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LANL* V2.0: global modeling and validation

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We describe in this paper the new version of LANL*. Just like the previous version, this new version V2.0 of LANL* is an artificial neural network (ANN) for calculating the magnetic drift invariant, L*, that is used for modeling radiation belt dynamics and for other space weather applications. We have implemented the following enhancements in the new version: (1) we have removed the limitation to geosynchronous orbit and the model can now be used for any type of orbit. (2) The new version is based on the improved magnetic field model by Tsyganenko and Sitnov (2005) (TS05) instead of the older model by Tsyganenko et al. (2003). We have validated the model and compared our results to L* calculations with the TS05 model based on ephemerides for CRRES, Polar, GPS, a LANL geosynchronous satellite, and a virtual RBSP type orbit. We find that the neural network performs very well for all these orbits with an error typically $\Delta L^* < 0.2$ which corresponds to an error of 3% at geosynchronous orbit. This new LANL-V2.0 artificial neural network is orders of magnitudes faster than traditional numerical field line integration techniques with the TS05 model. It has applications to real-time radiation belt forecasting, analysis of data sets involving decades of satellite of observations, and other problems in space weather.

Introduction

The Earth's radiation belts or Van Allen belts are describing a donut shaped region surrounding Earth that is filled with highly energetic and charged particles which are trapped in the Earth's magnetic field. These particles are energetic enough to penetrate the surfaces of spacecraft and/or instruments and can pose a significant hazard to these assets in space. For example, electron fluxes in the radiation belts can vary by six orders of magnitudes within hours and days. There is a significant need for measuring and modeling this environment and to accurately understand the physical processes causing the time variations in the electron and ion fluxes.

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The large scale motion of charged particles in the Earth's magnetosphere are dominated by the structure of the global geomagnetic and geoelectric fields. At sufficiently high energies (tens or hundreds of keV) the influence of the electric field can be neglected and particle motion can be described by three periodic motions: gyration around the magnetic field, bounce along the magnetic field between magnetic mirror points, and gradient/curvature drift across the magnetic field in an azimuthal direction around the Earth. Each periodic motion has a Hamiltonian invariant and in the Earth's field they are well separated by the adiabatic time scales. The gyro-invariant is the magnetic moment, μ , which is invariant on millisecond time scales. The bounce invariant, given by K, is related to the magnetic field line integrated along the field between mirror points and has time scales of seconds. The drift invariant, Φ , is related to the integral along the azimuthal drift shell around the Earth. If the magnetic field changes slowly relative to a drift period (hours) then the drift path is closed and Φ is adiabatically conserved. A more convenient quantity is L^* (L-star) which is defined as

$$L^* = -\frac{2\pi k_0}{\Phi R_{\rm E}},\tag{1}$$

where k_0 is the Earth's dipole moment, $R_{\rm E}$ is the radius of the Earth (6370 km) and Φ is defined as

$$\Phi = \int_{S} \mathbf{B} \cdot d\mathbf{S}. \tag{2}$$

In a dipole magnetic field, L^* is the distance from the center of the Earth to the equatorial point of a given field line, in units of Earth radii. In a dipole field, particles of any pitch angle would also have the same L^* for a given point in space (see also, Roederer, 1970; Schulz and Lanzerotti, 1974; Schulz, 1991). However, a simple dipole magnetic field is not a sufficiently accurate representation, especially for geosynchronous orbit and beyond. Geosynchronous orbit, for example is at $L^* = 6.6$.

One important challenge for modeling of the radiation belts (and other populations in space) is that the charged particles moving in space form complex current systems

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that in turn distort the geomagnetic field. The interaction of the solar wind, magnetospheric, and ionospheric current systems form an interconnected dynamic system that produces strong distortions of the Earth's field such that it no longer approximates a dipole and, indeed, requires sophisticated numerical field models that are themselves subject of intensive research.

