

Sensitivity of predicted ultrafine particle size distributions in Europe to different nucleation rate parameterizations using PMCAMx-UF v2.2

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Abstract. The three-dimensional chemical transport model, PMCAMx-UF v2.2, designed to simulate the ultrafine particle size distribution, was used to investigate the impact of varying nucleation mechanisms on the predicted aerosol number concentration in Europe. Two basic case scenarios were examined: the original ternary H₂SO₄-NH₃-H₂O parameterization and a biogenic vapor-sulfuric acid parameterization. Using the organic-based parameterization, PMCAMx-UF predicted higher N_{10} (particle number above 10 nm) concentrations over Europe by 40 %-60 % on average during the simulated period, which is a relatively small difference, given the differences in the two assumed mechanisms. The low sensitivity of the particle number concentrations to the nucleation mechanisms used in this study may not exist in other regions outside Europe. Adjusting the nucleation rate by an order of magnitude for both mechanisms led to an average change of ± 30 % in N_{10} for the ternary ammonia case and -30% to 40% for the biogenic vapor case. In the biogenic organic nucleation scenario, reducing the fresh nuclei diameter from 1.7 to 1 nm resulted in reductions in N_{10} and N_{100} by -13% and -1%, respectively. Incorporating extremely low-volatility organic compounds (ELVOCs) as the nucleating species resulted in a predicted increase in the N_{10} concentration by 10 %–40 % over continental Europe compared to the ammonia parameterization. Model predictions were evaluated against field measurements from 26 stations across Europe during the summer of 2012. For N_{10} , the ternary ammonia and ELVOC-based parameterizations were in better agreement with the field data compared to the other tested mechanisms. In the case of N_{100} , all used parameterizations resulted in predictions that were consistent with the available field measurements.

1 Introduction

Aerosol nucleation and direct emissions from sources are the two principal processes for the introduction of new particles in the atmosphere. New particles formed by nucleation can either grow to larger sizes or can be lost by coagulation with existing particles (Kulmala et al., 2004; Merikanto et al., 2009; Pierce and Adams, 2009). New particle formation (NPF) through the condensation of vapors (e.g., sulfuric acid, organics, ammonia, and nitric acid) is estimated to be responsible for up to half of the global cloud condensation nuclei (CCN) and consequently affects considerably the cloud droplet number concentration (Adams and Seinfeld, 2002; Makkonen et al., 2009; Wang and Penner, 2009).

Various nucleation mechanisms have been proposed to describe the initial step of NPF. These mechanisms include sulfuric acid–water ($H_2SO_4-H_2O$) binary nucleation (Nilsson and Kulmala, 1998; Vehkamäki et al., 2002), sulfuric acid–ammonia–water ($H_2SO_4-NH_3-H_2O$) ternary nucleation (Bianchi et al., 2016; Kulmala et al., 2002; Napari et al., 2002; Yu, 2006), ion-induced nucleation (Jokinen et al., 2018; Kirkby et al., 2016; Laakso et al., 2002; Modgil et al., 2005), halogen oxide nucleation (Hoffmann et al., 2001), nucleation involving organic compounds (Li et al., 2019; Metzger et al., 2010; Weber et al., 2020), sulfuric aciddimethylamine nucleation (Yao et al., 2018), and iodine oxides (Sipilä et al., 2016). The corresponding nucleation rates depend on the sulfuric acid vapor concentration, with numerous studies indicating a strong correlation between sulfuric acid levels and the rate of new particle formation (Kuang et al., 2008; Lee et al., 2019; Sihto et al., 2006).

While NPF can be often explained by a simplified acidbase model (Chen et al., 2012) in sulfur-rich environments, model simulations (Anttila and Kerminen, 2003) and field measurements have showed that the condensation of sulfuric acid alone is often not enough to explain the observed growth rates of newly formed particles (Kuang et al., 2008). In environments with low sulfur dioxide levels, new particle growth has been linked to organic vapors (Olenius et al., 2018; Yli-Juuti et al., 2020). To explain the growth of the fresh nuclei, the condensation of organic species (Anttila and Kerminen, 2003) and heterogeneous reactions (Zhang and Wexler, 2002) have been proposed. Condensing lowvolatility organic vapors assist freshly formed particles in overcoming the Kelvin effect growth barrier which appears for particles with diameters of a few nanometers (Semeniuk and Dastoor, 2018).

Organic aerosol (OA) is an important constituent of submicrometer particulate matter, contributing more than 50 % in many locations around the world (Reyes-Villegas et al., 2021; Ripoll et al., 2015). Secondary organic aerosol (SOA) is formed during the oxidation of biogenic and anthropogenic volatile organic compounds (VOCs) and often accounts for most of the submicrometer OA (Hallquist et al., 2009; Jimenez et al., 2009; Schulze et al., 2017). VOCs of biogenic origin include terpenes such as isoprene (C₅H₈), monoterpenes (C10H16), and sesquiterpenes (C15H24) (Curci et al., 2009; Vermeuel et al., 2023). The oxidation of terpenes leads to highly oxygenated organic molecules (HOMs) that can participate in NPF and contribute to the growth of preexisting particles (Ehn et al., 2014; Jokinen et al., 2015; Weber et al., 2020). HOMs, sulfuric acid, and ammonia exhibit a synergistic effect in NPF and growth, while nitrogen oxides (NO_x) can suppress these processes, revealing complex interactions between biogenic and anthropogenic pollutants (Lehtipalo et al., 2018). Chemical transport models integrate our understanding of atmospheric processes and when combined with atmospheric measurements can help us evaluate if this understanding is satisfactory. There have been a number of efforts to simulate the ultrafine particle number concentration and NPF from ground-level and airborne observations (Leinonen et al., 2022; Lupascu et al., 2015; Matsui et al., 2013).

