g6str	6.4±3.9	10.0±7.3	9.0±6.3	8.4±5.0	4.9±3.5	7.5±5.7	3.4±2.6
NICAM-g6	6.7±3.0	6.7±3.3	6.7±3.4	6.7±3.0	4.7±4.1	5.4±3.0	3.5±2.3
	Daytime (9am-4pm) mean PM <sub>2.5</sub> [ $\mu\text{g}/\text{m}^3$ ] and standard deviation [ $\mu\text{g}/\text{m}^3$ ]						
Observation	28.6±14.1	19.4±12.1	15.8±9.0	21.0±10.0	15.0±5.2	11.3±5.4	9.7±5.7
NICAM-g6str	5.9±3.8	7.1±4.3	6.8±4.4	7.2±4.5	5.3±2.8	3.5±2.3	1.6±0.8
NICAM-g6	5.0±1.7	4.0±2.1	4.0±2.4	4.4±1.9	7.4±4.5	2.4±0.9	1.4±0.5
	Nighttime (9pm-4am) mean PM <sub>2.5</sub> [ $\mu\text{g}/\text{m}^3$ ] and standard deviation [ $\mu\text{g}/\text{m}^3$ ]						
Observation	24.4±11.9	24.5±11.8	16.9±9.6	18.5±10.3	10.7±6.6	19.1±8.2	15.4±6.7
NICAM-g6str	7.5±3.6	14.2±9.2	12.1±7.6	10.8±5.5	4.1±3.9	12.0±4.6	5.1±3.1

NICAM-g6	7.5±2.3	9.1±1.5	8.8±2.1	8.4±3.0	2.6±3.1	7.8±1.3	4.4±2.2
	Ratio of daytime-mean PM2.5 to nighttime-mean PM2.5						
Observation	1.8±0.8	1.7±0.5	1.3±0.4	1.2±0.4	1.0±0.4	1.3±0.4	1.1±0.3
NICAM-g6str	1.1±0.6	1.3±0.7	1.1±0.6	1.1±0.5	0.9±0.3	1.2±0.9	1.0±0.6
NICAM-g6	0.9±0.2	0.8±0.1	0.8±0.1	0.8±0.1	0.8±0.2	0.9±0.2	0.8±0.2

1238

1239 **Figure captions**

1240 Figure 1 Topographical maps of (a) East Asia and (b) Eastern Japan, including the  
1241 observation sites for the model validation. The topography is based on GTOPO30 (the  
1242 horizontal resolution is 30 arc seconds, that is approximately 1 km) courtesy of the U.S.  
1243 Geological Survey.

1244

1245 Figure 2 (a) EC and (b) SO<sub>2</sub> emission inventories in 2007.

1246

1247 Figure 3 Horizontal distributions of temperature and winds in August averages at the  
1248 surface and the model height of approximately 5 km over Asia region using reanalysis  
1249 data from NCEP-FNL, simulation by NICAM-g6str, and simulation by NICAM-g6.

1250

1251 Figure 4 Temporal variations in the NICAM-g6str and NICAM-g6 simulated and  
1252 observed air temperature for a height of 2 m at (a) Yokohama, (b) Chiba, (c) Tsuchiura,  
1253 (d) Adachi, (e) Maebashi, (f) Machida and (g) Kisai in August 2007.

1254

1255 Figure 5 Same as Figure 4 but for relative humidity (RH).

1256

1257 Figure 6 Same as Figure 4 but for wind direction.

1258

1259 Figure 7 Same as Figure 4 but for wind speed.

1260

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1262 Figure 8 Horizontal distributions of precipitation in August averages derived from (a)  
1263 simulation by NICAM-g6str, (b) simulation by NICAM-g6, (c) reanalysis data from  
1264 MSM by JMA and (d) reanalysis data from GSMaP.

1265

1266 Figure 9 Temporal variations in the NICAM-g6str and NICAM-g6 simulated and  
1267 observed precipitation amounts at 21 Japanese sites in August 2007. The comparison  
1268 includes 10 sites in the Kanto area; (a) Maebashi, (b) Konosu, (c) Huchu, (d) Tsukuba,  
1269 (e) Tokyo, (f) Yokohama, (g) Abiko, (h) Saitama, (i) Chiba, and (j) Nerima, 3 sites in  
1270 the northern Japan; (k) Niigata, (l) Sendai, and (m) Sapporo, 5 sites in the western  
1271 Japan; (n) Nagoya, (o) Osaka, (p) Himeji, (q) Fukuoka, and (r) Hyuga, and 3 remote  
1272 islands (s) Hachijo-jima, (t) Oshima, and (u) Naha.

1273

1274 Figure 10 Predictive values of daily precipitation using the NICAM-g6str and  
1275 NICAM-g6 simulations and the AMeDAS measurements during August 2007 at the  
1276 sites defined at Figure 9, in units of percentage.

1277

1278 Figure 11 Temporal variations in the simulated (NICAM-g6str, NICAM-g6, and  
1279 WRF-CMAQ) and observed EC mass concentrations near the surface at (a) Maebashi,  
1280 (b) Kisai, (c) Komae and (d) Tsukuba in August 2007. The WRF-CMAQ results are  
1281 given by Shimadera et al. (2013). The left axis in red represents the simulated values,  
1282 and the right axis in black represents the observed values, in units of  $\mu\text{g}/\text{m}^3$ .

1283

1284 Figure 12 Same as Figure 11 but for sulfate.

1285

1286 Figure 13 Same as Figure 12 but for SO<sub>2</sub> without the WRF-CMAQ results, in units of  
1287 ppbv.

1288

1289 Figure 14 Scatterplot of August mean concentrations for EC, sulfate and SO<sub>2</sub> between  
1290 the simulations by NICAM-g6str and NICAM-g6 and the observations at the sites  
1291 shown in the left panels. The statistics parameters, relative bias (Br) and correlation  
1292 coefficient (R), calculated by the simulated and observed concentrations at all the sites,  
1293 are also shown in each panel.

1294

1295 Figure 15 Horizontal distributions of concentrations for EC, sulfate and SO<sub>2</sub> near the  
1296 surface using NICAM-g6str and NICAM-g6 in August averages. The circles in color  
1297 shows the observation results at the sites.

1298

1299 Figure 16 (a) EC and (b) sulfate mass concentrations at the FAMIKA four sites using  
1300 NICAM-g6str under the sensitivity experiments, WRF-CMAQ results shown by  
1301 Shimadera et al. (2013) and the FAMIKA observations in averages of August 6-11. The  
1302 bar represents the range of the sensitivity.

1303

1304 Figure 17 Temporal variations in the NICAM-g6str and NICAM-g6 simulated and  
1305 observed PM<sub>2.5</sub> near the surface at 18 Japanese sites in August 2007. The left axis in

1306 red represents the simulated values, and the right axis in black represents the observed  
1307 values, in unit of  $\mu\text{gm}^{-3}$ .

1308

1309 Figure 18 Extinction coefficients in August averages for the spherical particles  
1310 simulated by NICAM-g6str and NICAM-g6 and the spherical particles observed by the  
1311 NIES-LIDAR network at (a) Tsukuba and (b) Chiba, in units of  $1/(\text{Mm})$ . The bars  
1312 represent the 25th and 75th percentiles of the LIDAR observations.