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The Community Land Model underestimates land-use CO₂ emissions by neglecting soil disturbance from cultivation

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submitted November 20th, 2013
revised March 17th, 2014
Geoscientific Model Development

18 Abstract

19 The Community Land Model (CLM) can simulate planting and harvesting of crops but
20 does not include effects of cultivation on soil carbon decomposition. The
21 biogeochemistry model DayCent does account for cultivation and provides a baseline for
22 evaluating the CLM. With the goal of representing cultivation effects on soil carbon
23 decomposition, we implemented the DayCent cultivation parameterization in the CLM
24 and compared CLM and DayCent simulations at eight Midwestern United States sites
25 with and without the cultivation parameterization. Cultivation decreases soil carbon by
26 about 1350 gC m⁻² in the CLM and 1660 gC m⁻² in DayCent across the eight sites from
27 first cultivation (early 1900s) to 2010. CLM crop simulations without cultivation have
28 soil carbon gain, not loss, over this period, in contrast to the expected declining trends in
29 agricultural soil carbon. A global cultivation simulation for 1973-2004 reduces ecosystem
30 carbon by 0.4 Pg yr⁻¹ over temperate corn, soybean, and cereal crop areas, which occupy
31 approximately 1/3 of global crop area. Earth System Models may improve their
32 atmospheric CO₂ and soil carbon simulations by accounting for enhanced decomposition
33 from cultivation.

34 **1. Introduction**

35 Earth System Models (ESMs) include carbon (C) cycle calculations to assess the
36 biogeochemical effects of changes in the global environment, including changes in the
37 land cover due to land use (Brovkin et al., 2013). The Community Land Model (CLM)
38 underestimates the global land use and land management (LULM) C flux to the
39 atmosphere in large parts of the 19th and 20th centuries in simulations coupled to the
40 Community Earth System Model (CESM), when compared to the Houghton (2003)
41 bookkeeping estimates (Lawrence et al., 2012).

42 The Houghton (2003) estimates of the LULM flux are based on a bookkeeping model
43 that assesses the cumulative C flux to the atmosphere from LULM relative to no LULM.
44 The bookkeeping model performs detailed accounting of carbon pools and fluxes in areas
45 of LULM based on meticulous usage of data combined with an empirical age-class
46 ecosystem model. In contrast, mechanistic land biogeochemistry models like the CLM
47 compute an instantaneous C flux to the atmosphere from the conversion of unmanaged to
48 managed (and vice versa) land; they do not include cumulative C effects of land cover
49 change in the calculated flux for the years following the change (Lawrence et al., 2012).
50 This difference in definition accounts for part of CLM's underestimation of the LULM
51 flux. Similarly the CLM estimate is low relative to estimates from an intercomparison of
52 ESM simulations that replicate the bookkeeping approach by comparing simulations with
53 and without LULM (Brovkin et al., 2013).

54 There are other inconsistencies in the ESM and bookkeeping communities' definitions
55 and usage of the LULM flux (Pongratz et al., 2013; Houghton, 2013; Gasser and Ciais,
56 2013). As just one example, the Houghton (2003) estimates also include more of the

57 LULM activities that contribute to the land use C flux. For example, Houghton (2003)
58 reports contributions from deforestation, afforestation, agricultural abandonment, wood
59 harvest, fire suppression, woody encroachment, and cultivation. CLM accounts for the
60 first four by prescribing annual changes in land cover type and harvesting. Woody
61 encroachment may be implicit in such data, and CLM's latest fire model accounts for fire
62 suppression (Li et al., 2013). Loss of soil C from cultivation is definitely a missing
63 process in the CLM and, as far as we know, in other land biogeochemical models used in
64 ESM simulations.

65 In simulations with the DayCent biogeochemical model at twenty-one sites around the
66 American Midwest, Hartman et al. (2011) account for the loss of soil C from cultivation.
67 Extrapolating to the Great Plains region and accounting for N₂O and CH₄ in their
68 greenhouse gas calculations, Hartman et al. calculate 1.73×10^{15} g CO₂-C equivalents
69 emitted from 1860 to 2003 (1×10^{15} g = 1 petagram = 1 Pg). This number includes
70 processes that partly mitigate the loss of soil C (irrigation, fertilization, grassland
71 restoration), so cultivation alone results in a larger number.