Many models of the Earth's geomagnetic field have been developed but both the pace of development and the numerical sophistication of the models has increased dramatically in the last several decades. Numerically simple models such as the static Olsen-Pfitzer model (Olson, 1974) have given way to dynamic, statistical models driven by a host of solar wind and geomagnetic inputs. The models developed by Tsyganenko and colleagues are representative and are among the most widely used (Tsyganenko et al., 2003; Tsyganenko and Sitnov, 2005). At an even higher level of complexity are global magnetohydrodynamic models or physics based plasma/field model (e.g. Zaharia et al., 2006) but both of these models are sufficiently computer-intensive that they are typically only used for analysis in limited and targeted studies.

The motion of particles in complex, realistic geomagnetic field configurations can be closely approximated using "guiding center" theory representing motion as functions of the three adiabatic invariants, μ , K, and L^* . The first two invariants are relatively easy to calculate even in sophisticated modern field models because they involve only the local field and a one-dimensional integral along a single field line. The third invariant L* is much more difficult and computationally expensive to calculate because it is both two-dimensional and global (McCollough et al., 2008). Typical integration requires on the order of 10⁵ calls to the magnetic field model for obtaining the magnetic field vector. The resulting long computation times often push researchers to compromise and force them to use simpler, less accurate magnetic field models which may produce large inaccuracies and even wrong conclusions (Huang et al., 2008).

Further development of radiation belt and space weather models requires techniques that are computationally feasible and still use the most accurate magnetic field models available. Direct numerical integration of the magnetic field can use standard numerical

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techniques together with the brute force of many processors but other approaches that do not sacrifice accuracy for speed are also possible as we describe below.

In this follow-on paper to Koller et al. (2009), we describes the recently improved and updated version of LANL* V2.0. The model is based on the same artificial neural network (ANN) technique that was used for the first version but now includes two major enhancements: (1) the model can now be used for any type of orbit including a low earth orbit or polar orbit and is not limited to geosynchronous orbit anymore. (2) The new model is now based on the improved magnetic field model by Tsyganenko and Sitnov (2005) (TS05) instead of the older model by Tsyganenko et al. (2003).

In the following sections we describe the ANN setup and the underlying TS05 model that was used for training the neural network. In Sect. 4 we show validation studies and conclude with Sect. 5.

2 The Tsyganenko and Sitnov 2005 Model – TS05

We used the recent magnetic field model TS05 by Tsyganenko and Sitnov (2005). This model is currently the most accurate model (Huang et al., 2008) following a series of models published by Tsyganenko and colleagues. The Tsyganenko magnetic field models are all semi-empirical models based on decades of magnetic field measurements. The models calculate quasi-static states of the Earth's dynamic magnetic field based on solar wind conditions and geomagnetic indices. The quasi-static state is a statistical average for a given set of solar wind conditions but is not a true equilibrium state. The TS05 model is based on space magnetometer data taken during 37 major geomagnetic storms in 1996–2000 and concurrent observations of the solar wind and the interplanetary magnetic field. It accounts for external contributions from the magnetopause, magnetotail current sheet, ring current, and Birkeland current. It also includes partial ring current with field-aligned closure currents which allows it to account for local time asymmetries of the inner magnetospheric field. These currents are driven by separate variables calculated as a time integral for a combination of geoeffective

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parameters of solar wind density, speed, and the magnitude of the southward component of the interplanetary magnetic field. As with the actual geomagnetic field, the TS05 model is compressed on the sunward side by the solar wind and extended on the antisunward side in a comet-like magnetic tail. The model also defines the boundary (a "magnetopause") between the Earth's geomagnetic field and the external solar wind fields. The model includes six parameters W_{1-6} representing the time-integrated driving effect of the solar wind on the magnetosphere (Tsyganenko and Sitnov, 2005). We used the <code>irbem-lib</code> (Boscher et al., 2010) implementation of the magnetic field model TS05 in SpacePy (Morley et al., 2011)

3 Artificial Neural Networks (ANN)

An artificial neural network consists of a number of non-linear processing units that are interconnected through weighted communication lines (see Reed and Marks, 1999, for an introduction). The units, called "neurons", receive input signals from a number of other nodes and produce a single scalar output which then can be used as input to other neurons via new weighted connections.