PMCAMx-UF is a three-dimensional regional chemical transport model (CTM) developed by Jung et al. (2010) specifically for simulating ultrafine particles. Baranizadeh et al. (2016) updated the nucleation parameterization in PMCAMx-UF by integrating the Atmospheric Cluster Dynamics Code, which is based on quantum chemical input

data. The observed number concentrations of particles larger than 4 nm could be reproduced within 1 order of magnitude for Europe at that stage, showing that there is room for improvement. Fountoukis et al. (2012) performed simulations over Europe and compared the model predictions against size distribution measurements from seven areas. The model successfully reproduced hourly number concentrations of particles larger than $10 \text{ nm} (N_{10})$ within a factor of 2 for more than 70% of the time. However, it regularly underpredicted the concentrations of particles larger than $100 \text{ nm} (N_{100})$ by 50 %. Notably, these early versions of the model did not account for SOA condensation on ultrafine particles. Patoulias et al. (2015) addressed this limitation by incorporating the condensation of organic vapors on nanoparticles through the development of a new aerosol dynamic model, DMANx (Dynamic Model for Aerosol Nucleation extended), demonstrating its significant impact on NPF. Julin et al. (2018) further extended the model by including the effects of amines on NPF and projected future changes in ultrafine particle emissions across Europe. The impact of secondary semi-volatile organic vapors on particle number concentrations was examined by integrating the volatility basis set (VBS) approach into PMCAMx-UF and applying the model over Europe (Patoulias et al., 2018). Including the VBS enabled the model to reproduce N_{10} and N_{100} ground measurements within a factor of 2 for 65 % and 70 % of the observations, respectively. The model was further enhanced to incorporate multiple generations of intermediate-volatility organic compound (IVOC) gas-phase oxidation, along with the formation and dynamic condensation of extremely low-volatility organic compounds (ELVOCs) from monoterpenes (Patoulias and Pandis, 2022).

Different nucleation parameterizations are used by global air quality and regional chemical transport models that present different parameter sensitivity. Riccobono et al. (2014) developed an empirical parameterization based on field measurements to describe the dependence of nucleation rates on sulfuric acid and oxidized biogenic compounds concentrations. Kirkby et al. (2016) found that highly oxidized organic compounds play a role in atmospheric particle nucleation comparable to that of sulfuric acid. Gordon et al. (2016) simulated the monoterpene HOMs formation using an empirical yield of HOMs during the oxidation of monoterpenes. Sartelet et al. (2022) simulated the heteromolecular nucleation of extremely low-volatility organic compounds (ELVOCs) from monoterpenes and sulfuric acid and reported improved predictive ability for suburban sites during the summer. Yu et al. (2020) developed a detailed kinetic nucleation model that includes H₂SO₄, H₂O, NH₃, and ions and created computationally efficient lookup tables that can be easily integrated into atmospheric chemical transport models.

The parameterizations of nucleation often involve adjusting the absolute nucleation rate with a nucleation tuner while maintaining its dependence on the concentrations of the participating vapors (Jung et al., 2010). Another important parameter is the initial nuclei diameter that is the size of newly formed particles. Paasonen et al. (2018) investigated particle growth in a boreal forest, highlighting the model's sensitivity to initial nuclei diameter variations, which substantially impacted growth dynamics and subsequent cloud condensation nuclei (CCN) formation.

In this study, we explore the impact of ammonia and organic vapor-based nucleation parameterizations on predicted particle number concentrations (e.g., N_{10} and N_{100}) and evaluate potential changes in model performance. Specifically, we investigate the effects of (a) altering the nucleation rate by an order of magnitude (both increase and decrease), (b) modifying the nuclei diameter, and (c) incorporating extremely low-volatility organic compounds (ELVOCs). Ground-level measurements from 26 European stations during the simulated period are used to evaluate PMCAMx-UF for the different used parameterizations.

2 Model description

The three-dimensional chemical transport model PMCAMx-UF simulates the chemically resolved mass distributions and particle number distributions down to the nanometer size range (Fountoukis et al., 2012; Jung et al., 2010; Patoulias and Pandis, 2022; Patoulias et al., 2018). PMCAMx-UF is based on the PMCAMx (Gaydos et al., 2007) air quality model that describes the processes of horizontal and vertical dispersion and advection, emissions, dry and wet deposition, aerosol dynamics and thermodynamics, and aqueous- and aerosol-phase chemistry. The simulation of the aerosol microphysics is handled in PMCAMx-UF by the updated version of the Dynamic Model for Aerosol Nucleation (DMANx), which simulates condensation, evaporation, new particle formation (NPF), and coagulation, assuming an internally mixed aerosol (Patoulias et al., 2015). DMANx is based on the TwO-Moment Aerosol Sectional (TOMAS) algorithm which independently tracks the aerosol number and mass distributions for each of the 41 logarithmically spaced size bins between 0.8 nm and 10 µm (Adams and Seinfeld, 2002). In each bin, the particle density is calculated and updated continuously as a function of the corresponding composition. Each successive size bin boundary has twice the mass of the previous one to simplify the simulation of coagulation. The lowest boundary is at 3.75×10^{-25} kg of dry aerosol mass per particle, corresponding to a dry diameter of 0.8 nm. The modeled particle components include ammonium, sulfate, nitrate, chloride, sodium, water, crustal material, elemental carbon, primary organic aerosol (POA), and eight surrogate SOA components.

In the current study, the base case nucleation rate was computed using a ternary $H_2SO_4-NH_3-H_2O$ parameterization, assuming a scaling factor of 10^{-7} (Fountoukis et al., 2012; Napari et al., 2002). For NH₃ concentrations below the threshold value of 0.01 ppt (parts per trillion), the binary

H₂SO₄-H₂O parameterization of Vehkamäki et al. (2002) was used. Coagulation is an important sink of the aerosol number in the atmosphere but also a mechanism by which freshly nucleated particles grow to larger sizes. Following Adams and Seinfeld (2002), the effects of gravitational settling and turbulence on coagulation are assumed negligible, and particles coagulate predominantly via Brownian diffusion. The coagulation coefficients were calculated based on the wet diameters of the particles, which were determined following the method of Gaydos et al. (2005). For smaller particles, the corrections of Dahneke (1983) for noncontinuum effects were used. The coagulation algorithm uses an adaptive time step, which does not allow an increase in the aerosol number or mass concentration in any size bin by more than an order of magnitude or a decrease by more than 25 % in each step.

