

Mid-Pliocene global climate simulation with MRI-CGCM2.3

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Mid-Pliocene global climate simulation with MRI-CGCM2.3: set-up and initial results of PlioMIP Experiments 1 and 2

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Abstract

The mid-Pliocene (3.3 to 3.0 million yr ago), a globally warm period before the Quaternary, is recently attracting attention as a new target for paleoclimate modelling and data-model synthesis. This paper reports set-ups and results of experiments proposed in Pliocene Model Intercomparison Project (PlioMIP) using with a global climate model, MRI-CGCM2.3. We conducted pre-industrial and mid-Pliocene runs by using of the coupled atmosphere-ocean general circulation model (AOGCM) and its atmospheric component (AGCM) for the PlioMIP Experiments 2 and 1, respectively. In addition, we conducted two types of integrations in AOGCM simulation, with and without flux adjustments on sea surface. General characteristics of differences in the simulated mid-Pliocene climate relative to the pre-industrial in the three integrations are compared in this study. Generally, difference of simulated surface climate between AGCM and AOGCM is larger than that between the two AOGCM runs, with and without flux adjustments. The simulated climate shows different pattern between AGCM and AOGCM particularly over low latitude oceans, subtropical land regions, and high latitude oceans. The AOGCM simulations do not reproduce wetter environment in the subtropics relative to the present-day, which is suggested by terrestrial proxy data. The differences between the two types of AOGCM runs are little over the land but evident over the ocean particularly in the North Atlantic and polar regions.

1 Introduction

The mid-Pliocene warm period (mPWP), 3.3 to 3.0 million yr ago, is a warm interval with small (or absent) Northern Hemisphere ice sheets relative to any period throughout the Quaternary (e.g. Zachos et al., 2001; Lisiecki and Raymo, 2005; Jansen et al., 2007). It is also the most recent period when global climate was substantially warmer than the present-day for a sustained time with the modern geographical distribution of continent and ocean (e.g. Haywood et al., 2009a). The effort for simulating the climate in this

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interval would advance validations of climate models predicting future climate change (e.g. Crowley, 1996; Salzmann et al., 2009) and to the estimation of “Earth system sensitivity” with the globally warmer climate accompanying drastic changes in land glaciers and vegetation pattern over the land (e.g. Knutti and Hegerl, 2008; Lunt et al., 2010; Pagani et al., 2010).

The paleoclimate conditions in mPWP are investigated by a series of studies that summarized conditions at a large number of marine and terrestrial sites and areas. As a part of the United State Geological Survey (USGS) Global Changes Research effort, the Pliocene Research Interpretation and Synoptic Mapping (PRISM) Project has documented the characteristics of climate in mPWP on a global scale by use of various types of proxy records (e.g. foraminifera, diatoms, ostracods, pollen and plant macrofossil data). The PRISM datasets (e.g. Dowsett et al., 1999) have been used to drive numerical simulations designed to explore the impact of climate forcing and feedback during the Pliocene and assess the reproducibility of climate simulations derived by general circulation models (GCMs) in this period (e.g. Haywood and Valdes, 2004; Jiang et al., 2005; Haywood et al., 2009b; Dowsett et al., 2011). The Paleoclimate Modeling Intercomparison Project (PMIP; e.g. Joussaume et al., 1999; Braconnot et al., 2007a,b), a worldwide framework for studying on the paleoclimate reconstructions and simulations, has focuses on mPWP as one of new target intervals in its latest phase (PMIP3; Otto-Bliesner et al., 2009). By applying the latest version of the PRISM dataset (PRISM3D, Dowsett et al., 2010) for prescribed boundary forcings, Pliocene modeling intercomparison project (PlioMIP) was proposed to access the reproducibility of global climate models for mPWP climate simulations (Haywood et al., 2010, 2011). PRISM3D dataset contains all surface boundary conditions, topography, vegetation, land ice, and also deep ocean temperature for initial condition in air-sea coupled climate simulations. Under the experimental protocols, several results derived by climate models had already been reported (e.g. Chan et al., 2011; Yan et al., 2011; Kamae et al., 2011; Koenig et al., 2012). In this paper, we represented the experimental setups and preliminary results of PlioMIP Experiments 1 (with atmospheric general

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circulation model, AGCM) and 2 (with atmosphere-ocean coupled general circulation model, AOGCM) for pre-industrial control (Control) and mid-Pliocene (Pliocene) climate simulations with MRI-CGCM2.3 (Yukimoto et al., 2001, 2006a), which was also used in the third phase of the Coupled Model Intercomparison Project (CMIP3) and the second phase of the PMIP. In the present study, we focused on comparisons of general patterns in surface climate parameters, temperature and precipitation, for the preliminary report. The paper gives fundamental information on the modeling of Pliocene climate with MRI-CGCM2.3 and its basic results for further studies on their physical mechanisms including atmospheric and/or ocean general circulations (e.g. Hadley circulation, Walker circulation, El-Niño and Southern Oscillation, Indian Ocean Dipole mode, wind-driven surface ocean circulation, and Atlantic Meridional Overturning Circulation; AMOC) and comparisons of the simulated climate with other models and reconstructed climate revealed by proxy data. We also represent Pliocene biome prediction over the land using with equilibrium biogeography model, BIOME4 (Kaplan et al., 2003), under the climate state simulated by the AGCM and AOGCMs. The biome simulations could help to compare general characteristics of surface climate patterns among the simulations by a single index translated from some elemental climate parameters (see Sect. 2.5) and facilitate quantitative model-model (e.g. Harrison et al., 1998; Wohlfahrt et al., 2008) or data-model comparisons (e.g. Haywood et al. 2009b; Pope et al., 2011; Kamae and Ueda, 2011) in forthcoming PliomIP papers.

2 Climate model description

The model used for the present study is air-sea coupled model developed at Meteorological Research Institute in Japan (MRI-CGCM2.3; Yukimoto et al., 2001, 2006a) for climate projections and paleoclimate simulations (e.g. Yukimoto et al., 2006b; Kitoh et al., 2007; Ohba and Ueda, 2010; Ueda et al., 2011). We employ a version with triangular truncation at zonal wave number 42 (T42; an approximately 280 km transform grid) and a 30-layer hybrid sigma-pressure coordinate system with the top at 0.4 hPa.

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5 The model treats cloud with diagnostic scheme applying different types of relationships between cloud amount and relative humidity depends on convective/layer cloud and land/ocean regions (Yukimoto et al., 2006a). For deep moist convection, the Arakawa-Schubert scheme (Arakawa and Schubert, 1974) with prognostic closure similar to that of Randall and Pan (1993) is used in the AGCM. The land component is based on the simple biosphere (SiB) model (Sellers et al., 1986; Sato et al., 1989), which includes the effects of vegetation. The land model has three soil layers with different field capacities depending on the vegetation type, in which the temperature, liquid water, and frozen water in each soil layer are predicted. Canopy and grass are treated for each of the 13 vegetation types, for which the parameters are dependent on the vegetation type and month of the year. Runoff from the soil layers, surface runoff plus underground runoff, is transferred to either a river-mouth grid or and inland-water grid, followed by river routing.

15 The oceanic component is a Bryan-Cox-type global ocean general circulation model (OGCM). The horizontal resolution is 2.5° longitude and 2.0° latitude poleward of 12° S and 12° N, with finer resolution up to 0.5° near the equator for representation of equatorial oceanic waves. The vertical 23 levels are unevenly placed between the surface and the deepest bottom at 5000 m. The sea-ice component is similar to the model developed by Mellor and Kantha (1989). Compactness and thickness are predicted based on the thermodynamics and the horizontal advection and diffusion. The freezing and melting rates of sea ice are calculated with balances of heat and fresh water at the sea-ice bottom, at the open sea surface (freezing only), and within seawater (creation of frazil ice). The advection velocities are determined from the surface ocean current multiplied by an empirical constant (set to one-third at present).

