

## ***Interactive comment on “Pliocene Model Intercomparison Project (PlioMIP): experimental design and boundary conditions (Experiment 2)” by A. M. Haywood et al.***

**A. M. Haywood et al.**

a.haywood@see.leeds.ac.uk

Received and published: 23 March 2011

When trying to specify CO<sub>2</sub> as far back in time as the Pliocene uncertainty is inevitable.

Evidence for Pliocene CO<sub>2</sub> comes from a number of sources: 1) The stomatal density of fossil leaves (Kürschner et al. 1996) 2) Carbon isotope analyses (e.g. Raymo et al. 1996) 3) Alkenone-based estimates (Pagani et al. 2010; Seki et al. 2010) 4) Boron isotope analyses (e.g. Seki et al. 2010).

The values of CO<sub>2</sub> from each of these proxies differ. However, within error they overlap.

The stomatal density records support a CO<sub>2</sub> of ~350 to 380 ppmv (Fig. 6 page 308)

C85

The average of the Raymo carbon isotope analyses is similar to the stomatal-based estimates but peaks above that value (beyond 450 ppmv) occur and are entirely plausible (Fig. 5 page 320).

The Pagani et al. study looked at reconstructed CO<sub>2</sub> from a number of different marine records and the work shows clearly that in three of the six marine records a CO<sub>2</sub> value of 400 or 405 is perfectly reasonable (sites 806, 1012 and 882; Fig. 2 page 28). The CO<sub>2</sub> range stated in the abstract of this paper is 365 to 415 ppmv.

In the Seki et al. study the alkenone-based CO<sub>2</sub> record (ODP Site 999; Fig. 5b page 206) is consistent with a value around 400 ppmv. The highest resolution section of the Seki et al. boron isotope record (Fig. 8 page 207) shows a decline in CO<sub>2</sub> from ~400 to 280 after 3 million years.

Given that range in CO<sub>2</sub> estimates within the literature, and that fact that a number of the records and techniques can easily support a value of 405 ppmv, we do not accept that we have in any way biased the PlioMIP ensemble "too hot".

The CO<sub>2</sub> value was chosen at the first PlioMIP workshop hosted by the Goddard Institute for Space Studies in New York in 2008. It was the subject of much discussion and debate, but finally a value of 405 was agreed by everyone. This choice has subsequently been validated by the Seki et al. and Pagani et al. studies.

The chosen value is the same as that presented in the experimental design for PlioMIP Experiment 1 already reviewed and published in GMD.

We wish to keep the CO<sub>2</sub> value the same between Experiment 1 and 2 and given (a) that the proxies are perfectly capable of supporting that choice (without even the need to infer changes in CH<sub>4</sub>) and (b) that groups have already completed the simulation, or they are running it, we find no justification in altering the experimental design at this stage.

Pliocene CO<sub>2</sub> is an important ongoing area of research with new records coming on

C86

line in the next few years. Beyond Experiment 2, CO<sub>2</sub> is an obvious choice for sensitivity tests as part of PlioMIP.

References: Kürschner, WM., Van der Burgh, J., Visscher, H., Dilcher, D.L. 1996. Oak leaves as biosensors of late Neogene and early Pleistocene paleoatmospheric CO<sub>2</sub> concentrations. *Marine Micropaleontology* 27, (1/4), 299-312.

Mark Pagani, Zhonghui Liu, Jonathan LaRiviere, Ana Christina Ravelo (2010). High Earth-system climate sensitivity determined from Pliocene carbon dioxide concentrations. *Nature Geoscience* 3, 27-30.

Raymo, M.E., Grant, B., Horowitz, M. & Rau, G.H. (1996). Mid-Pliocene warmth: stronger greenhouse and stronger conveyor. *Marine Micropaleontology* 27 (1/4), 313-326.

Seki, Osamu, Foster, Gavin L., Schmidt, Daniela N., Mackensen, Andreas, Kawamura, Kimitaka and Pancost, Richard D. (2010). Alkenone and boron-based Pliocene pCO<sub>2</sub> records. *Earth and Planetary Science Letters*, 292, (1-2), 201-211.

---

Interactive comment on Geosci. Model Dev. Discuss., 4, 445, 2011.