During the last years, PMCAMx-UF has been extended to include the chemical aging of semi-volatile anthropogenic organic vapors, the oxidation of intermediate-volatility organic compounds (IVOCs), and the production of extremely low-volatility organic compounds (ELVOCs) by monoterpenes (Patoulias and Pandis, 2022). Additional information describing the evolution and evaluation of PMCAMx-UF model can be found in previous publications (Fountoukis et al., 2012; Jung et al., 2010; Patoulias et al., 2018; Patoulias and Pandis, 2022).

The extended Statewide Air Pollution Research Center (SAPRC) gas-phase chemical mechanism is used in PMCAMx-UF (Carter, 2000; ENVIRON, 2013). SAPRC contains 219 reactions of 64 gases and 18 free radicals. The SAPRC version used for the current study includes five lumped alkanes (ALK1–5), two lumped aromatics (ARO1 and ARO2), two lumped olefins (OLE1 and OLE2), a lumped monoterpene (TERP), isoprene (ISOP), and a lumped sesquiterpene species (SESQ).

A pseudo-steady-state approximation (PSSA) is used for the simulation of sulfuric acid vapor concentration. This allows a significant increase in the computational speed, with a minor loss in accuracy (Pierce and Adams, 2009). The condensation of ammonia on ultrafine particles is modeled following Jung et al. (2010) and ends when sulfate is entirely neutralized forming ammonium sulfate. The assumption that the system is always in equilibrium is used for the partitioning of nitric and hydrochloric acids (as nitrate and chloride, respectively) to particles in the accumulation mode range in PMCAMx-UF. In this version of PMCAMx-UF, the water content of the organic aerosol is neglected, and the aerosol water is associated with the inorganic aerosol components.

2.1 Nucleation mechanisms

PMCAMx-UF has the option of using a number of nucleation parameterizations (Baranizadeh et al., 2016; Fountoukis et al., 2012). In this work, we investigate two types of parameterizations, namely a ternary H_2SO_4 – NH_3 – H_2O parameter-

ization (ammonia parameterization) and a second including the products of the biogenic VOC oxidation or the H_2SO_4 – bSOA– H_2O parameterization (biogenic secondary organic aerosol (bSOA) parameterization). Several variations within these schemes are examined.

The ammonia parameterization, based on the scaled approach of Napari et al. (2002), has been the default parameterization in PMCAMx-UF and serves as the basis for our analysis of sulfuric acid-ammonia-water nucleation for easier comparison with the results of previous PMCAMx-UF applications. This approach was selected over the Baranizadeh et al. (2016) parameterization, which is one of the parameterizations available in PMCAMx-UF, because the latter has shown a tendency to overpredict concentrations of particles with diameters between 10 and 100 nm. In the base case, the selected nucleation tuner is equal to the value of 10^{-7} . The fresh nuclei diameter d_p ranges between 0.8 and 1.2 nm as a function of ammonia, sulfuric acid, temperature, and relative humidity (RH) (Napari et al., 2002). The parameterization is valid for temperatures between 240 and 300 K, RH of 5 %-95 %, ammonia mixing ratios of 0.1-100 ppt, sulfuric acid concentration of 10⁴-10⁹ molec. cm⁻³, and nucleation rates between 10^{-5} - 10^{6} cm⁻³ s⁻¹ (Napari et al., 2002).

The participation of biogenic secondary organic compounds in the nucleation mechanism together with sulfuric acid is based on the semi-empirical parameterization by Riccobono et al. (2014):

$$J_{1,7} = k \left[\text{BioOxOrg} \right] \left[\text{H}_2 \text{SO}_4 \right]^2, \tag{1}$$

where $J_{1.7}$ is the nucleation rate (in cm⁻³ s⁻¹) for particles with mobility diameter equal to 1.7 nm, k is a fitted parameter that was originally set equal to 3.27×10^{-21} molec.⁻³ cm⁶ s⁻¹, [BioOxOrg] is the concentration of monoterpene oxidation products (in molec. cm⁻³), and [H₂SO₄] is the concentration of sulfuric acid (in molec. cm⁻³) in the atmosphere. In this study, the Riccobono et al. (2014) parameterization replaced the Napari et al. (2002) parameterization in the corresponding simulations. Both parameterizations can be implemented simultaneously in PMCAMx-UF to examine their combined effects, but this is outside the scope of the present work.

The above parameterization needs to be adjusted to be compatible with the VBS parameters. PMCAMx-UF lumps all monoterpenes such as α -pinene, β -pinene, and limonene into one surrogate species. The monoterpene atmospheric oxidation products, using the VBS, are represented by four surrogate species with effective volatility at 298 K and $C^* = 1$, 10, 100, and 1000 µg m⁻³. We assume here that only the product with the lowest volatility ($C^* = 1 \mu g m^{-3}$) participates in new particle formation. This species is used effectively as a surrogate for the compounds with much lower volatility participating in the process. The sensitivity of our results to this choice will be examined in a subsequent section. To calculate the corresponding nucleation rate constant (instead of the value used in Eq. 1), we used the available nucleation rate measurements summarized in the work of Chen et al. (2012) and the corresponding maximum and minimum boundary of observed nucleation rates derived in that study. To get a zeroth-order estimation of an appropriate rate constant value, the predicted concentrations of sulfuric acid vapor and biogenic SOA ($C^* = 1 \ \mu g \ m^{-3}$) vapor during the PMCAMx-UF simulation were used to calculate the nucleation rate constant using Eq. (1). Least square fitting of the predicted nucleation rate to the average of the maximum and minimum boundaries of atmospheric measurements shown in Chen et al. (2012) yields a rate constant (k) of $0.1 \times 10^{-21} \ molec.^{-3} \ cm^6 \ s^{-1}$. This value was used in our PMCAMx-UF parameterization, resulting in Eq. (2):

$$J_{1.7} = 1 \times 10^{-22} \left[\text{bSOA}_{C_1^*} \right] [\text{H}_2 \text{SO}_4]^2,$$
(2)

where $bSOA_{C_1^*}$ corresponds to the concentration of the biogenic secondary organic vapor from the oxidation of monoterpenes with a saturation concentration (C^*) of $1 \mu g m^{-3}$ at 298 K. The use of this surrogate VBS species instead of the BioOxOrg of Riccobono et al. (2014) results in a different rate constant. Figure S1 in the Supplement depicts these boundaries of Chen et al. (2012) together with nucleation rates predicted by PMCAMx-UF, showing that our chosen constant renders the nucleation parameterization in Eq. (2) consistent with the available field measurements.