72 Here we investigate the feasibility of accounting for the direct loss of soil C from
73 cultivation in the CLM, which would reduce the current underestimation in land use
74 emissions shown by Lawrence et al. (2012). CESM's existing land use emission term
75 only accounts for the removal of C from replacing unmanaged vegetation with crops and
76 has no direct effect on the soil C.

77 **2. Methods**

78 We use the models DayCent (Parton et al., 1998; Del Grosso et al., 2006) and
79 CLM4.5bgc, the most recent version of the CLM with biogeochemistry (Oleson et al.,

80 2013; Koven et al., 2013). We perform simulations at eight sites distributed across the
81 Great Plains of the American Midwest that span much of the region's climate variations:
82 Cherry and Hamilton Counties in Nebraska; Dewey County, Oklahoma; Dunn County,
83 North Dakota; Hutchinson County, Texas; Kingsbury and Lyman Counties in South
84 Dakota; and Yuma County in Colorado (see Hartman et al. (2011) Fig. 1 for map). At
85 each site we drive the models with nearest neighbor 0.5° in latitude and longitude CRU-
86 NCEP version 4 atmosphere data (Climate Research Unit-National Centers for
87 Environmental Prediction)
88 (http://dods.extra.ccea.fr/store/p529viov/cruncep/V4_1901_2012), available for years
89 1901-2010. We also drive the models with transient atmospheric CO₂ concentration and
90 transient nitrogen (N) deposition data as done by Bonan and Levis (2010), available for
91 years 1860-2010. We prescribe soil texture per site with percent sand/silt/clay values
92 from Hartman et al. (2011). Grass is the native vegetation at all sites (Hartman et al.,
93 2011) (Table 1).

94 We spin up the models at each site with boundary conditions for the year 1860, 100%
95 grass cover, and repeating atmospheric conditions for 1901-1920. We continue with
96 transient simulations that cycle the 1901-1920 atmospheric conditions from 1861 to 1920
97 and proceed with the remaining time series from 1921 to 2010. We perform three such
98 1861-2010 transient simulations:

- 99 1) GRASS with grass cover as in the spin-up but with transient forcings,
- 100 2) CROP where grasses switch to rainfed corn on a site-specific conversion year,
- 101 3) CLTV same as CROP but with direct effect of cultivation on the decomposition of soil
102 C (Table 2). We expect that the first order effect of cultivation on the soil carbon

103 decomposition will not depend on the crop type present in the simulations (rainfed corn
104 rather than the more common at these sites rainfed winter wheat and spring grains).

105 We also perform global CROP and CLTV simulations from 1973 to 2004 to assess
106 large-scale signals of cultivation-enhanced soil C decomposition. As boundary conditions
107 we use transient meteorology (Qian et al., 2006) and transient N deposition and
108 atmospheric CO₂ values as done by Bonan and Levis (2010). We initialize the
109 simulations from a 1972 CROP simulation as a proxy for starting with native soils in
110 1973. In contrast to the site simulations, here we assume that cultivation begins in 1973
111 on all temperate corn, soybean, and cereal crop areas. This is a first evaluation of the
112 potential biogeochemical effect of enhanced C decomposition from soils disturbed by
113 agricultural practices.

114 **2.1. DayCent**

115 DayCent is well documented and well tested (Parton et al., 2005; Del Grosso et al.,
116 2008) in simulations of agricultural, grassland, and forest systems and of various
117 cultivation practices. Hence we treat the DayCent model as a baseline for comparisons
118 with the CLM at the eight sites.

119 Hartman et al. (2011) show that DayCent's crop yields compare very well against
120 observations at the twenty-one sites chosen for their study. Here we select the eight sites
121 where DayCent performs best against observations. We do not expect this selection
122 approach to bias the CLM simulations.

123 DayCent is designed primarily for local/regional applications, while the CLM is
124 designed mainly for global scale applications. Hence DayCent includes a level of detail in
125 the representation of crop management not included in the CLM (Bonan et al., 2013). For

126 example, the DayCent simulations apply increasing N fertilizer over time beginning in
127 1950. The CLM site simulations do not apply N fertilizer.