25 In AOGCM simulations, we performed Control and Pliocene runs using two types of model settings, with flux adjustment or without flux adjustment (see Sect. 3.1.3). In Yukimoto et al. (2006a), the climatology of present-day climate simulation was kept close to the observed conditions by using flux adjustments for heat and fresh water obtained from the last part of the 334 yr calibration run, in which the surface temperature

and salinity have been restored to the observed climatology (Levitus et al., 1994; Levitus and Boyer, 1994). Adjustment was also applied for wind stress only in the equatorial region to reproduce a realistic climatological thermocline along the equator, the structure of which plays an important role in the model performance in simulating tropical variabilities, such as El-Niño. The detailed model description, performance in the present-day control simulation, and impact of flux adjustment to control simulation are described in Yukimoto et al. (2006a) and Kitoh et al. (2007).

3 Experimental design

3.1 Control and Pliocene climate simulations

3.1.1 Dataset for boundary and initial conditions

We applied the “PMIP3-style” boundary conditions for AGCM and AOGCMs Control simulations as detailed below. For Pliocene simulation, we added the differences between PRISM3D Pliocene and PRISM3D modern of boundary conditions (sea surface temperature; SST, and topography) converted to the model resolution (atmosphere and land: T42, ocean: 2.5° longitude and 0.5° to 2.0° latitude) to the boundary condition used in the Control runs. All integrations were conducted as “alternate experiment” in which changes in land/sea mask from default setting in climate model are not required (Haywood et al., 2010, 2011). The details of experimental designs are described in Sects. 3.1.2 and 3.1.3, and summarized in Tables 1 and 2.

For surface boundary conditions in Control simulations, we use present-day realistic distribution of topography and land-sea mask (Fig. 1a) which were also used in the previous studies of climate projections for the past, present, and future (e.g. Yukimoto et al., 2006a, b, Kitoh et al., 2007). In order to allow different horizontal resolutions between the atmospheric and the ocean model, the fractional coverage of land-sea is considered in each AGCM grid cell (Fig. 1a). Because of the relatively coarse resolution

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of AGCM (~ 280 km), the orography in the model is smoothed from that in the real (e.g. the highest altitude in the Himalayas is 5536 m). The land cover (vegetation distribution) in Control run is derived from PRISM3D data for modern biomes. For Control simulation with AGCM, SST was specified to modern climatology, LEVITUS94 (Levitus and Boyer, 1994). In AOGCMs, bathymetry is basically derived from ETOPO5 (NOAA, 1998) but the Denmark Strait is made slightly deeper and broader than it to represent the subgrid-scale overflow of waters that form in the Nordic Seas. This modification contributes to the realistic present-day thermohaline circulation in the North Atlantic (Yukimoto et al., 2006a).

The boundary conditions used in this study to reproduce Pliocene climate are derived from PRISM3D dataset (Dowsett et al., 2010). The dataset contains: SST (Fig. 8a) and deep ocean temperature (DOT; Dowsett et al., 2009), vegetation (Fig. 2, Salzmann et al., 2008), ice sheet (Figs. 1b and 2, Hill et al., 2007), and topography (Fig. 1b, Sohl et al., 2009). Reconstructed SST based on 86 sites of proxy data in mPWP (Fig. 8a) shows vastly warmed condition in the mid and high latitude regions (Dowsett et al., 2009), especially in the Northern North Atlantic ($\sim 18^\circ\text{C}$, Robinson, 2009). Over the off the western coast of the continent, the sea surface condition is warmed up significantly (Fig. 8a). The vegetation pattern in mPWP reconstructed by 202 sites of paleobotanical evidences and BIOME simulation (Salzmann et al., 2008) shows poleward shift of vegetation types in mid and high latitude, expansion of wetter biomes in the subtropics, and expansion of extratropical grassland on the Eurasian Continent in mid latitude (Fig. 2b). We also used PRISM3D modern vegetation data for Control simulation in AGCM and AOGCMs (Fig. 2a). The PRISM3D topography (Fig. 1b) reconstructed by palaeobotanical and palaeoelevation evidences represents that the East African Rift Valley were 500 m higher in mPWP relative to the present value (Thompson and Fleming, 1996; Dowsett et al., 1999; Sohl et al., 2009). In contrast, the western cordilleras of the North America and the Northern South America were lower than the present day. Stable isotopic, stratigraphic sea level records, and pollen data on the land reveal significant reduction of the continental ice sheet on Greenland and Antarctica (Thompson

and Fleming, 1996; Dowsett et al., 1999). The land ice extent over Greenland was reduced by the half (Dowsett et al., 1999), with the ice restricted to the high-altitude regions of East Greenland. In East Antarctica, significant ice sheet reduction was recognizable in the Wilkes and Aurora Subglacial basins (Hill et al., 2007; Haywood et al., 2010). The dataset of SST and sea ice is only applied for AGCM simulation. DOT is used for initial condition for AOGCM simulations. The concentration of atmospheric CO₂ is estimated as relatively higher level (350 to 415 ppmv) from that in the pre-industrial (e.g. Kürschner et al., 1996; Raymo et al., 1996; Pagani et al., 2010; Seki et al., 2010; Van De Wal et al., 2011).

Initial condition of ocean for AOGCM Control integration were taken from 31 December of present-day control simulation conducted by Yukimoto et al. (2006a) which was prescribed with present-day boundary forcings (atmospheric CO₂ concentration at 348 ppmv, N₂O at 306 ppbv, CH₄ at 1650 ppbv, solar constant at 1367 W m⁻², eccentricity at 0.016715, obliquity at 23.441°, and precession at 102.70°). To setup the initial conditions of ocean temperature for Pliocene simulations, we used DOT data product in PRISM3D (Dowsett et al., 2009). It was based on estimates in 27 locations using Mg/Ca paleothermometry shows generally warmer condition relative to today.

3.1.2 AGCM simulation

The experimental design for the AGCM simulation follows that presented in Haywood et al. (2010) as “PlioMIP Experiment 1”. The integrations are conducted as “alternate experiment” of PlioMIP in which changes in land/sea mask from default setting in climate model are not required. Figure 1a shows map of land fraction used in the all experiments (Control and Pliocene simulations with AGCM and AOGCMs) conducted in this paper. As with the present-day condition, any modifications were applied to the land-sea mask in the Pliocene simulations (e.g. Central American seaway, Bering Strait, Hudson Bay, and West Antarctica). In Control simulation, SST specified into the model is LEVITUS94 (Levitus and Boyer, 1994). For Pliocene simulation, anomalies of SST (Dowsett et al., 2009), topography, and ice sheet height (Sohl et al., 2009)

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derived from PRISM3D dataset are added to those used in Control. Sea ice cover is incorporated within the SST and is represented at grids where SST is -1.8°C . Types of land cover (including vegetation and land ice) are prescribed, which are PRISM3D (Salzmann et al., 2008) biome data translated into 13 types of SiB classification (Sellers et al., 1986; Sato et al., 1989). The translation scheme is detailed in Table 3. Figure 2 shows SiB distribution prescribed in the model for Control and Pliocene runs. The SiB distribution for the Control run well reproduced that for the present-day represented in Sato et al. (1989). In this study, none of the grids is classified into three types of vegetation (deciduous broadleaf, broadleaf deciduous shrubs, cultivated land). The experimental settings about atmospheric composition and external forcings in this study are summarized in Table 1. The concentration of CO_2 in the atmosphere was set to 280 and 405 ppmv in Control and Pliocene. The latter is chosen to account for possible additional contributions to greenhouse warmth from non- CO_2 greenhouse gases (Haywood et al., 2010, 2011). The seasonally varying ozone profile is prescribed to present-day climatological data taken from Wang et al. (1995). In Pliocene run, all other trace gases, orbital parameters, and solar constant were specified as the same values as the Control settings.