2.2 Description of sensitivity tests

A series of sensitivity tests have been performed for the ammonia and biogenic organic parameterizations described above (Table 1). To evaluate the impact of the absolute nucleation rates, we increased the rate constant by an order of magnitude for the ammonia and bSOA parameterizations in Cases 2 and 5. Similarly, the rate constant was also decreased by an order of magnitude in Cases 3 and 6. For the organic nucleation scenario, two additional cases have been investigated. In Case 7, the initial nuclei diameter was reduced from 1.7 to 1 nm.

In Case 8, extremely low-volatility organic compounds (ELVOCs) with a saturation concentration (C^*) of $10^{-5} \mu g m^{-3}$ were introduced as the organic component in the nucleation mechanism. The ELVOCs are assumed to be produced by the oxidation of monoterpenes with a yield of 5 % (Patoulias and Pandis, 2022; Rissanen et al., 2014) and a rate constant (k) of 1.0×10^{-21} molec.⁻³ cm⁶ s⁻¹, with the corresponding nucleation rate being calculated by

$$J_{1.7} = 1 \times 10^{-21} \left[\text{bSOA}_{C_{10^{-5}}^*} \right] [\text{H}_2 \text{SO}_4]^2.$$
(3)

3 Model application

The modeling domain of PMCAMx-UF in this application covers a $5400 \times 5832 \text{ km}^2$ region in Europe, with a grid resolution of $36 \times 36 \text{ km}$ and 14 vertical layers extending up to

 Table 1. Nucleation parameterization scenarios.

	Case 1	Case 2	Case 3	Case 4	Case 5	Case 6	Case 7	Case 8
Third species	Ammonia	Ammonia	Ammonia	Organic	Organic	Organic	Organic	Organic
$C^* (\mu g m^{-3})$				1	1	1	1	10^{-5}
$k \text{ (molec.}^{-3} \text{ cm}^6 \text{ s}^{-1}\text{)}$	$10^{-7^{a}}$	$10^{-6^{a}}$	$10^{-8^{a}}$	10^{-22}	10^{-21}	10^{-23}	10^{-22}	10^{-21}
Particle diameter (nm)	0.8-1.2	0.8-1.2	0.8-1.2	1.7	1.7	1.7	1	1.7

^a Nucleation tuner for ternary ammonia nucleation (dimensionless).

7.5 km in a terrain-following grid (Table S1 in the Supplement). The modeling period focuses on the PEGASOS campaign and includes a total of 34 d in 2012, starting on 5 June until 8 July 2012.

A rotated polar stereographic map projection was used for the simulations by PMCAMx-UF to simulate all of Europe efficiently. This projection minimizes distortions in spatial representation across the region and the rotation ensures that Europe is accurately represented while maintaining consistency in horizontal grid spacing, which is essential for atmospheric modeling (ENVIRON, 2013). To minimize the effect of the initial conditions on the results, the first 2 d of each simulation were excluded from the analysis. Relatively low and constant values have been used for the boundary conditions allowing the predicted particle number concentrations over central Europe to be determined by the emissions and corresponding processes simulated by the model. The boundary conditions and their effects on the predicted number concentrations by PMCAMx-UF in this domain have been discussed in previous publications (Patoulias et al., 2018; Patoulias and Pandis, 2022).

Meteorological inputs to PMCAMx-UF include temperature, pressure, horizontal wind components, water vapor, vertical diffusivity, clouds, and rainfall. The above inputs correspond to hourly data and were generated by the Weather Research and Forecasting (WRF) model (Skamarock et al., 2005). The WRF was driven by geographical and dynamic meteorological data generated by the Global Forecast System (GFSv15) of the National Oceanic and Atmospheric Administration/National Centers for Environmental Prediction. The layers of WRF and PMCAMx-UF were aligned with each other with the two models using the same coordinate. The WRF simulation was periodically re-initialized every 3 d with observed conditions to ensure accuracy in the corresponding fields used as inputs in PMCAMx-UF. Each field was provided with fidelity appropriate to the chosen grid resolution of the model as the measurements were pre-processed by the WPS (WRF Preprocessing System) package.

The particle emissions were based on the pan-European anthropogenic particle number emission inventory and the carbonaceous aerosol inventory (Kulmala et al., 2011) developed during the European Integrated project on Aerosol, Cloud, Climate, and Air Quality Interactions (EUCAARI) project. The resulting number and mass inventories contain number emissions and consistent size-resolved composition for particles over the size range of approximately 10 nm to 10 μ m. The frequency of output of PMCAMx-UF is selected by its user. Hourly output was used in the present simulations.

Measurements

The model results were compared against measurements in 26 ground sites, which are available in the European Supersites for Atmospheric Aerosol Research (EUSAAR), EBAS databases (https://ebas.nilu.no, last access: 23 February 2025), and the Aerosol, Clouds and Trace Gases Research Infrastructure (ACTRIS) (https://actris.nilu.no, last access: 23 February 2025). Particle size distribution measurements at all sites were made using either a differential mobility particle sizer (DMPS) or a scanning mobility particle sizer (SMPS). Information about all the measurement stations can be found in Table S2 in the Supplement.

4 Results

4.1 Base case ammonia and organic parameterizations

The average ground level (first vertical layer) number concentrations for both base case nucleation parameterizations are shown in Fig. 1. For the ammonia parameterization, the N_{tot} and N_{10} have the highest concentrations in the Iberian Peninsula, the Netherlands, Poland, and Türkiye due to nucleation. For N_{50} and N_{100} , the highest concentrations are predicted in the Balkans and the Mediterranean Sea due to the high emissions of sulfur dioxide in the surrounding areas and the intense photochemistry. High N_{50} and N_{100} values are also predicted in Poland, Russia, and Ukraine due to urban and industrial emissions. Nucleation is predicted to increase the total average number concentration by 160 %. For N_{10} and N_{100} , the enhancement due to nucleation was 140 % and 45%, respectively. The predicted ammonia concentration exceeded 8 ppb in Germany, the Netherlands, France, northern Italy, Poland, and Russia, as shown in Fig. 2a, primarily due to intensive agricultural activities in these regions. Figure 2b presents the average sulfuric acid concentration, which, unlike ammonia, was higher over marine areas such as the Mediterranean Sea and particularly the Aegean Sea, as well as coastal regions of the Atlantic Ocean including the Portuguese, Spanish, and French coasts. These elevated levels are attributed to significant SO₂ emissions from maritime shipping activities and high OH levels in these high relative humidity sunny regions. Figure 2c depicts the average predicted nucleation rate, with values exceeding $1 \text{ cm}^{-3} \text{ s}^{-1}$ in parts of Portugal, northern Spain, the United Kingdom, the Balkans, Türkiye, Poland, and Russia. In contrast, the average nucleation rate across the remainder of Europe generally remained below $0.2 \text{ cm}^{-3} \text{ s}^{-1}$.