128 Here we assess the potential biogeochemical effect of adding to the CLM the DayCent
129 representation of agricultural disturbance to soil C by crop cultivation. Cultivation in
130 DayCent refers to a list of plowing or planting events that disturb the soil according to a
131 decomposition enhancement factor (ϵ) for two litter C pools (metabolic and structural)
132 and three soil C pools (active, slow, and passive) (Table 2). $\epsilon > 1$ indicates a
133 corresponding increase in the C decomposition rate due to cultivation; 1.0 indicates no
134 effect. A site-specific DayCent schedule file prescribes the timing of cultivation events
135 per simulation year (Table 3). A cultivation event is assumed to have a 30-day effect on
136 soil decomposition and this replaces the effect of previous cultivation events when 30-
137 day periods overlap.

138 **2.2. The Community Land Model (CLM)**

139 The CLM is the land component of the CESM (Hurrell et al., 2012), though used here
140 in offline mode, i.e. not coupled to interactive models of the atmosphere, ocean, and sea-
141 ice. The CLM is a state of the art biogeophysics and biogeochemistry model that
142 simulates interactions among land surface, soil, and canopy processes. The CLM is
143 widely tested and documented in global, regional, and point simulations and is among the
144 most advanced models of its kind for coupling to an ESM for climate change research.

145 Lawrence et al. (2011, 2012) describe the CLM4.0 and Oleson et al. (2013) describe
146 the CLM4.5 in great detail, including updates relative to the CLM4.0, such as:

147 1) Revised calculation of canopy conductance, gross primary production, and
148 transpiration, consistent with FLUXNET eddy covariance flux towers (Bonan et al.,
149 2011; Sun et al., 2012).

150 2) Revised hydrology (Swenson et al., 2012), snow fraction (Swenson and Lawrence,
151 2012), and representation of lakes (Subin et al., 2012), wetlands, and rivers.

152 3) Revised soil biogeochemistry that includes DayCent-like litter and soil carbon pools
153 and transfers among pools, vertically-resolved soil carbon dynamics, and N-gaseous
154 emissions (Koven et al., 2013).

155 4) Updates to the crop model. CLM4.5 crops use the interactive N algorithm instead of
156 prescribed N as in CLM4.0 (Levis et al., 2012). CLM4.5 accounts for N retranslocation
157 during the grain-fill stage of crops by releasing N stored in the leaves and stems for grain
158 development. To support the retranslocation process, CLM4.5 varies C-to-N ratios in
159 crop C pools, prescribing lower ratios in early stages of the crop development. CLM4.5
160 also includes a simple crop fertilization scheme (Drewniak et al., 2013) that we use here
161 in the global CLTV and CROP simulations but not in the site simulations.

162 We implement DayCent's enhancement of soil C decomposition due to cultivation in
163 the CLM and prescribe the same site-specific DayCent enhancement factors and schedule
164 files (Tables 2 and 3). The CLM partitions structural litter into cellulose and lignin pools.
165 We apply the DayCent structural litter enhancement factor to both of these pools. The
166 CLM performs biogeophysics and biogeochemistry calculations in 10 soil layers to a
167 depth of 3.8 m. In the comparisons with DayCent simulations, we analyze CLM output in
168 the top five soil layers because they cumulatively reach about 29 cm of depth, closer to
169 the depth of DayCent's soil profile calculations (top 20 cm).

170 In CLM's global simulations we simplify the effect of cultivation to one that repeats
171 every year rather than changing according to a schedule file. We designate model grid
172 cells as economically more or less developed and assign soil C decomposition
173 enhancement factors (ϵ) accordingly (Table 4). This protocol was developed for global
174 DayCent simulations (not shown) and the ϵ values differ from those specified for the site-
175 specific simulations (Table 2).