Neural networks are organized in layers. The first layer provides a node for each input element. In our case the input layer for LANLstar consists of 19 nodes (Table 1), one for each input parameter for the TS05 model plus additional nodes (node 14-15) to help specify the drift shell especially for low earth orbit. The hidden layer in our neural network contains 20 neurons that are connected to each input node and one output node to produce L^{*} .

Typical ephemerides of orbits in the inner magnetosphere are located on closed drift shells. However, during geomagnetic storm conditions, the magnetosphere can be compressed by the solar wind and higher drift orbits can connect to the magnetopause for which the magnetic flux integral (Eq. 2) is not defined. Such a discontinuity in L^* requires a separate ANN. Otherwise, the ANN would simply extrapolate L^* into regions where it is in reality not defined. Therefore, we created a second neural network

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LANLmax to describe the maximum possible value for L^* under the specified solar wind conditions. The maximum L^* is often referred to as the last closed drift shell. Table 2 lists the input parameters necessary for calculating $L_{\rm max}$. Since $L_{\rm max}$ is describing a global configuration, the neural network LANLmax is also independent of specific ephemerides (r, θ, φ) but still a function of equatorial pitch angle $\alpha_{\rm eq}$.

Similar to the real nervous system, artificial neural networks have to be trained by learning from examples. We used the latest version of the <code>irbem-lib</code> library in SpacePy to generate the input-output database (see also Koller et al., 2009). We created one million samples with a uniform random number generator for ephemeris (1.03 < $r_{\rm GEO}$ < 11, -180° < θ < +180°, -90° < φ < +90°) and pitch angle values (10° < $\alpha_{\rm loc}$ < 90°). The solar wind conditions for all samples are based on randomly selected conditions during a full solar cycle from 1996 to 2007. Solar wind conditions and magnetic field parameters were taken from http://virbo.org based on Qin et al. (2007) using the implementation in SpacePy (Morley et al., 2011).

The second neural network LANLmax was trained with the last closed drift shell calculated with irbem-lib as well. We generated 10 000 training examples by using a bisection search algorithm stepping radially outward towards dusk in a cartesian solar magnetic (SM) coordinate system. The accuracy of the bisection search algorithm for L_{max}^* was set to 0.01.

We used a constrained truncated Newton algorithm in the ffnet python module (Wojciechowski, 2009) to train an ANN on the input-target data for both neural networks Lanlstar and Lanlmax. The training algorithms were specified with a tolerance of 10^{-6} .

Neural networks can show a degree of fault tolerance due to the redundant parallel structure. Many nodes draw information from a number of other nodes to produce one overall output. In the case that a certain input value is not available, the system will degrade but not necessarily completely fail because the correlation functions are distributed over several other nodes (Reed and Marks, 1999). Our neural network requires the magnetic field strength $B_{\rm mirr}$ at the mirror point. For a self-consistent calculation,

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one would use the TS05 model to obtain this value. However, one could also use the graceful degradation property of neural networks and calculate this number based on a simpler magnetic field model or even a dipole field. We will perform a detailed study in a future publication. We have used the self-consistent calculation using TS05 for the validation study below.

4 Validation study

We validated both neural networks with completely independent data using a selection of actual satellite ephemerides. These ephemerides and solar wind conditions where not part of the training data set. The first example is shown in Fig. 1 using LANL geosynchronous satellite 1990-095. We selected 1000 random ephemeris locations between 15 October 2001 and 30 June 2005 and calculated L^* based on both the TS05 model and LANL*. They agree quite well with a standard error of $\Delta L^* \approx 0.088$ which corresponds to 1.3%. Pitch angles for this and all other validation were randomly selected between $10^\circ < \alpha_{loc} < 90^\circ$.