Figure 3a shows time series of global-mean surface air temperature (SAT) in Control and Pliocene runs. Both simulations equilibrate in a few years and show low inter-annual variabilities. We integrated the AGCM for 60 yr and calculated climatological means from the last 50 yr. Choice of the averaging period 50 yr or 30 yr does not show significantly differences in the results.

3.1.3 AOGCM flux- and non-flux-adjusted runs

For the PlioMIP Experiment 2, the simulations with AOGCMs, the experimental design is based on Haywood et al. (2011). The topography, land ice, vegetation, land-sea mask, atmospheric trace-gas concentrations, and the other forcings for Control and Pliocene runs are identical to the AGCM simulation (Table 1). For initial condition, difference in ocean temperature at December between mPWP and modern derived

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from Dowsett et al. (2009) is added to the modern ocean temperature used in the Control simulation (Table 2). Any modifications are applied for initial condition of the sea salinity.

We performed the PlioMIP experiment with two types of AOGCMs, with or without flux adjustment, after the second phase of the PMIP (Braconnot et al., 2007a,b). For analysis of the climate reaching near-equilibrium, we integrate the model for different terms between flux-adjusted run (AOGCM_FA) and non-flux-adjusted run (AOGCM_NFA), summarized in Table 2. In AOGCM_FA, Control and Pliocene runs with the modern and anomalous boundary conditions are integrated for 500 yr and the last 50 yr are used for calculation of the climatological means. In AOGCM_NFA, similar boundary and initial condition are prescribed and integrate for 500 yr without the flux adjustment at first. Then continue the integration for another 500 yr and the last 50 yr are referred to as “AOGCM_NFA Control run”. On the other hand, the anomalous ocean temperature is superimposed on the 500th yr condition in AOGCM_NFA run, then integrate for another 500 yr prescribed with the Pliocene boundary forcings. The last 50 yr is referred as “AOGCM_NFA Pliocene run”.

Figure 3b and c shows time series of global mean SAT and DOT in the four integrations (Control and Pliocene runs with AOGCM_FA and AOGCM_NFA). Both of the surface conditions in Pliocene runs are warmer than those in Control runs because of the initial warming of ocean temperature imposed on the model and the boundary forcings in the Pliocene runs. The individual runs reach their near-equilibrium with the integrations which are suggested by modest trends of SST and DOT during their last 100 yr (Table 5) and small radiative imbalance at top of the atmosphere (TOA) during their last 50 yr (Table 4, detailed in Sect. 4.1). Control run in AOGCM_NFA shows relatively large drift (-0.110 yr^{-1}) than the other integrations.

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3.2 Simulation of biome distribution in Pliocene

3.2.1 BIOME4 model

The equilibrium biogeography model, BIOME4 (Kaplan et al., 2003), is used to translate the climate data from the AGCM and AOGCM experiments into vegetation distributions. The model predicts the most-prevalent vegetation types (biomes) as a function of the seasonal cycle of climate parameters by physiological considerations that place constraints on the growth and regeneration of different plant functional types. These constraints are calculated through the use of limiting factors for plant growth, which include mean SAT of the coldest and warmest month, annual minimum SAT, growing degree days (GDD) above 5 and 0 °C, and calculation of a coefficient (Priestley-Taylor coefficient) for the extent to which soil moisture supply satisfies atmospheric moisture demand. GDDs are calculated by linear interpolated daily SAT from monthly climate data.

3.2.2 Experimental design for Pliocene biome simulation

In this study, we conduct one modern and three Pliocene biome simulations by using the results of climate model simulations described above. We adopted an anomaly procedure (e.g. Kaplan et al., 2003; Wohlfahrt et al., 2008; Haywood et al., 2009b; Kamae and Ueda, 2011) for three (AGCM, AOGCM.FA, and AOGCM.NFA) Pliocene biome simulations. The use of an anomaly procedure, in which the changes in climate between Pliocene and Control simulations are superimposed on an observation-based modern climatology, enables us to compensate the first-order biases in the climate control simulation (e.g. Harrison et al., 1998; Wohlfahrt et al., 2008) and extract general discrepancies in biome types caused by the uncertainties in the boundary conditions during mPWP. To apply this procedure, differences in the climatological values of monthly mean SAT, precipitation, and percent of potential sunshine hours between Pliocene and Control climate experiments were linearly interpolated to the 0.5° grid of

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the BIOME4 model and then added to a long-term mean climatology for the 20th century, CLIMATE 2.2 (Kaplan et al., 2003). This modern climatology is an improved version of the Leemans and Cramer (1991) data set, based on a greatly expanded weather station network. Soil properties were specified from a dataset derived from FAO global soil map (FAO, 1995). Atmospheric CO₂ concentrations of 324 and 405 ppmv are used to force BIOME4 for Control and Pliocene simulations, respectively. The former is approximately the mean value during the period of measurement of the climate station data used in CLIMATE 2.2. The CO₂ concentration in the latter experiment is based on the several types of proxy data during the mPWP (Dowsett et al., 2010) and was also used in Pliocene climate simulations.

4 Results

General characteristics of surface climate conditions during Pliocene simulated by the climate models and potential natural vegetation simulated by BIOME4 model are represented in this paper. We mainly focused on geographical patterns of them and difference between Pliocene and Control, and qualitatively comparisons among the models (AGCM, AOGCM_FA, and AOGCM_NFA). In this Chapter, simulated global mean values and spatial patterns of SAT and precipitation in annual, December, January, and February (DJF) and June, July and August (JJA) mean field, in addition, associated changes in potential vegetations are represented.

4.1 Global mean values

Values of global mean SAT, precipitation rate, and energy balance at TOA in Pliocene and its differences from Control simulated by the models are listed in Table 4. Except the TOA energy balance, the changes of global mean values between Pliocene and Control simulated in AGCM, AOGCM_FA, and AOGCM_NFA are comparable. The surface warming in three simulations are ranged from +1.8 to 2.1 °C. The Pliocene

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warmings simulated with MRI-CGCM2.3 are moderate relative to those with MIROC4m under the similar experimental settings (+2.8 °C in AGCM and +3.5 °C in AOGCM, Chan et al., 2011). The global mean SAT in Control simulations with AGCM is higher (+1.1 °C) than those with AOGCMs in this study. Increase in precipitation rates are ranged from +0.08 to 0.10 mm day⁻¹. The increase of precipitation are also modest relative to MIROC4m simulations (from +0.14 to 0.21 mm day⁻¹, Chan et al., 2011).

Radiative balances at TOA in Pliocene relative to Control change for +3.8, -0.3, and -0.2 W m⁻² (positive downward) in AGCM, AOGCM_FA, and AOGCM_NFA, respectively. The large increase of TOA radiative balance in AGCM Pliocene simulation is consistent with the warmer surface condition during mPWP, including ocean. The small changes of global mean TOA energy balance in AOGCM simulations reveals that the models are reached their near-equilibrium state. Major factors for the positive changes in TOA radiative balance in AGCM run are (1) increase in atmospheric CO₂ concentration (280 to 405 ppmv), (2) changes in cloud radiative forcing, (3) change in surface albedo corresponding with changes in cryosphere (sea ice cover, land glacier, and snow cover over the land), and (4) vegetation cover. The TOA radiative balances and their dependencies on boundary forcings and SST patterns in AGCM run are detailed in Kamae and Ueda (2012).