176 **3. Results and discussion**

177 At all eight sites, GRASS has the smallest 1901-2010 trend in soil C of the three
178 simulations because all eight sites start in equilibrium for the GRASS simulation (e.g.,
179 Dunn County shown in Fig. 1). Small trends in soil C in GRASS are due to competing
180 processes, including CO₂ fertilization due to rising CO₂ concentration and increasing soil
181 decomposition by heterotrophic respiration due to warming. Moreover, increased soil
182 decomposition and increased N deposition over time increase the N available to plants
183 and this can increase plant productivity.

184 At all the sites except for Dewey County, the CLM simulates increasing soil C in the
185 CROP simulations (Fig. 1). This is inconsistent with the expectation that pasture-to-crop
186 conversion should lead to loss of soil C due to biomass removal at harvest, in part
187 because crop biomass is returned to the soil as litter at harvest in the CLM. At all eight
188 sites DayCent simulates a decline of 1000-3000 g C m⁻² over the 20th century. DayCent's
189 rainfed corn generates less plant litter than the native grass, especially before fertilization
190 begins around 1950, in part because crop biomass is removed at harvest. Even in Dewey
191 County where the CLM simulates a slow decline in soil C, this is an order of magnitude
192 less than the loss simulated in the equivalent DayCent CROP simulation.

193 At all eight sites the CLM simulates reduced soil C when accounting for the effect of
194 cultivation on the decomposition of soil C (CLTV) relative to when the CLM does not
195 account for this effect (CROP) (Fig. 1). Compared to the DayCent simulations, which
196 were calibrated for each site individually, the CLM performs best in Cherry, Dewey, and
197 Dunn Counties. Here CLM's soil C declines by about 1200, 1500, and 1700 g C m⁻² from
198 1901 to 2010 due to enhanced soil decomposition from cultivation.

199 In Dewey and Dunn Counties the CLM also captures the eventual reduction in soil C
200 loss simulated by DayCent with the adoption of less intensive cultivation practices by
201 farmers (Fig. 1). DayCent shows these declining soil C losses also for Counties where the
202 CLM does not, e.g., Cherry and Hamilton, because DayCent's fertilization effect
203 enhances plant litter inputs. We do not apply fertilizer in these CLM simulations, so we
204 miss the increase in productivity that compensates for increased soil C decomposition
205 from cultivation.

206 At the four other sites, Hutchinson, Kingsbury, Lyman, and Yuma, the CLM
207 underestimates the cultivation-enhanced decomposition and the resulting soil C decline.
208 We attribute this to higher clay contents at these sites (Table 1), resulting in suppressed
209 soil C decomposition and reduced heterotrophic respiration. The CLM also simulates
210 lower NPP and LAI at these sites because clay inhibits plant access to soil moisture.
211 Reduced productivity at these sites contributes to reduced apparent sensitivity of the soil
212 C to cultivation. In other words, less C produced leads to less decomposed, even under
213 cultivation.

214 Consistent with the site simulation results, CLM's global CLTV simulation loses more
215 than 120 g C m⁻² from 1973 to 2004 in the top 29 cm of soil relative to CROP and 800 g

216 m^{-2} in the central United States (Fig. 2). The global ecosystem C declines by 0.4 Pg C yr^{-1}
217 from 1973 to 2004 in CLTV relative to CROP due to the enhanced soil C decomposition
218 over temperate corn, soybean, and cereal areas.

219 **4. Conclusions**

220 Past work has investigated potential biogeophysical effects from not tilling agricultural
221 soils after harvest. For example, in a version of the CCSM3, Lobell et al. (2006)
222 prescribed increased surface albedo to represent the presence of crop residue after harvest
223 and found cooling as a result. Here we address a potential biogeochemical effect from
224 land cultivation.

225 We perform CLM simulations at eight sites in the American Midwest to examine
226 whether accounting for the direct effect of cultivation on soil C decomposition may
227 reduce an underestimation in land use emissions simulated by the CLM (Lawrence et al.,
228 2012).

229 We implement in the CLM the cultivation-enhanced soil C decomposition algorithm
230 used in DayCent (Hartman et al., 2011). According to this algorithm, soil C
231 decomposition responds to farming activities known to disturb the soil and leads to
232 reduced soil C in both the CLM and DayCent relative to simulations without this effect.
233 This simple change brings the CLM closer to simulating the declining trends in
234 agricultural soil C supported by observations (Schlesinger, 1991).