Another independent validation is shown in Fig. 2. We selected randomly 1000 ephemeris points between 1 March 1991 and 1 April 1991 from the CRRES satellite orbit. Note that this is outside of the solar cycle data used for the training data. This also includes the extreme storm from 24 March 1991 with a $Dst_{min} \approx -300\, nT$. The neural network performs very well for this geo-transfer orbit with a standard error of $\Delta L^* \approx 0.1$.

We achieved the best result with GPS-ns41 ephemerides (Fig. 3). GSP-ns41 is in a circular orbit with $r \approx 4.2\,R_{\rm E}$. Again, we selected 1000 random ephemeris locations between 1–30 April 2004. The standard error is $\Delta L^* \approx 0.04$, which corresponds to a 1% error.

The 1000 ephemeris locations for the Polar satellite (Fig. 4) were selected randomly between 1996–2005 covering a wide range of ephemerides due the precession of the

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spacecraft. The standard error for this polar type orbit is $\Delta L^* \approx 0.1$ for randomly chosen local pitch angles.

Figure 5 shows the validation study for a "virtual" RBSP satellite (Radiation Belt Storm Probes) which is a future NASA mission, planned to launch in May 2012. The future ephemeris was mapped back in time to start in February 2000. We selected 1000 random point between February 2000 and January 2002 for the validation. The standard error is $\Delta L^* \approx 0.06$.

We also provide a time series figure for a particular storm on 25 October 2002. This geomagnetic storm commenced due to a high speed solar wind stream resulting in a $Dst = -100\,\text{nT}$. Figure 6 shows a comparison of the neural network LANL* versus the TS05 model for three different local pitch angles $\alpha_{\text{loc}} = 90^\circ$, 60° , 30° . Figure 7 shows a time series for the last closed drift shell model LANLmax for the same time period from 23 October 2002 to 4 November 2002 for $\alpha_{\text{loc}} = 90^\circ$.

5 Conclusion and summary

We have presented a new version of our LANL* model to calculate L^* , which is a computationally intensive but important input parameter for radiation belt models. Accurate magnetic field models have been neglected for calculating this value due to the computational burden and often a simple dipole field is used instead. Our model can calculate L^* based on a sophisticated dynamic magnetic field model at a fraction of the time required for full drift shell integration using a neural network technique. Once all input parameters for the neural network are assembled, one million calculation will only take a few seconds to run on a modern desktop computer. This is a speedup of almost six orders of magnitudes while adding only a few percent of error to the output which is negligible considering the uncertainty in the underlying magnetic field model TS05 itself.

This new, computationally efficient model is particularly important for real-time processing of space weather applications and studies involving solar cycles of data sets.

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While this particular version has only been trained with the TS05 model, the technique itself is applicable to other magnetic field models as well. We are currently working on developing a neural network for a variety of empirical magnetic field models and self-consistent physics based numerical models like RAM-SCB (Zaharia et al., 2006).

The LANL* neural network model is available for download at http://www.lanlstar.net.

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Table 1. Input parameters for the neural network LANLstar.

Number	Parameter	Description
1	t _Y	Integer number representing the year
2	t_{DOY}	Day of the year [int]
3	t_{UT}	Universal Time in units of hours [float]
4	Dst	Disturbance storm time index [nT]
5	$p_{\sf sw}$	Solar wind dynamic pressure [nPA]
6	B_{v}	Y component of the IMF field [nT]
7	$B_z^{'}$	Z component of the IMF field [nT]
8–13	W_{1-6}^{-}	See (Tsyganenko and Sitnov, 2005)
14	L_{m}	McIllwain value (Roederer, 1970)
15	B_{mirr}	Magnetic field strength at mirror point [nT]
16	$\alpha_{ m loc}$	Local pitch angle [deg]
17	r_{GSM}	Radial coordinate in GSM system $[R_E]$
18	$ heta_{GSM}$	Latitudinal coordinate in GSM [deg]
19	$arphi_{GSM}$	Longitudinal coordinate in GSM [deg]

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Table 2. Input parameters for the neural network ${\tt LANLmax}$.