For the biogenic parameterization, the predicted N_{tot} and N_{10} have the same spatial patterns as with the ammonia parameterization but with higher predicted levels especially in Italy, Russia, the Balkans, and parts of the Mediterranean Sea (Fig. 1e–f). The highest predicted concentrations of N_{50} and N_{100} are almost identical with those predicted by the ammonia parameterization (Fig. 1c-d). When these predictions were compared to the no-nucleation scenario, the enhancement attributable to nucleation in this simulation was approximately 300 % for N_{tot} , 180 % for N_{10} , and 50 % for N_{100} . The gas-phase concentration of the bSOA component with $C^* = 1 \,\mu \text{g} \,\text{m}^{-3}$ was predicted to be elevated in forested regions of central and northern Europe, including the Scandinavian countries, northern Russia, and Georgia (Fig. 2d). The average sulfuric acid concentration remained similar to that in the previous case. Incorporating bSOA ($C^* = 1 \,\mu g \, m^{-3}$) as a third species resulted in an increased nucleation rate, with higher average values (above $4 \text{ cm}^{-3} \text{ s}^{-1}$) in regions such as Portugal, northern Spain, the Mediterranean Sea, Greece and the Aegean Sea, the Balkans, Türkiye, Poland, and Russia (Fig. 2f). Despite the relatively low concentration of semivolatile biogenic organics in the Mediterranean Sea region, the high concentrations of sulfuric acid resulted in an elevated predicted nucleation rate (Fig. 2e).

The PMCAMx-UF number concentration predictions using the biogenic nucleation parameterization are higher than those predicted using the ammonia parameterization in most areas. More specifically, the predicted N_{tot} is 80%–150% higher, and the N_{10} is 30%–60% higher in regions with intense nucleation (Fig. 3a–b, e–f). On the other hand, N_{tot} decreased by approximately 25% and N_{10} by 10% in southern England, northern France, and the Netherlands when the organic parameterization replaced the ammonia one. N_{50} increased by about 10% in Greece and Russia, while it decreased by about 10% in the United Kingdom and Germany (Fig. 3g). The changes in N_{100} were minor, ranging from 5%–10% across the European domain (Fig. 3h).

4.2 Evaluation of the model

The predictions of the two simulations (ammonia and bSOA base case parameterizations) were compared against hourly N_{10} and N_{100} field measurements. The overall hourly normalized mean bias (NMB) for N_{10} was found to be -16% for ammonia and 2% for the bSOA case, while the N_{100} NMB

was close to 7 % for both cases (Fig. 4). The overall normalized mean error (NME) for the N_{10} was 54 % in the ammonia case and 61 % in the bSOA one. This indicates that the overall performance of the two parameterizations is comparably effective. Despite their inherent differences, both parameterizations demonstrated robust performance, even when evaluated on an hourly basis. At 15 stations (ANB, CBW, DSN, DSW, HYY, ISP, KPU, MLP, PRG, USM, VSM, WLD, ZUG, PAT, and SPC) (see Table S2 for definitions), the NMB for N_{10} is lower in the case of the bSOA parameterization than in the case with ammonia. However, there are six stations (ASP, FNK, GDN, VRR, NEO, and THE) for which the use of biogenic organic nucleation significantly increases the N_{10} NMB compared to the ammonia case (Fig. 4a). In both simulations, the N_{10} NME remains below 60 % for most of the stations, with a difference of less than 5% between the two cases. Notably, the predictions in six stations (ASP, FNK, GDN, VRR, NEO, and THE) had significantly lower NME values for the ammonia mechanism (Fig. 4b). Of these, three stations are located in Greece, one in Malta, one in Finland, and one in Sweden.

For N_{100} , the NME for both cases was similar and equal to 48 %. The hourly N_{100} NMB for all stations ranged between -40 % and 80 % (Fig. 5a). No significant differences (less than 10 %) appear in the NMB of N_{100} for the two simulations, with the only exceptions being those of the ASP, FNK, and THE stations. For the FNK and THE stations in Greece, the ammonia parameterization shows less of an error, while the opposite is the case for ASP in Sweden (Fig. 5b).

4.3 Results of sensitivity tests

4.3.1 Effect of scaling the ammonia and biogenic nucleation rate parameterizations

For the ammonia parameterization, two additional cases were investigated (Cases 2 and 3; Table 1). In Case 2, an increase in the nucleation rate by a factor of 10 caused a 70 %– 100 % increase in N_{tot} (4000–6000 cm⁻³) and a 40 %–60 % increase in N_{10} (over 2000 cm⁻³) in the regions with intense nucleation like the Iberian Peninsula, central Europe, the Balkans, and Türkiye (Fig. S2). For N_{50} an increase of about 10 %–20 % (300–500 cm⁻³) was predicted in the Balkans, eastern Mediterranean, Poland, and Russia. For N_{100} , the change was small (less than 10 %) with the most significant increase of 5 %–8 % in the Balkans, the eastern Mediterranean Sea, and Russia (Fig. S2).