235 We do not calibrate the CLM against observations or DayCent simulations in this
236 study, so the general agreement between the CLM4.5 and DayCent gives us confidence
237 in the reliability of the CLM4.5 as a biogeochemical and crop model. However, we
238 acknowledge that greater agreement at some of the sites and lesser agreement at others is

239 a function of each model's response to site-specific boundary conditions, e.g. the effect of
240 soil texture as discussed above with regard to clay. More generally we find that CLM
241 productivity (e.g. NPP) tends to be more sensitive to site-specific characteristics than
242 DayCent productivity. As a result CLM's soil decomposition responds to cultivation with
243 more sensitivity to such site-specific characteristics than DayCent's and the same is true
244 of the models' responses to the interannual variability of climate (Fig. 1).

245 Global CLM simulations put the site results in large-scale perspective. Enhanced soil
246 C decomposition in areas of temperate corn, temperate soybean, and temperate cereals
247 leads to a loss of ecosystem C at a rate of 0.4 Pg yr^{-1} . If all crop areas – the ones that the
248 CLM represents as crops and the ones that the CLM currently represents as grasses – lost
249 C at this rate, the ecosystem C lost could exceed 1.2 Pg yr^{-1} .

250 This loss rate declines with time as soils affected by the enhanced decomposition
251 gradually approach a new equilibrium. In our global simulations we activate the process
252 of enhanced soil C decomposition in 1973 using present-day crop distributions rather
253 than using transient crop areas and starting from the emergence of agriculture to the
254 present. Given that humans have significantly disturbed present-day crop areas for years
255 to centuries, we assume that true CO_2 emissions from cultivation have been more evenly
256 distributed through time and that soil C losses have declined with time since the initial
257 disturbance.

258 There are concerns of consistency on multiple levels regarding our community's
259 varying definitions and usage of the LULM C flux (Pongratz et al., 2013; Houghton
260 2013; Gasser and Ciais, 2013). As just one example, current generation land and
261 biogeochemical models used in assessments of the global C budget (Le Quéré et al.,

262 2013) are typically compared against bookkeeping models (Houghton 2003) that account
263 for the loss of soil C from cultivation. We propose that land and biogeochemical models
264 have the potential of improving their simulations of soil C and land use emissions by
265 accounting for the loss of C from cultivation. By extension, in this way ESM simulations
266 of atmospheric CO₂ trajectories also have the potential of improving.

267 **Acknowledgments**

268 The National Center for Atmospheric Research is sponsored by the United States
269 National Science Foundation (NSF). This work was supported by NSF grant AGS-
270 1020767. The authors thank R. Anderson and R. Houghton for helpful comments.

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- 385

386 **Tables**387 **Table 1.** Site Characteristics

County	%sand	%clay	Native grass type
Cherry, NE	65	15	50% C3, 50% C4
Dewey, OK	20	15	25% C3, 75% C4
Dunn, ND	20	15	50% C3, 50% C4
Hamilton, NE	20	15	50% C3, 50% C4
Hutchinson, TX	10	35	25% C3, 75% C4
Kingsbury, SD	10	35	75% C3, 25% C4
Lyman, SD	5	45	75% C3, 25% C4
Yuma, CO	40	20	50% C3, 50% C4

388

389 **Table 2.** Cultivation events and their corresponding decomposition enhancement factors for
 390 litter (ϵL) and soil C (ϵS). Litter and soil are indexed in order from labile to more recalcitrant
 391 C pools for three litter pools (metabolic, cellulose, and lignin) and three soil pools (active,
 392 slow, and passive). DayCent combines cellulose and lignin into a single structural litter
 393 pool. Site-specific schedule files prescribe the timing of events in each year (Table 3) in
 394 CLM and DayCent simulations.