4.2 SAT

Latitudinal and spatial distributions of differences in SAT between Pliocene and Control simulated by AGCM and AOGCMs are shown in Figs. 4 and 5. All runs (AGCM, AOGCM_FA, and AOGCM_NFA) predict predominant warming in the higher latitude (~ +16 °C) and relatively small change in low latitude (~ +2 °C, Fig. 4b). SAT increases in all runs especially over the high-latitude oceans (Arctic, Northern North Atlantic, and Antarctic Oceans) and neighboring land regions (Greenland and Antarctic Continent, Fig. 5a, d and g). Over the mid-latitude land regions in the Northern Hemisphere, North America and Eastern Eurasian Continents, surface warmings are weak relative to their surrounding regions. They would be related to decreases of surface albedo in

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those regions. Figure 6 shows changes in surface albedo between Pliocene and Control calculated by upward and downward shortwave radiation at surface. In all runs, decreases of surface albedo are dominated over the high-latitude lands (Greenland and edges of Antarctica) and the oceans (Arctic, Northern North Atlantic, and Antarctic Oceans, Fig. 6a–i). The former regions are match with the reduction of the land glaciers (Figs. 1b and 2a and b) prescribed in the models. The surface warming in Arctic, Northern North Atlantic, and Antarctic Oceans (Fig. 5a, d and g) are corresponding with reduction of sea-ice covers (Fig. 6a–i) and increasing SST (Fig. 8a, c and e). The relatively modest warming over the mid-latitude North America and Eastern Eurasian Continents are corresponding with the changes in the prescribed land covers from forest to grassland (Fig. 2a and b) and associated decreases of surface albedo (Fig. 6b, e and h). The above characteristics are common within the three runs. The relationship between the TOA energy budget and surface boundary conditions are represented in Kamae and Ueda (2012).

Spatial characteristics of the SAT change in Pliocene from Control simulated in AGCM are different from those in AOGCMs. The surface warming in mid and high latitude (low latitude) predicted in AGCM is generally larger (smaller) than those in AOGCMs. In AGCM run, SAT in low latitude increases only a little, and decreased in the case over the land (Figs. 4d and 5a). Particularly, decrease of SAT is predominant over the subtropical North Africa, Arabian Peninsula, and North India in JJA (Fig. 5c). The surface cooling over the subtropical land regions simulated in AGCM is corresponding with precipitation increase in those regions (Fig. 6c, see Sect. 4.3). For AOGCM Pliocene simulations, in contrast, SAT increases in all latitude relative to Control (Figs. 4b and 5). The zonal mean SATs simulated in AOGCM runs increase almost uniformly (+1 to 2°C) in low and mid latitude (Fig. 4b). AOGCMs-predicted SATs over the land are relatively higher than those over the ocean (Figs. 4c and f and 5b and c). It is also noting that the reduction of the east-west temperature gradient in AGCM run (Fig. 5a) associated with the prescribed SST (Fig. 8a) is not evident in AOGCM runs (Fig. 5d and g, detailed in Sect. 5).

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Different characteristics of SAT changes in Pliocene from Control are also evident between two AOGCMs, particularly in high latitude. Despite SAT simulated in AOGCM_FA is similar with that in AOGCM_NFA in low latitude, the former is relatively higher (~1 to 2°C) than the latter in high latitude (Fig. 4b). In contrast, SAT over the Northern North Atlantic (between 45° N and 70° N) in AOGCM_NFA run is higher than that in AOGCM_FA run (Figs. 4f and 5b and c). The difference of SAT between AOGCM_FA and AOGCM_NFA in those regions might be closely related to the sea-ice covers (Fig. 6) and the region where deep water formation is dominated. Further detailed discussion for difference in intensity of AMOC between the two AOGCM runs and possible mechanisms for them are necessary in future works.

4.3 Precipitation

Precipitation rates in all Pliocene runs increase relative to Control runs in global mean values (Table 4), but their spatial patterns are not similar. In all runs, the precipitation changes between Pliocene and Control in low latitude are larger than those in mid and high latitude (Fig. 7a–i). In AGCM run (Fig. 7a), the pattern of precipitation change in low latitude is characterized with decreasing (increasing) on the inside (outside) of the tropical rainfall zone in Control run, which was detailed in Kamae et al. (2011). For example, precipitation decreases over the tropical Indian Ocean, the Western North Pacific, the equatorial Central Pacific, the Central South Pacific, and the equatorial Western Atlantic, and increases over the tropical and subtropical Africa, the Arabian Peninsula, the subtropical South Asia and subtropical Oceania. These patterns reveal that inter-tropical convergence zone (ITCZ) in Pliocene is broadening meridionally and zonally the rainfall accompanying with large-scale convergence in ITCZ is weakening compared to Control. On the outside (inside) of the convergence regions in the tropics, the Pliocene climate is more rainy (arid) compared to Control. The simulated wetter condition in the subtropical regions, especially in Africa and Australia, are consistent with the terrestrial proxy evidences (Salzmann et al., 2008). The systematical changes in the precipitation pattern are corresponding with changes in the atmospheric general

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circulations, i.e. weakening and meridional broadening of ascending motion in Hadley circulation and weakening of Walker cells over the Pacific and Indian Oceans (Kamae et al., 2011).

In AOGCM runs, changes of precipitation in low latitude, i.e. increase in the subtropics and decrease in the tropics, are modest relative to that in AGCM (Fig. 7d and g). Except the South Asia, the precipitation increases in the subtropical regions appeared in AGCM (North Africa, Arabian Peninsula, and Australia) are not evident in AOGCM runs (Fig. 7d and g). The intensification of South Asian rainfall simulated by AOGCMs, particularly during JJA (Fig. 7f and i), indicates the strengthening of the Asian summer monsoon in Pliocene. The increase of precipitation in Northern North Atlantic are also modest in AOGCM runs relative to AGCM run, which would be associated with relatively small SAT increase (Figs. 5 and 8). Over the tropical Atlantic Ocean, any significant changes are not appeared in AGCM run, but increases/decreases of precipitation in Northern/Southern Hemisphere are evident in AOGCM runs (Fig. 7d–g), particularly in AOGCM_NFA. The meridionally-asymmetric difference of precipitation over the tropical Atlantic Ocean is suggestive of meridional shift of ITCZ in Pliocene.

4.4 Biome distribution

General characteristics of the surface climate conditions predicted in AGCM and AOGCMs are compared by the single climate index, biome. Figure 9 shows present-day natural vegetation (hereafter, PD_biome) and biome distributions in Pliocene simulated with BIOME4 driven by the climate parameters predicted in AGCM, AOGCM_FA, and AOGCM_NFA runs (hereafter, A_biome, AOFA_biome, and AONFA_biome). In low latitude, A_biome (Fig. 9b) generally shows larger area of moister biomes than PD_biome (Fig. 9a) in accordance with the precipitation increase simulated in AGCM run (Fig. 7a). Tropical forest in equatorial and Eastern South America, equatorial and South Africa, and Southeastern Asia expand for wider regions in A_biome relative to PD_biome. In subtropical regions (North and South Africa, Central Asia, and Australia), savanna and grassland of A_biome shift poleward, thus reducing the fractional coverage of desert. The fraction of the areas for tropical forest and desert in Pliocene relative

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to the present-day on African continent change for +65 % and –30 %, respectively. In mid and high latitude, A_biome (i.e. warm-temperate forest, temperate forest, boreal forest, and tundra) reveal poleward displacements relative to PD_biome. The increases in the coverages of warmer-temperate biomes in mid and high latitude are mainly corresponding with the warmer surface condition in AGCM run (Fig. 5a). The relative contributions of differences of the potential sunshine hours and atmospheric CO₂ concentration between Pliocene and Control for the alterations in the biome patterns are minor (figures not shown) than those of the SAT and precipitation. Comparisons of the biomes with the terrestrial evidences in low and mid latitude during mPWP are detailed in Kamae and Ueda (2011).