Number	Parameter	Description
1	t _Y	Integer number representing the year
2	t_{DOY}	Day of the year [int]
3	t_{UT}	Universal Time in units of hours [float]
4	Dst	Disturbance storm time index [nT]
5	$ ho_{\sf sw}$	Solar wind dynamic pressure [nPA]
6	B_{ν}	Y component of the IMF field [nT]
7	$B_z^{'}$	Z component of the IMF field [nT]
8–13	W_{1-6}	See (Tsyganenko and Sitnov, 2005)
14	$lpha_{ m eq}$	equatorial pitch angle [deg]

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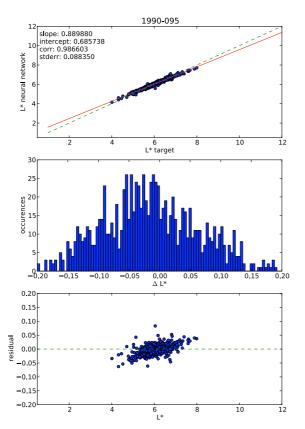


Fig. 1. Validation of neural network with LANL geosynchronous satellite 1990-095 ephemerides. Top: scatter plot of L^* values from model (target) against neural network results. Green-dashed line indicates a perfect fit and red solid line indicates the least-square fit to the scatter plot. One thousand ephemeris points and solar wind conditions were chosen randomly between 15 October 2001 and 30 June 2005. Middle: histogram plot of overall residuals from scatter plot. Bottom: residual plot as a function of L^* .

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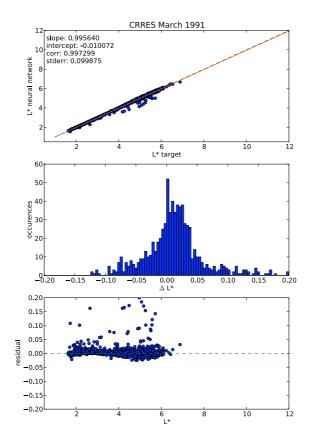
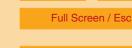


Fig. 2. Validation of neural network with CRRES satellite ephemerides. Top: scatter plot of L* values from TS05 model (target) against neural network results. Green-dashed line indicates a perfect fit and red solid line indicates the least-square fit to the scatter plot. One thousand ephemeris points and solar wind conditions were chosen randomly between 1 March 1991 and 1 April 1991 and include the geomagnetic superstorm on 24 March 1991 with $Dst \approx -300$. Middle: histogram plot of overall residuals from scatter plot. Bottom: residual plot as a function of L^* .

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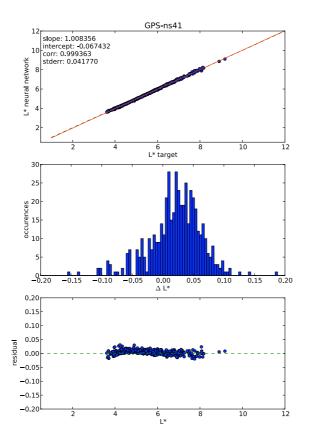


Fig. 3. Validation of neural network with GPS-ns41 satellite ephemerides. Top: scatter plot of L* values from TS05 model (target) against neural network results. Green-dashed line indicates a perfect fit and red solid line indicates the least-square fit to the scatter plot. One thousand ephemeris points and solar wind conditions were chosen randomly between 1-30 April 2004. Middle: histogram plot of overall residuals from scatter plot. Bottom: residual plot as a function of L^* .