In Case 3, a reduction by a factor of 10 in the nucleation rate resulted in an overall reduction in all investigated number concentrations for the modeled domain. A 40%–60% reduction in N_{tot} (about 2000–3000 cm⁻³) and a 30%–40% decrease in N_{10} (over 15 000 cm⁻³) was predicted in the regions with intense nucleation (Fig. S3). The N_{50} decreased by about 15%–20%, mainly in the Balkans, Mediterranean, eastern Europe, Türkiye, and parts of Scandinavia. For N_{100} ,

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Figure 1. Average ground level number concentrations (in cm⁻³) for the ternary ammonia nucleation simulation during 5 June–8 July 2012 for (a) all particles (N_{tot}) and particles above (b) 10 nm (N_{10}), (c) 50 nm (N_{50}), and (d) 100 nm (N_{100}). Average ground level number concentrations (in cm⁻³) for the biogenic semi-volatility organic nucleation simulation during 5 June–8 July 2012 for (e) all particles (N_{tot}) and particles above (f) 10 nm (N_{10}), (g) 50 nm (N_{50}), and (h) 100 nm (N_{100}). Different scales are used.



Figure 2. Ground level average concentration of (**a**) ammonia (NH₃) (in ppb) and (**b**) sulfuric acid (in ppt) and (**c**) the nucleation rate J (in cm⁻³ s⁻¹) for the ternary ammonia nucleation. Ground level average mass concentration of (**d**) biogenic semi-volatility secondary organic compounds with $C^* = 1 \,\mu \text{g m}^{-3}$ (in ppt) and (**e**) sulfuric acid (in ppt) and (**f**) the nucleation rate J (in cm⁻³ s⁻¹) for the organic nucleation during 5 June–8 July. Different scales are used.

there was a 5%–10% decrease in the Balkans, Russia, and the eastern Mediterranean. PMCAMx-UF predicted a 10% increase in N_{100} in the United Kingdom (Fig. S3).

The increase in the biogenic nucleation rate by a factor of 10 in Case 5 resulted in a significant increase of 150 %–200 % for the N_{tot} (15 000–20 000 cm⁻³) in the areas with intense nucleation and a 50 %–70 % increase in N_{10} (over 3000 cm⁻³) in western Europe, Türkiye, and Scandinavia

(Fig. S4). In the case of N_{50} , there was an increase of about 15 %–20 % in the regions of Scandinavia and northern Russia and 10 %–15 % in the eastern Mediterranean. For N_{100} , there was a small increase for almost all of the domain, with a peak change of 5 %–8 % in the Balkans and Türkiye (Fig. S4).

The reduction by a factor of 10 in the nucleation rate in Case 6 led to a 50%–70% reduction in N_{tot} (5000– 7000 cm⁻³) and a 35%–50% reduction in N_{10} (2500–



Figure 3. Average ground change [biogenic–ammonia parameterization] of the number concentration (in cm⁻³) (**a**–**d**) and fractional increase (f_{N_x}) of the number concentration (in %) (**e**–**h**) during 5 June–8 July 2012 for (**a**, **e**) all particles ($f_{N_{tot}}$), (**b**, **f**) particles above 10 nm ($f_{N_{10}}$), (**c**, **g**) particles above 50 nm ($f_{N_{50}}$), and (**d**, **h**) particles above 100 nm ($f_{N_{100}}$). Different scales are used. Positive values indicate higher concentrations for the biogenic vapor case.



Figure 4. The hourly (a) normalized mean bias (NMB) (in %) and (b) normalized mean error (NME) (in %) of N_{10} for 26 stations. Blue bars are used for the simulation with ternary ammonia nucleation and red bars for the biogenic parameterization.



Figure 5. The hourly (a) normalized mean bias (NMB) (in %) and (b) normalized mean error (NME) (in %) of N_{100} for 26 stations. Blue bars are used for the simulation with ternary ammonia nucleation and red bars for the biogenic parameterization.

 3500 cm^{-3}) for the entire simulated area (Fig. S5). For N_{50} , there was a decrease of about 20%-25% in Scandinavia and northern Russia and about 15%-20% reduction in the eastern Mediterranean. In the case of N_{100} , there was a small decrease of 5%-8% in the eastern Mediterranean Sea (Fig. S5).

4.3.2 Effect of the initial nuclei diameter in the biogenic nucleation parameterization

The reduction in the nuclei diameter from 1.7 to 1 nm in Case 7 resulted in a 25%–35% reduction in N_{tot} (2500–3500 cm⁻³) and a 20%–25% decrease in N_{10} (1500–2000 cm⁻³) in the Balkans, Poland, and Russia, where intense nucleation events were predicted (Fig. S6). For N_{50} and

 N_{100} , a reduction of about 5% is predicted. The reduction in the nuclei diameter mainly affects the number of particles between 1–10 nm. The smaller initial diameter leads to an acceleration in coagulation and leads to faster losses of those fresh particles. For this reason, a significant reduction in N_{tot} is predicted in the eastern Mediterranean Sea and the Balkans, where the highest concentrations of the largest (N_{50} and N_{100}) particles are found.

The spatial variability in the average and fractional changes in the number concentration of N_{1-10} particles (reflecting nucleation rates), as well as the condensational sink (CS) and coagulation sink, resulting from a decrease in nuclei diameter to 1 nm is shown in Fig. S7. Reducing the nuclei diameter from 1.7 to 1 nm decreases the coagulation sink by 8%–14% in regions experiencing intense nucleation events, while the condensation sink remains largely unchanged. Across most of the domain, N_{1-10} concentrations decrease, with reductions ranging from 30%–70% in southeastern Europe.

The decrease in the concentration of 1–10 nm particles (Fig. S7) and N_{10} particles (Fig. S6) is primarily driven by an increased probability of coagulation for newly formed particles. This, in turn, significantly reduces the likelihood that these particles will grow large enough to survive, as growth strongly depends on interactions with pre-existing particles. Pierce and Adams (2007) demonstrated that, under most conditions, condensation is the dominant growth mechanism, while coagulation with larger particles acts as the primary sink for ultrafine particles. Consistent with this, our study finds that the probability of a new ultrafine particle growing to generate a CCN can vary widely from less than 0.1 % to approximately 90 %, depending on atmospheric conditions.