In mid and high latitude, AOFA_biome and AONFA_biome also show poleward shifts of biomes relative to PD_biome. In contrast, the expansions of wetter biomes in the subtropical regions (North and South Africa, Central Asia, and Australia) are not evident in AOFA_biome and AONFA_biome. The distributions of AOFA_biome and AONFA_biome in low latitude (Fig. 9c and d) are similar to PD_biome (Fig. 9a), except the South Asia. As a general character, the low-latitude AOFA_biome and AONFA_biome are similar to that of PD_biome, but in mid and high latitude, they show similar pattern to A_biome. In the South Asia, the tropical forest and savanna expand for wider area in AOFA_biome and AONFA_biome than PD_biome and A_biome (Fig. 9c and d).

The reconstructed vegetations during mPWP (Salzmann et al., 2008) in tropics and subtropics are more consistent with A_biome relative to AOFA_biome and AONFA_biome. The expansion of wetter vegetation over the Africa is well simulated in A_biome (Kamae and Ueda, 2011) but those of AOFA_biome and AONFA_biome are spatially restricted. The above characteristics are corresponding with the smaller increase of precipitation over the subtropical Africa in AOGCM runs relative to AGCM run (Fig. 7a–c). The poleward shift of biomes during mPWP in mid and high latitude is qualitatively well simulated in all runs (AGCM, AOGCM_FA, and AOGCM_NFA) which are ascribed to the simulated surface warming in mid and high latitude in all runs (Fig. 5a–c).

5 Discussions

The differences of the simulated climate conditions between AGCM and AOGCM runs would be closely related to SST patterns, which is prescribed as PRISM3D SST (Dowsett et al., 2009) in AGCM run but predicted in AOGCM runs. Figure 8 shows annual mean SST pattern in PRISM3D and those predicted in AOGCM runs. Comparing to Control, Pliocene SST in mid and high latitude in PRISM3D and AOGCM runs are generally higher, particularly in the Northern Hemisphere (Fig. 8a, c and e). The prominent warming of SST in the Northern North Atlantic prescribed in AGCM run (Fig. 7a) is also appeared in AOGCM runs (Fig. 8c and e), but the warmings are much weaker in AOGCMs ($\sim +4^{\circ}\text{C}$) than PRISM3D ($\sim +15^{\circ}\text{C}$). In low latitude, SST difference between mPWP and present-day in PRISM3D shows (1) generally small changes ($\pm 0 \sim 1^{\circ}\text{C}$ for average, Fig. 4f) and (2) decreasing of east-west gradient because of significant warming ($\sim +4^{\circ}\text{C}$) in the western coast of the continents (eastern edges of ocean basins). Above two characteristics are not appeared in AOGCM runs. The low-latitude-mean changes in SST are $+1 \sim 2^{\circ}\text{C}$ (Fig. 4f) and the east-west SST gradient changes little ($\pm 0^{\circ}\text{C}$). These discrepancies between data and model had also been pointed out in the AOGCM Pliocene simulations in Yan et al. (2011) and Chan et al. (2011). In addition, the simulated SST changes in the Northern North Atlantic show different patterns between AOGCM_FA and AOGCM_NFA runs. In AOGCM_FA run, the increasing of SST are dominated in Western and Northern North Atlantic, especially between 50° to 60° N and in Barents Sea, but the SST warming in AOGCM_NFA run is evident in eastern or southern regions relative to AOGCM_FA (e.g. between 30° to 40° N and around Iceland).

The characteristics of the changes in the meridional and east-west climate patterns represented above would be associated with anomalous atmospheric and/or ocean general circulations, which are insufficiently discussed in this paper. The changes of meridional and east-west atmospheric general circulation in PlioMIP experiment 1 and its dependencies on the SST patterns have already pointed out by Kamae

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et al. (2011). The simulated changes of atmospheric general circulations including Hadley and Walker circulations during Pliocene (Brierley et al., 2009; Kamae et al., 2011) are also suggested by proxy data (e.g. Etourneau et al., 2010). Changes of thermohaline and wind-driven ocean general circulations (e.g. Philander and Fedorov, 2003; Fedorov et al., 2010), including AMOC (e.g. Yan et al., 2011; Robinson et al., 2011), during Pliocene would play fundamental roles in formation and sustainment of the Pliocene climate system. The sea surface condition predicted in the AOGCMs (Fig. 8c and e) are also suggestive to drastic changes in AMOC in the Pliocene simulations in this study. Comparisons of the changes in AMOC simulated in AOGCMs and discuss their mechanisms under PlioMIP frameworks are necessary issues. Changes in surface water cycle including precipitation, evaporation, and runoff, and associated sea water salinity, sea-ice cover in Pliocene from Control, and reproducibilities of AMOC in the Control simulations might be key factors for the issues. In addition, comparison of the DOT in multi-model outputs from PlioMIP framework with the proxy data during mPWP (Dowsett et al., 2009) could help to access reproducibility of oceanic general circulations including AMOC in Pliocene in individual models.

6 Conclusions

The global climate model MRI-CGCM2.3 with three types of setting, AGCM, AOGCM_FA and AOGCM_NFA, are applied to PlioMIP Experiments 1 and 2. The initial results, particularly focusing on the surface climate conditions, are reported in this paper. The calculated climatologies with the small drifts in all runs reveal that the changes in the surface climate in Pliocene from Control are quite different between AGCM run and AOGCM runs, particularly in low latitude, the North Atlantic Ocean, and the subtropical land regions. In AOGCMs, the differences are evident for the surface warming in the Northern North Atlantic and the precipitation patterns in the equatorial regions, which are suggestive of different characteristics in simulated meridional circulations. Further studies for the comparisons of simulated structures of atmosphere and ocean

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general circulations during Pliocene among multi-models and performances of them evaluated by comparisons with proxy data would contribute for investigate the detailed aspects of the warm Earth's climate system.

Appendix A

5

List of abbreviations

A_biome:	biome distributions in Pliocene simulated with BIOME4 driven by the climate parameters predicted in AGCM
AGCM:	Atmospheric General Circulation Model
10 AMOC:	Atlantic Meridional Overturning Circulation
AOFA_biome:	biome distributions in Pliocene simulated with BIOME4 driven by the climate parameters predicted in AOGCM_FA
AOFA_biome:	biome distributions in Pliocene simulated with BIOME4 driven by the climate parameters predicted in AOGCM_NFA
15 AOGCM:	Atmosphere-Ocean coupled General Circulation Model
AOGCM_FA:	Flux-Adjusted AOGCM
AOGCM_NFA:	Non-Flux-Adjusted AOGCM
CMIP3:	the third phase of the Coupled Model Intercomparison Project
DJF:	December, January, and February
20 DOT:	Deep Ocean Temperature
GCM:	General Circulation Model

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GDD:	Growing Degree Days
ITCZ:	Inter-Tropical Convergence Zone
JJA:	June, July and August
mPWP:	mid-Pliocene Warm Period
5 MRI-CGCM2.3:	air-sea Coupled General Circulation Model developed at Meteorological Research Institute version 2.3
OGCM:	Ocean General Circulation Model
PD_biome:	Present-Day natural vegetation
PlioMIP:	Pliocene Modeling Intercomparison Project
10 PMIP:	Paleoclimate Modeling Intercomparison Project
PRISM:	Pliocene Research Interpretation and Synoptic Mapping project
SiB:	Simple Biosphere model
SAT:	Surface Air Temperature
SST:	Sea Surface Temperature
15 TOA:	Top Of the Atmosphere
USGS:	the United State Geological Survey

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Table 1. Summary of boundary conditions and external forcings prescribed in the experiments.