395

Index	Description	$\epsilon L1$	$\epsilon L2$	$\epsilon L3$	$\epsilon S1$	$\epsilon S2$	$\epsilon S3$
A	RodWeed Row Plant	1.000	1.100	1.100	1.000	2.554	2.554
B	Planters and Cultivators	1.000	1.200	1.200	1.000	2.815	2.815
C	Field Cultivators and Planters	1.000	1.241	1.241	1.041	3.041	3.041
D	Field and Row Cultivators	1.000	1.500	1.500	1.000	3.500	3.500
E	Sweeps and Tandem Disks	1.000	1.600	1.600	1.100	3.691	3.691
F	Field Cultivator and Tandem Disk	1.000	1.649	1.649	1.149	3.849	3.849
G	Multiple Tandem	1.000	1.735	1.735	1.235	4.435	4.435
H	DisksPoint Chisel Tandem Disk	1.000	1.800	1.800	1.200	4.800	4.800
I	Offset and Tandem Disks	1.000	2.034	2.034	1.234	5.434	5.434
J	Pint Chisel Offset Disk	1.000	3.396	3.396	1.396	7.396	7.396
K	Moldboard Plow	1.000	3.500	3.500	8.000	8.000	8.000

396

397 **Table 3.** Example cultivation schedule for Dunn County, North Dakota. Farming
 398 activities map by indices (column 3) to 30-day enhancement effects on soil
 399 decomposition (Table 2). Farming activities that occur before the 30 days have completed
 400 take full effect and replace previous activities. Farming activities do not combine.

Year	Day of year	index	explanations
1917	159	G	
1917	189	G	
1917	220	G	
1918	111	G	
1918	118	K	
1919	136	K	
1919	141	C	
1919	197	C	
1920	159	G	The previous 3-year period of farming activities repeats...
[...]			
1937	197	C	...and this event completes this phase
1938	159	G	
1938	189	G	
1938	220	G	2 years outside of any 3-year cycle
1939	111	G	
1939	118	K	
1940	131	G	Activity added to beginning of the previous 3-year cycle...
[...]			
1954	197	C	...and this event completes this phase
1955	159	G	Original 3-year cycle resumes...
[...]			
1966	197	C	...and this event completes this phase
1967	159	E	
1967	189	E	
1967	220	E	
1968	111	E	New 3-year cycle
1968	118	I	
1969	136	J	
1969	141	C	
1969	197	C	
1970	159	E	3-year cycle repeats...
[...]			
2008	197	C	...and this is the last event of a complete 3-year cycle
[...]			
2010	118	I	Partial 3-year cycle and simulation end in 2010

402 **Table 4.** CLM's global annual cultivation events and corresponding decomposition
 403 enhancement factors for litter and soil C (pools as in Table 2) in different countries.

404

Date	Crop	Description	ϵ_{L2}	ϵ_{L3}	ϵ_{S1}	ϵ_{S2}	ϵ_{S3}
<i>More developed country</i>							
15 Apr-14 May	All	Offset & Tandem Disks	6.67	6.67	6.67	6.67	1.00
15 May-13 Jun	All	Drill	3.41	3.41	3.41	3.41	1.00
14 Jun-13 Jul	Corn & Soybean	Row Cultivator	3.41	3.41	3.41	3.41	1.00
<i>Less developed country</i>							
15 Apr-29 Apr	All	Plowing	10.00	10.00	10.00	10.00	1.00
30 Apr-14 May	All	Cultivator	2.69	2.69	2.69	2.69	1.00
15 May-13 Jun	All	Drill	3.41	3.41	3.41	3.41	1.00
14 Jun-13 Jul	Corn & Soybean	Hand Weeding	1.10	1.10	1.10	1.10	1.10

