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Supplement of

Improving Antarctic Bottom Water precursors in NEMO for climate applications

Katherine Hutchinson et al.

Correspondence to: Katherine Hutchinson (kath.hutchinson@gmail.com)

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1 Supplementary Material

2 S1 A note on the NEMO namelists

3

4 In the Zenodo data repository associated with this manuscript (10.5281/zenodo.7561767), the NEMO reference namelist
5 (namelist_ref), “Open” configuration namelist (namelist_core_ia_cfg) and sea ice namelists (namelist_ice_ref and
6 namelist_ice_cfg) are given. The reference namelist is the default provided with the NEMO code. Unless stated otherwise in
7 the “cfg”, the simulation uses the choices selected in the “ref” namelist. The namelist_core_ia_cfg is specific to a global
8 ocean configuration (with modifications adapted to eORCA1) forced by interannual core winds. For more information on all
9 the parameters included in these namelists, please refer to the NEMO reference manual available on Zenodo
10 (10.5281/zenodo.6334656). Of specific interest may be Chapter 6.10 on “Interaction with ice shelves (ISF)” where the
11 various options to represent ice-shelf/ ocean fluxes, heat and salt exchange coefficients and melt parameterization choices are
12 explained.

13

14 The differences in namelist_core_ia_cfg for the “Open” and “Closed” cavity runs are listed in Table S1. Note that these
15 differences are minor as the adaptations are made mostly to the input files (explained under “DOMAIN FILES AND
16 INITIAL CONDITIONS” in Zenodo data repository description 10.5281/zenodo.7561767).

17

Namelist parameter	Brief description	Closed	Open
ln_isfcav_mlt	ice shelf melting into the cavity	false	true
sn_isfpar_fwf	namelist block for freshwater flux from ice shelf melt	Read in fixed freshwater flux from melt as estimated from Depoorter et al. (2013) for all ice shelves	Use a new input file where this freshwater flux is removed in front of FRIS, LCIS and ROSS.
ln_hpg_sco	s-coordinate formulation of pressure gradient	true	false
ln_hpg_isf	s-coordinate formulation of pressure gradient adapted for under ice shelves.	false	true

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19

20 Table S1 Namelist differences when FRIS, LCIS and ROSS cavities are open.

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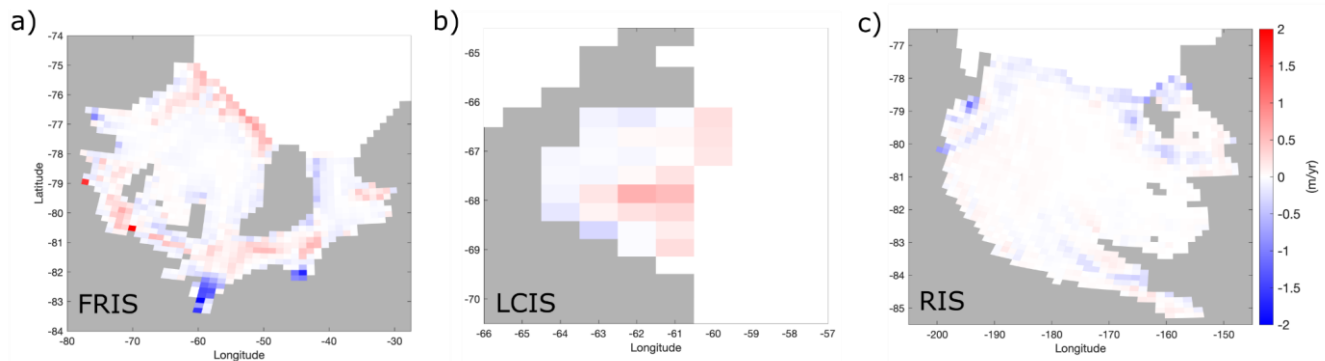
22 **S2 An alternative methodology to parameterize the effect of tides under the ice shelves**

23 **S2.1 Rationale**

24 The influence of tides on ice shelf basal melt is parameterized in NEMO using a constant background kinetic energy, set to
25 the value of $2.5 \times 10^{-3} \text{ m}^2 \text{ s}^{-2}$ everywhere (namelist parameter `rn_ke0`). As discussed in Jourdain et al. (2019), within ice shelf
26 cavities tides play an important role in modulating basal melt by imposing an added current velocity along the ice shelf base.
27 The magnitude of the tidal currents are, however, not constant everywhere, and so a single kinetic energy value (as is the
28 default option in NEMO) can be improved upon by using a two dimensional field. To inform this, we follow the
29 methodology of Jourdain et al. (2019) and use the Circum-Antarctic Tidal Simulation CATS2008 tidal map interpolated onto
30 the eORCA1 grid (Howard et al., 2019). Additionally, some of the NEMO code had to be adapted to allow for this type of
31 tidal parameterization and so the following files were amended: `isf_oce.F90`, `isfcavgam.F90`, `isfstp.F90`, `zdfdrd.F90`. The
32 simulation was run for 124 years and the differences in melt rate between this simulation and the reference “Open” cavity
33 simulation are presented in Fig. S1.
34