	Control		Pliocene (alternate)	
	AGCM	AOGCMs	AGCM	AOGCMs
Sea surface temperature	LEVITUS94	–	Control + PRISM3D anom	–
Land/sea mask			MRI modern	
Topography and ice sheet height	MRI modern		Control + PRISM3D anom	
Vegetation and ice sheet cover	PRISM3D modern		PRISM3D Pliocene	
Atmospheric CO ₂ concentration (ppmv)	280		405	
Other trace gases	CH ₄ : 760 ppbv, N ₂ O: 270 ppbv, CFC: none, ozone: Wang et al. (1995)			
Solar constant (W m ⁻²)	1365			
Orbital parameters	Eccentricity: 0.016724, Obliquity: 23.446°, Precession: 102.04°			

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Table 2. Integration length, initial condition and period for analysis in each run.

	AGCM		AOGCM_FA		AOGCM_NFA	
	Control	Pliocene	Control	Pliocene	Control	Pliocene
Ocean temperature for initial condition	–	–	Present-day control	Control + PRISM3D anom	Present-day control	Control (500th yr) + PRISM3D anom
Total integration	60 yr	60 yr	500 yr	500 yr	1000 yr	500 yr
Averaging period				Last 50 yr		

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Table 3. Converting scheme of BIOME4 vegetation types to SiB classification used in this study.

BIOME4	SiB
28. Land ice	13. Land ice
27. Barren	11. Desert
26. Cushion-forb, lichen, moss tundra	10. Tundra
25. Prostrate shrub tundra	11. Desert
24. Dwarf-shrub tundra	11. Desert
23. Shrub tundra	10. Tundra
22. Steppe tundra	11. Desert
21. Desert	11. Desert
20. Temperate grassland	7. Grass land
19. Tropical grassland	6. Grass land + deciduous conifer
18. Boreal parkland	11. Desert
17. Open conifer woodland	9. Semi-desert
16. Temperate broadleaved savanna	6. Grass land + deciduous conifer
15. Temperate sclerophyll woodland	6. Grass land + deciduous conifer
14. Temperate xerophytic shrubland	7. Grass land
13. Tropical xerophytic shrubland	7. Grass land
12. Tropical savanna	6. Grass land + deciduous conifer
11. Deciduous taiga/montane forest	5. Deciduous conifer
10. Evergreen taiga/montane forest	5. Deciduous conifer
9. Cold mixed forest	4. Evergreen conifer
8. Cool conifer forest	4. Evergreen conifer
7. Cool mixed forest	3. Deciduous broadleaf + evergreen conifer
6. Warm-temperate mixed forest	6. Grass land + deciduous conifer
5. Temperate conifer forest	3. Deciduous broadleaf + evergreen conifer
4. Temperate deciduous forest	6. Grass land + deciduous conifer
3. Tropical deciduous forest/woodland	6. Grass land + deciduous conifer
2. Tropical semi-deciduous forest	1. Evergreen broadleaf
1. Tropical evergreen forest	1. Evergreen broadleaf

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Table 4. Pliocene global annual mean values and anomaly from Control (parentheses). Energy budgets at top of the atmosphere (TOA) are calculated as sum of shortwave and longwave, downward minus upward.

	TOA energy budget (W m^{-2})	SAT ($^{\circ}\text{C}$)	Precipitation (mm day^{-1})
AGCM	335.1 (+3.8)	15.6 (+2.1)	2.70 (+0.08)
AOGCM_FA	337.3 (−0.3)	14.5 (+1.9)	2.65 (+0.10)
AOGCM_NFA	337.8 (−0.2)	14.5 (+1.8)	2.68 (+0.09)

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Table 5. Trends of global annual mean values ($^{\circ}\text{C } 100\text{ yr}^{-1}$) during the last 100 yr in individual integrations.

	SAT (Control)	SAT (Pliocene)	DOT (Control)	DOT (Pliocene)
AOGCM_FA	−0.012	0.048	−0.014	0.028
AOGCM_NFA	−0.110	0.008	0.020	0.020

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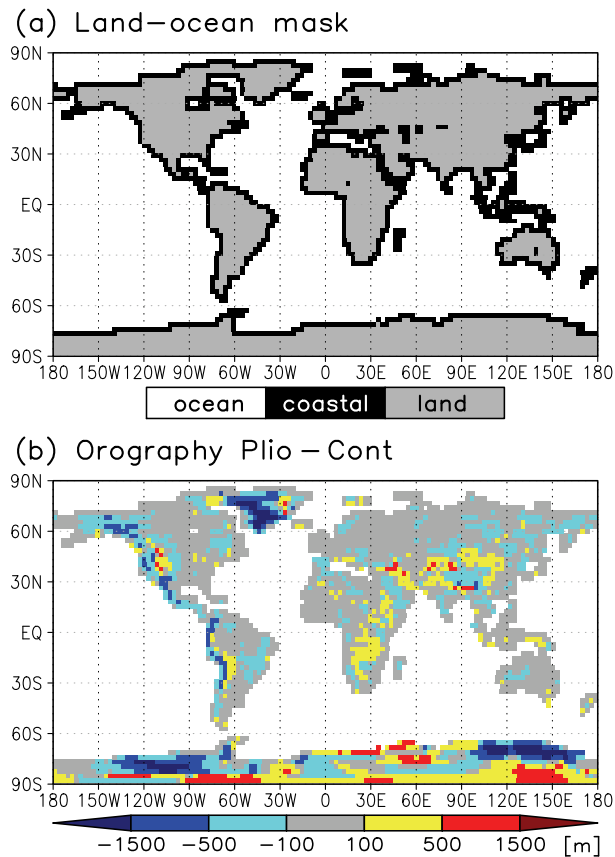


Fig. 1. Boundary conditions used in AGCM and AOGCM. **(a)** Land-ocean mask, **(b)** difference of orography (m) between Pliocene and Control.

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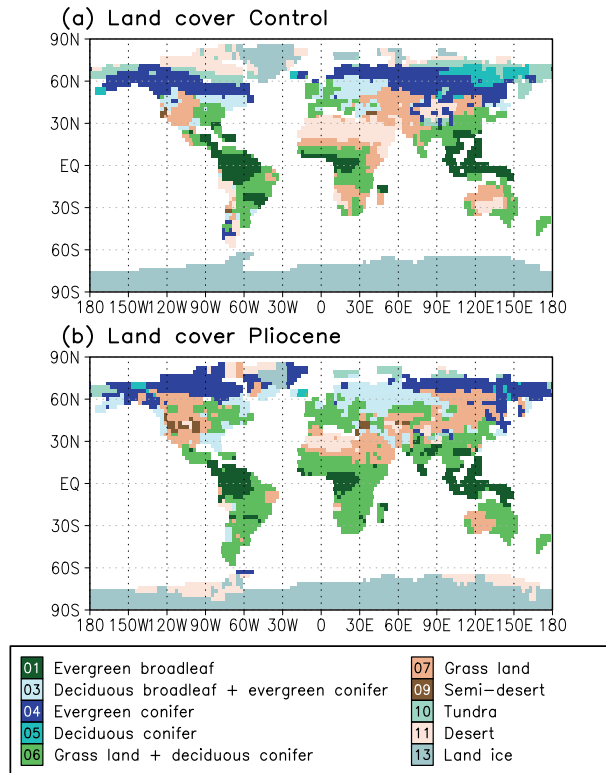


Fig. 2. Land covers prescribed in AGCM and AOGCM. **(a)** Control and **(b)** Pliocene SiB distribution.