4.3.3 Effect of ELVOCs in nucleation

In this case, the semi-volatile biogenic organics ($C^* =$ $1 \,\mu g \, m^{-3}$) were substituted by the biogenic ELVOCs ($C^* =$ $10^{-5} \,\mu g \,m^{-3}$) in the parameterization. This was accompanied by an increase in the scaling factor from 10^{-22} to 10^{-21} molec.⁻³ cm⁶ s⁻¹. This modification resulted in a predicted increase of 40 %–100 % in N_{tot} (2000–4000 cm⁻³) and a 10%–40% increase for N_{10} (500–2000 cm⁻³) compared to the base bSOA parameterization across regions including Portugal, northern France, the United Kingdom, Germany, Poland, southern Scandinavia, the Balkans, and Russia. Conversely, a reduction of approximately 30 % in N_{tot} and 20 % in N_{10} is predicted for the Mediterranean region (Fig. 6). For N_{50} , an increase of 5 %–10 % (100–200 cm⁻³) was predicted in Poland and Scandinavia, while a slight decrease of 5% is shown for the Mediterranean Sea. The change in N_{100} was less than 10%, with the most significant differences occurring in Portugal, Türkiye, Scandinavia, and the United Kingdom.

For the case of the sulfuric acid ELVOC nucleation, high nucleation rates are predicted in the United Kingdom, Portugal, northern Spain, northern Italy, Poland, the Balkans, Türkiye, and Russia (Fig. 7c). In these areas, there are high concentrations of ELVOCs and sulfuric acid according to PMCAMx-UF (Fig. 7a, b).

4.4 Evaluation of all simulation cases

A scenario excluding nucleation has been included in the evaluation for comparative purposes. This no-nucleation scenario significantly underestimates N_{10} concentrations, whereas the incorporation of nucleation significantly improves model predictions across all investigated cases (Cases 1–8; Fig. 8). All simulations with nucleation result in predicted distributions of N_{10} concentrations that are consistent with the observed measurement range. The exception is the scaled up biogenic–sulfuric acid parameterization (Case 5) that overpredicts the N_{10} concentrations in a lot of the stations. The median observed concentration of N_{10} is close to Cases 1 and 2, both of which employ ammonia as a third species, but also Cases 4, 7, and 8, which are based on biogenic organic vapors.

The N_{50} concentrations are clearly underestimated in the no-nucleation simulation. In both scenarios in which the nucleation rate was reduced by an order of magnitude (Cases 3 and 6), the predicted N_{50} concentration is closer to the measurements in terms of median N_{50} and the range of values (Fig. 8b). The remaining cases (1, 2, 4, 5, 7, and 8) overestimate the median N_{50} ; however, the corresponding ranges of values are close to the measurements.

In the case of N_{100} , the no-nucleation case significantly underestimates its concentration. Conversely, in all nucleation tests, the predicted median N_{100} concentration is close to the measurement values. At the same time, PMCAMx-UF predicts a broader range of N_{100} values for Cases 1–8 in relation to the measurements (Fig. 8c).

The no-nucleation simulation underestimates all number concentrations with a NMB of -60% for N_{10} , a NMB of -30% for N_{50} , and a NMB of -27% for N_{100} (Fig. 9). Cases 1, 2, 4, 7, and 8 exhibit a NMB of $\pm 20\%$ for N_{10} , and Cases 3 and 6 (both involving a 10-fold reduction in nucleation rate) show a NMB between -30% and -40% (Fig. 9a). Case 5, which involves bSOA and an increased nucleation factor, has the highest NMB of all at 50%.

For N_{50} , the simulations in which the nucleation rate was reduced by a factor of 10 exhibit the lowest NMB which was close to zero. The cases where the nucleation rate was increased by 10 times (Cases 2 and 5) presented the maximum NMB among all the simulated scenarios with a NMB of 22 % and 30 %, respectively (Fig. 3b). For the rest of the cases (1, 4, 7, and 8), the NMB varies between 0 % and 20 % (Fig. 9b).

For N_{100} , all cases incorporating ammonia or bSOA nucleation exhibit a NMB of less than 10%. The cases in which the nucleation rate was reduced by an order of magnitude (Cases 3 and 6) demonstrate the lowest NMB which was close to zero (Fig. 9c).



Figure 6. Average ground increase in the number concentration (in cm⁻³) (**a**–**d**) and fractional increase (f_{N_x}) of number concentration (in %) (**e**–**h**) for case 8 (ELVOCs as third species) of organic nucleation during 5 June–8 July 2012 for (**a**–**e**) all particles $(f_{N_{tot}})$ and (**b**–**f**) particles above 10 nm $(f_{N_{10}})$, (**c**–**g**) particles above 50 nm $(f_{N_{50}})$, and (**d**–**h**) particles above 100 nm $(f_{N_{100}})$. Different scales are used.



Figure 7. Ground level average gas concentration for Case 8 of (**a**) extremely low-volatility organic compounds (ELVOCs) with $C^* = 10^{-5} \,\mu \text{g m}^{-3}$ (in ppt) and (**b**) sulfuric acid (in ppt) and (**c**) the nucleation rate J (in cm⁻³ s⁻¹) for the organic nucleation during 5 June–8 July. Different scales are used.

The normalized mean error (NME) for N_{10} ranges between 50% and 60% for nearly all examined parameterizations, with the only exception being Case 5 (biogenic and increased scenario) for which NME exceeds 80% (Fig. S8a). For N_{50} , the lowest NME was found for the reduced scaling factor for ammonia (Case 3) and biogenic (Case 6) parameterization (Fig. S8b). Regarding N_{100} , all cases presented a NME of less than 50% (Fig. S8c).

Soccer plots, which depict fractional bias as a function of fractional error, are utilized to illustrate model performance (Morris et al., 2005). In Fig. 10, the performance of PMCAMx-UF is shown for the examined parameterizations and for all measurements in all stations using the daily temporal resolution. For the no-nucleation scenario, the model performance for N_{50} and N_{100} was average ($F_{\text{bias}} < \pm 60 \%$ and $F_{\text{error}} < \pm 75 \%$). However, for the N_{10} , the performance fell outside this range, indicating the fundamental errors and underscoring the necessity of incorporating nucleation processes for accurate N_{10} prediction. For N_{10} , the ammonia (Case 1) and biogenic (Case 4) parameterization, along with the nuclei size adjustment (Case 7) and the use of ELVOCs (Case 8), show good performance. The scenarios involving scaling factor adjustments (either increased or decreased by an order of magnitude) border on the good and excellent performance regions (Cases 2, 3, 5, and 6). For N_{50} and N_{100} , all eight investigated parameterizations have good performance ($F_{\text{bias}} < \pm 30\%$ and $F_{\text{error}} < \pm 50\%$) and are very close to the criteria for excellent performance ($< \pm 30\%$ and $F_{\text{error}} < \pm 50\%$).