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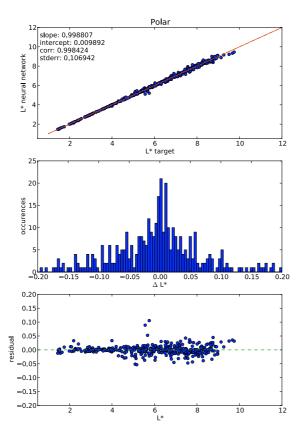


Fig. 4. Validation of neural network with Polar satellite ephemerides. Top: scatter plot of L^* values from TS05 model (target) against neural network results. Green-dashed line indicates a perfect fit and red solid line indicates the least-square fit to the scatter plot. One thousand ephemeris points and solar wind conditions were chosen randomly between 1996–2005 covering a wide range of ephemerides due the precession of the spacecraft. Middle: histogram plot of overall residuals from scatter plot. Bottom: residual plot as a function of L^* .

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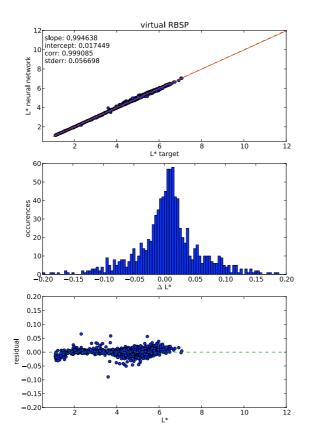


Fig. 5. Validation of neural network with ephemerides from one of the RBSP satellites using an example ephemeris file of this future mission and mapping them to solar wind condition between February 2000 and January 2002. Top: scatter plot of L* values from TS05 model (target) against neural network results. Green-dashed line indicates a perfect fit and red solid line indicates the least-square fit to the scatter plot. One thousand ephemeris points and solar wind conditions were randomly chosen for the above time period. Middle: histogram plot of overall residuals from scatter plot. Bottom: residual plot as a function of L^* .

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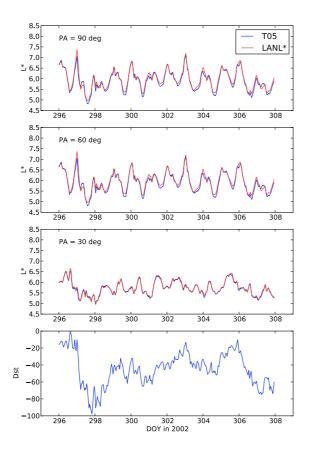


Fig. 6. Calculating the L^* time series during a moderate storm with $Dst = -100 \, \text{nT}$ following the ephemerides of LANL geosynchronous satellite 1990-095. The geomagnetic storm had its Dst minimum close to 25 October 2002 (DOY = 298). The L^* values from the underlying TS05 model (blue) and from the LANL* neural network (red) are compared for 90, 60, 30° pitch angles. The bottom panel shows the disturbance storm time index (*Dst*) for this period.

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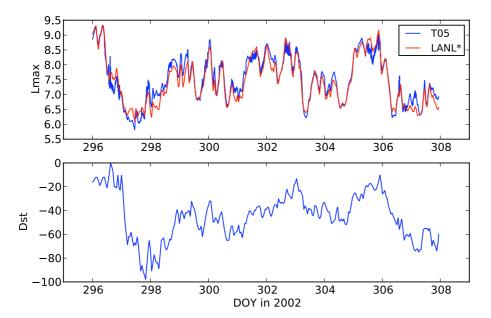


Fig. 7. Time series plot of the last closed drift shell $L_{\rm max}$ as it was calculated for solar wind conditions from 23 October 2002 to 4 November 2002 using TS05 (blue) and the LANL* neural network (red) for pitch angle $\alpha_{\rm loc} = 90^{\circ}$. Even for a moderate storm, the last closed drift shell can vary between $L^* = 6 - 9.5$.

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