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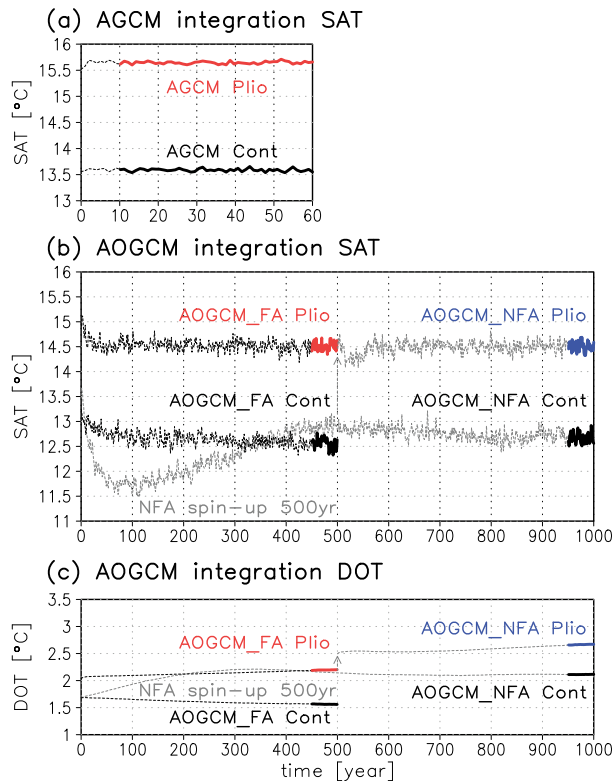


Fig. 3. Time series of the model integrations including spinups. Bold solid lines show the averaging periods of the individual simulations. **(a)** Globally averaged surface air temperature (SAT, °C) for the AGCM and **(b)** AOGCM. **(c)** Same as **(b)** but for deep ocean temperature (DOT, °C) averaged for depth below 1000 m.

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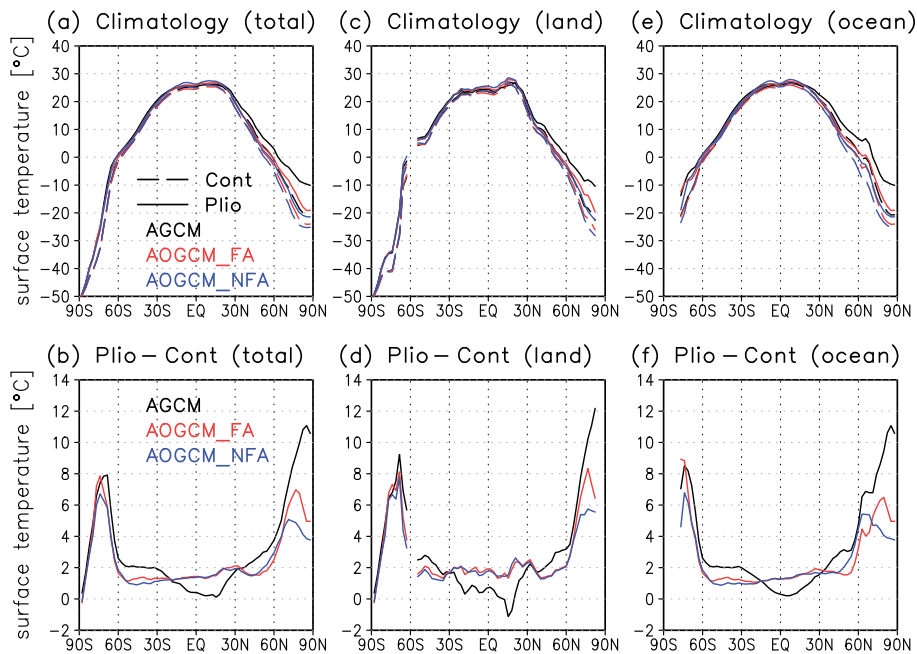


Fig. 4. (a, c, e) Zonal-mean SAT ($^{\circ}\text{C}$) for the Control (dashed lines) and Pliocene (solid lines) and (b, d, f) their differences.

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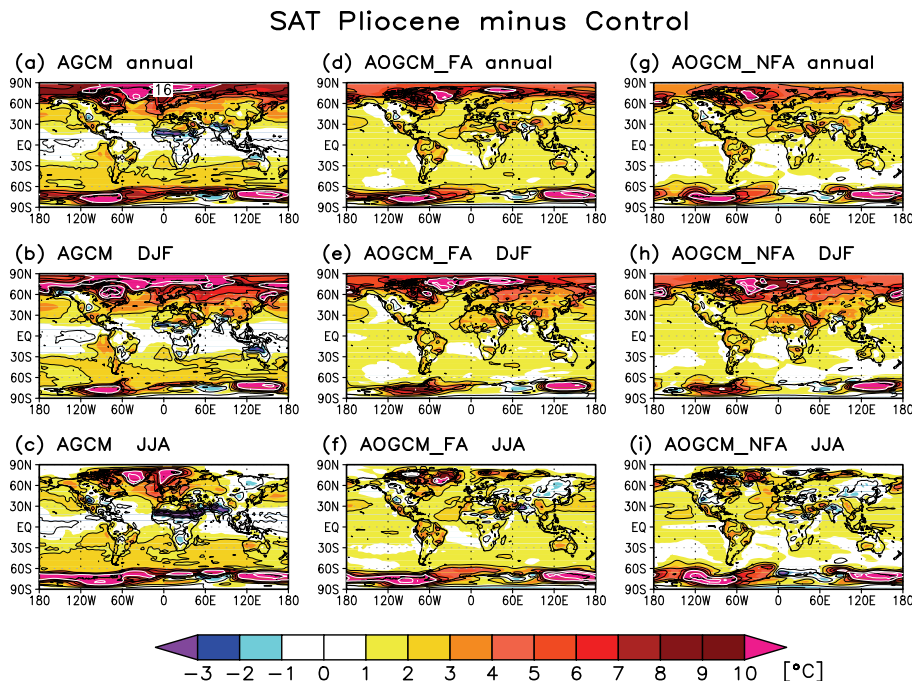


Fig. 5. Difference of SAT (°C) between Pliocene and Control. **(a–c)** AGCM, **(d–f)** AOGCM with flux adjustment (AOGCM_FA), and **(g–i)** AOGCM without flux adjustment (AOGCM_NFA). Black and white contours represent ± 2 , 4, 6, 8 °C and ± 10 , 16 °C, respectively. **(a, d, g)** Annual, **(b, e, h)** December, January, and February (DJF) and **(c, f, i)** June, July and August (JJA) mean field.