The performance of PMCAMx-UF for various cases was also analyzed using the soccer plots for each 1 of the 26 sites across Europe using once more daily temporal resolution (Fig. S9). The PMCAMx-UF performance for N_{100} for most stations is good or excellent for the ammonia and biogenic organic nucleation cases. For N_{10} , the ammonia (Case 1) parameterization performs a little better than the biogenic cases (4 and 8).



Figure 8. Measurements from 26 ground stations, against the simulation without nucleation, the ammonia ternary parameterization (Case 1) and the change by an order of magnitude in the scaling factor (Cases 2 and 3); the biogenic parameterization (Case 4) with the change by an order of magnitude in the scaling factor (Cases 5 and 6); the decrease in the nuclei diameter (Case 7); and the ELVOC addition as the third species (Case 8) for (a) N_{10} , (b) N_{50} , and (c) N_{100} . The lower and upper lines in each box represent the 25% and 75% of the results, respectively, while the middle line corresponds to the median value.

The lack of sensitivity to the nucleation parameterization was also observed at higher altitudes. The model was additionally evaluated against airborne measurements obtained from the Pan-European Gas–AeroSOI–climate interaction Study (PEGASOS) campaign. The PEGASOS dataset includes vertical aerosol measurements conducted aloft using a Zeppelin over the Po Valley (Italy). The predicted vertical profiles for particle concentrations, simulated using the two nucleation parameterizations, were very similar and in good agreement with the observations, highlighting the consistency of the model predictions aloft (Fig. S10).

5 Conclusions

In this study, we considered two nucleation parameterizations involving sulfuric acid and water, namely one in which ammonia was the third reactant and one in which semivolatile biogenic organics participated in the critical cluster. The parameters of both expressions were selected so that the predicted rates would be generally consistent with available ambient nucleation rate measurements. Nucleation enhanced the N_{tot} by 160 %–300 %, the N_{10} by 140 %–180 %, and the N_{100} by 45 %–50 % during the simulated period.



Figure 9. The NMB for hourly (a) N_{10} , (b) N_{50} , and (c) N_{100} for the no-nucleation scenario and the ammonia ternary parameterization (Case 1); the change by an order of magnitude in the scaling factor (Cases 2 and 3); the biogenic parameterization (Case 4) with the change by an order of magnitude in the scaling factor (Cases 5 and 6); the decrease in the nuclei diameter (Case 7); and the ELVOC addition as the third species (Case 8).

The base case organic parameterization, when implemented in PMCAMx-UF, tended to predict higher N_{10} concentrations over Europe that were, on average, 40 %–60 % higher compared to the ammonia case. This is a relatively small difference, given the substantial differences between the two nucleation mechanisms. The biogenic organic parameterization predicted values of N_{10} that were 30 %–50 % higher over the Mediterranean, more than 50 % higher in Russia, and 20 % higher in Scandinavia compared to the ammonia parameterization predictions. There were a few areas in central and western Europe in which the opposite was true, with 20 % lower N_{10} values predicted when the biogenic organic parameterization was used.

Despite the significant differences in the used parameterizations, the average predicted N_{100} concentrations over the domain differed by less than 5 %. This suggests surprisingly low sensitivity of the current concentrations of these larger particles (a proxy for CCN) to the details of the nucleation mechanism, provided the parameterizations are consistent with the available ambient observation dataset.

Both parameterizations demonstrated good performance on average against hourly measurements at 26 stations, with similar accuracy. The simulation with ternary ammonia nucleation had a NMB for N_{10} of -16% and for N_{100} equal to 6%. The performance for the biogenic organic parameterization had a lower NMB of 2% for N_{10} but a little higher (8%) for N_{100} . The relatively low NMB is partially due to the tendency of the model to overpredict at some stations and underpredict at others, leading to some cancellation of biases



Figure 10. Model evaluation using the fractional error (%) versus fractional bias (%) of daily number concentrations for (**a**) N_{10} , (**b**) N_{50} , and (**c**) N_{100} for the no-nucleation scenario, the ammonia ternary parameterization (Case 1); the change by an order of magnitude in the scaling factor (Cases 2 and 3); the biogenic parameterization (Case 4) with the change by an order of magnitude in scaling factor (Cases 5 and 6); the decrease in the nuclei diameter (Case 7); and the ELVOC addition as the third species (Case 8).

in the overall average. The NMEs in N_{10} for both simulations were below 60% for most of the stations and were quite similar for the two parameterizations. While this study finds limited sensitivity of particle number concentrations to the nucleation mechanism in the European region, this conclusion may not be applicable in vastly different environments like tropical rainforests, oceans, deserts, or polar regions.

Modifying the ammonia nucleation rate parameterizations by an order of magnitude led to average changes in predicted N_{10} concentrations by ± 30 % and N_{100} by -5 % to 2 %. Similar adjustments in biogenic aerosol nucleation rates resulted in average changes from -30 % to 40 % for N_{10} and from -5 % to 2 % for N_{100} . Decreasing the nuclei diameter for biogenic organic nucleation from 1.7 to 1 nm caused a significant decrease in N_{10} , particularly over the Mediterranean Sea and central Europe, with average changes of -20 %. Incorporating ELVOCs as a third species resulted in an average change of 3 % in N_{10} and 0.4 % in N_{100} , aligning well with observed number concentrations at most stations. These adjustments represent relatively modest differences, given the divergent nucleation mechanisms involved.

Code availability. The model code base used to generate the results for ammonia ternary nucleation (PMCAMx-UF version 2.1) can be found on Zenodo at https://doi.org/10.5281/zenodo.10078189 (Pandis and Patoulias, 2023). The model code base used to generate the results for biogenic nucleation (PMCAMx-UF version 2.2) can be found on Zenodo at https://doi.org/10.5281/zenodo.12720811 (Patoulias and Pandis, 2024). The analysis codes and data used to prepare the paper can be found on Zenodo at https://doi.org/10.5281/ zenodo.13348332 (Patoulias et al., 2024).

Data availability. Data will be made available on request to the corresponding author.

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