405

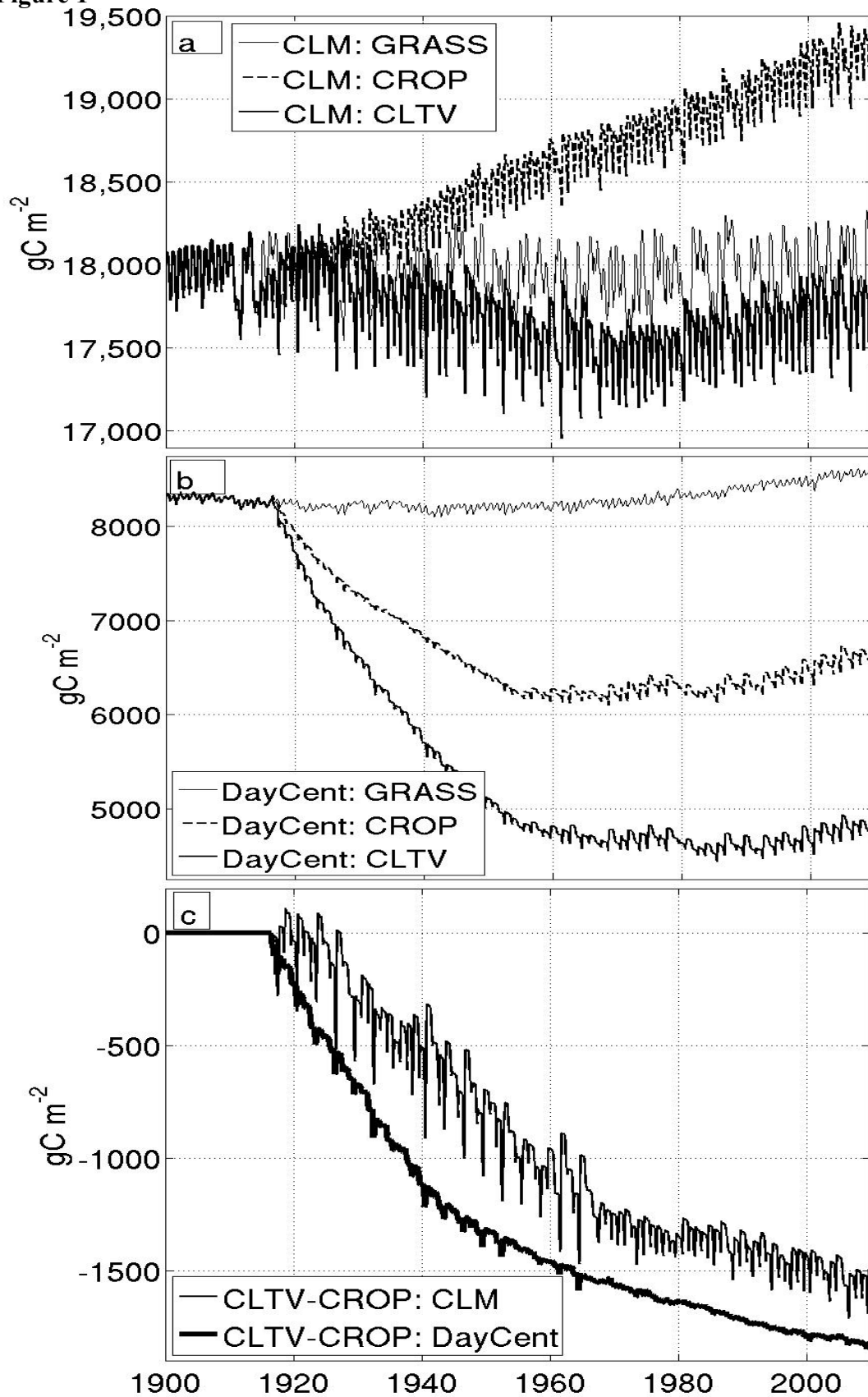
406 Figure Captions

407 **Figure 1.** Soil carbon (g m^{-2}) from the GRASS, CROP, and CLTV simulations in Dunn
408 County, North Dakota. Showing years 1900-2010 (a) from CLM's top ~29 cm of soil depth;
409 values range 1.70×10^4 - 1.95×10^4 gC m^{-2} , (b) from DayCent's ~20-cm soil, and (c) the data
410 from (a) and (b) as the CLTV-CROP difference.

411 **Figure 2.** Area averaged soil carbon (g m^{-2}) in CLM's top ~29 cm of soil depth from the global
412 CROP and CLTV simulations for the years 1973-2004. (a) Central United States at 30-45°N
413 85-105°W and (b) global. Soil carbon increases by about 120 g m^{-2} in both the central U.S. and
414 globally in the CROP simulation. Soil carbon decreases by about 900 g m^{-2} in the central U.S.
415 and by about zero globally in the CLTV simulation. This difference in simulated trends is
416 because the enhanced soil carbon decomposition due to cultivation applies to a much larger
417 fraction of the total land area in the central U.S. than on the global scale.