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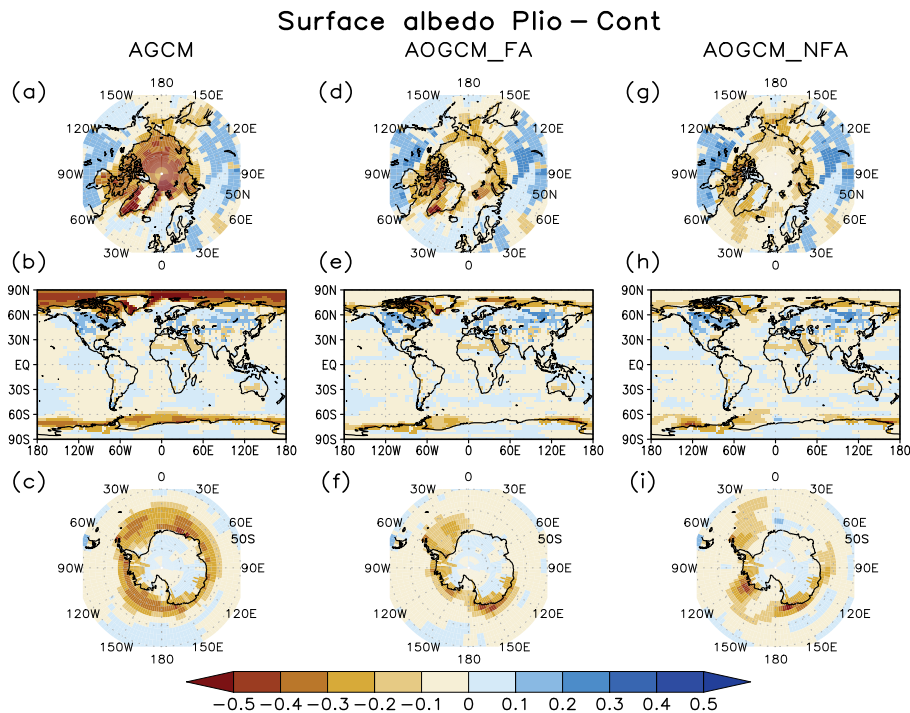


Fig. 6. Difference of surface albedo between Pliocene and Control in (a–c) AGCM, (d–f) AOGCM_FA, (g–i) AOGCM_NFA.

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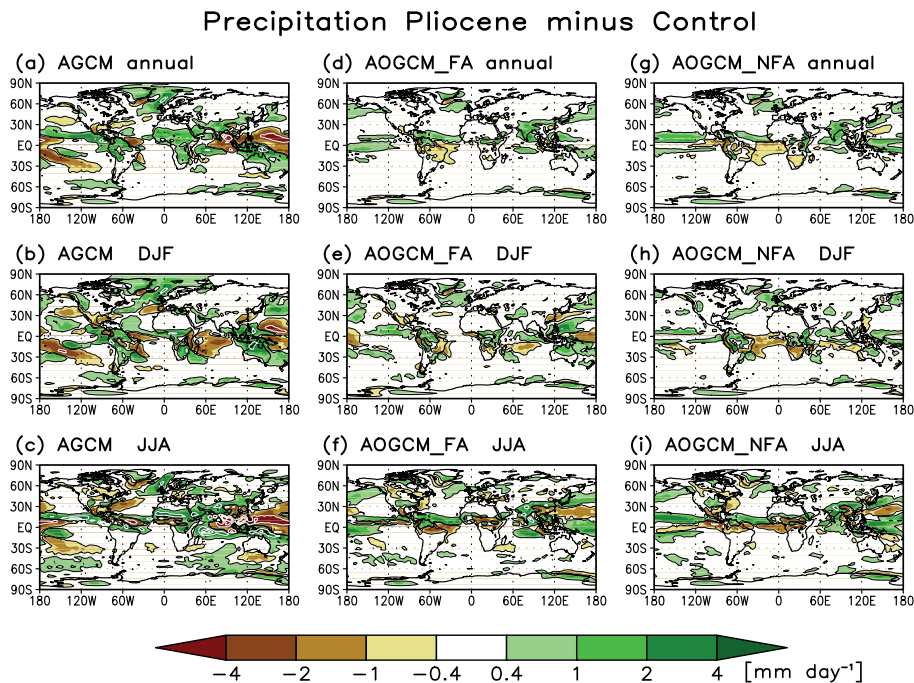



Fig. 7. Same as Fig. 5 but for precipitation rate (mm day^{-1}). Thin black contour represents ± 0.8 , $\pm 2 \text{ mm day}^{-1}$ and white contour interval is 4 mm day^{-1} .

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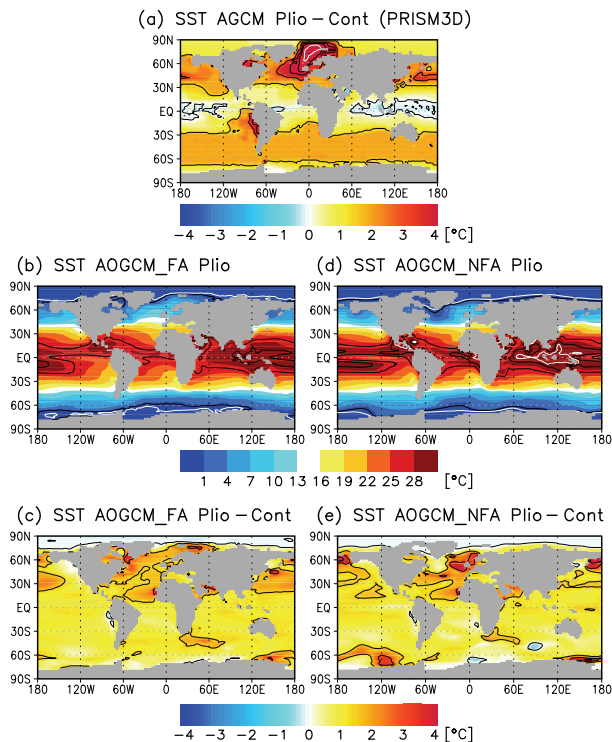


Fig. 8. (a) Difference of sea surface temperature (SST) between Pliocene and Control imposed on AGCM. (b) SST in AOGCM.FA and (d) AOGCM.NFA. (c) Difference of SST between Pliocene and Control in AOGCM.FA and (e) AOGCM.NFA. Black and white contours in (a, c, e) represent $\pm 0, 1.5, 3, 4.5, 6^\circ\text{C}$ and $\pm 10, 15^\circ\text{C}$, respectively. Black and white contours in (b, d) represent $0, 27, 29^\circ\text{C}$ and $-1, 30^\circ\text{C}$, respectively.

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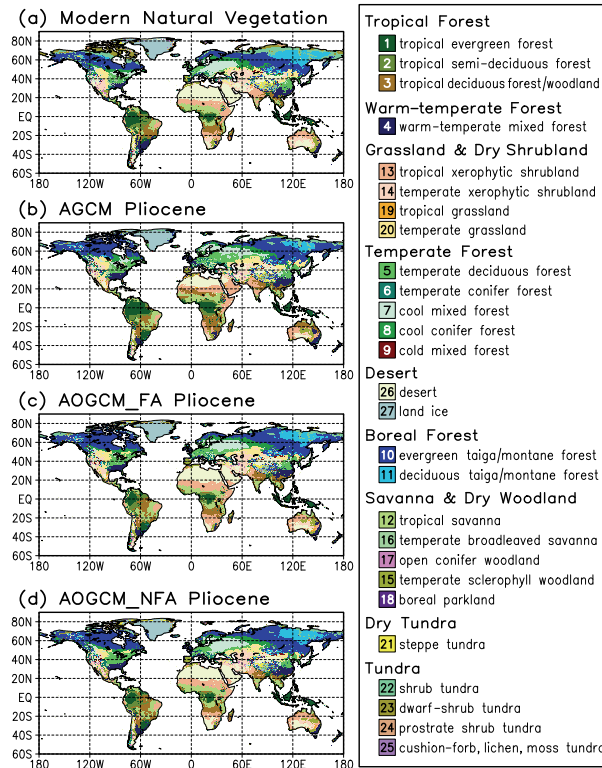


Fig. 9. (a) Present-day biomes simulated by BIOME4 model. (b) Biome distributions in Pliocene simulated with BIOME4 driven by the climate parameters predicted in AGCM (A.biome), (c) AOGCM_FA (AOFA.biome), and (d) AOGCM_NFA (AONFA.biome). Biomes were classified into 9 types of MEGABIOME category (bold) after Harrison and Prentice (2003).

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