418

Figure 1

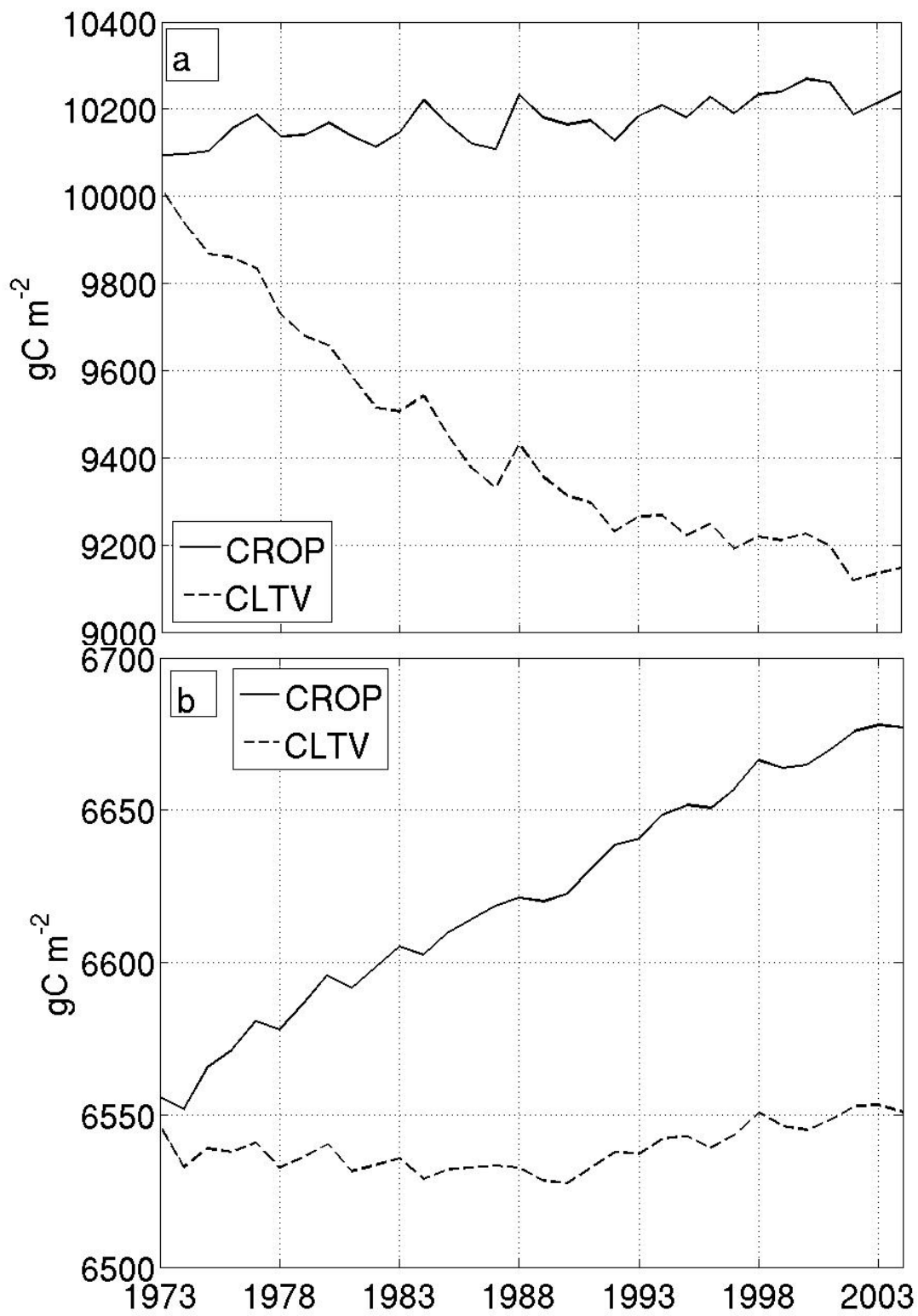


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420

421

422 **Figure 2**



423

424