35 **S2.2 Impact of alternative tidal parameterization on basal melt**

36
37 Using the two-dimensional CATS tidal atlas to parameterize the effect of tides slightly increases melt for FRIS (total mean
38 melt flux over 1995-2009 of $120 \pm 22 \text{ Gt/yr}$) and LCIS ($39 \pm 8 \text{ Gt/yr}$) and reduces net melt for ROSS ($102 \pm 18 \text{ Gt/yr}$)
39 compared to results shown in Table 1. In general, the tidal velocities for CATS under FRIS and LCIS are faster than the
40 default constant and for RIS are slower. The spatial differences in yearly basal melt rate can be seen in Fig. S1. The marked
41 differences for FRIS are an increase in melt at the ice shelf front and a decrease within the deep fjords along the grounding
42 line. An explanation for this is that the elevated tidal velocities increase the rate of melting as warm offshore water enters the
43 cavity, causing elevated melt along the western ice shelf front. This water then loses its heat, and thus potential for melt, and
44 slows down as it travels into the southernmost extremities of the cavity where it induces less melt than in the default
45 simulation. The converse is true for RIS where the CATS tidal map shows slower induced velocities than the default
46 parameterization, meaning a decrease in the melt rate all along the ice shelf front. To explore the impact of these changes in
47 melt rate on water mass properties, we also compared with two cross sections across the ice shelf fronts (FRIS February
48 1995 and RIS February 2000) and found temperature differences of less than 0.1°C and salinity differences of less than 0.05
49 psu using this alternative method to represent the tidal effect. These plots are not included here as it is impossible to see the
50 difference compared to Figs. 6 and 7 with the naked eye, and another anomaly plot adds no value to the reader.
51



52
 53 Figure S1. Difference in melt rates (CATS tides - default parameterization) for (a) Filchner-Ronne Ice Shelf, (b) Larsen C
 54 Ice Shelf and (c) Ross Ice Shelf. The results are mean values for the model equivalent period 1995-2009. A positive
 55 difference indicates more melting for the “CATS tides” run in areas of melt in Figure 4, and less freezing in areas of re-
 56 freezing in Figure 4.
 57

58 S2.3 Conclusion

59 This simulation using a two dimensional map of tidal velocities informed by CATS2008 shows minor changes in net melt
 60 flux for each cavity (<10 Gt/yr) and small adjustments in the melt rate pattern (<2 m/yr). These changes are not as large as
 61 one would expect when tides are explicitly simulated as in that case, the basin wide circulation and water mass distribution
 62 would be affected. Explicit tides were not explored in this study as the eORCA1 configuration we use is designed for climate
 63 applications (explicit tides do not fit this purpose as they contribute too much numerical mixing).
 64

65 S3 An evaluation of sea ice production and polynya activity in the NEMO simulations

66 In this section, we analyze polynya activity in the Ronne and Ross polynya regions and explore corresponding changes when
 67 FRIS and RIS cavities are opened.

68 S3.1 Polynya realism in the NEMO simulation without cavities

69 Ice production in the Ronne and Ross polynya regions in the present NEMO v4.2 eORCA1 configuration is found to overall
 70 align well with observed coastal patterns. ‘Ice production’ is diagnosed as the annual integral of sea ice generated over a
 71 domain spanning 73-80 °S and 30-60 °W for the Ronne Polynya region, and 160 °E to 155 °W south of 74 °S for the Ross
 72 Polynya region. When FRIS melt is parameterized (“Closed” run), the Ronne Polynya region produces $24 \times 10^9 \text{ m}^3$ of ice per
 73 year, compared to $58 \pm 21 \times 10^9 \text{ m}^3$ reported from the satellite-based estimates of Nakata et al. (2021). The Ross Polynya
 74 region produces $368 \times 10^9 \text{ m}^3$ of ice per year in the “Closed” simulation, compared to $387 \pm 41 \times 10^9 \text{ m}^3$ reported in Nakata
 75 et al. (2021). It is important to note here that model output and satellite-based estimates are not directly comparable due to
 76 differing definitions for the region of interest between the two sources. If we look at the patterns of sea ice production (Fig.
 77 S2), we see the largest values of around 5 m yr^{-1} at the expected locations along the coasts of Antarctica (Nakata et al.,
 78 2021). Terra Nova Bay Polynya does not correspond exactly to the observed position, likely due to the absence of simulated
 79 landfast sea ice.

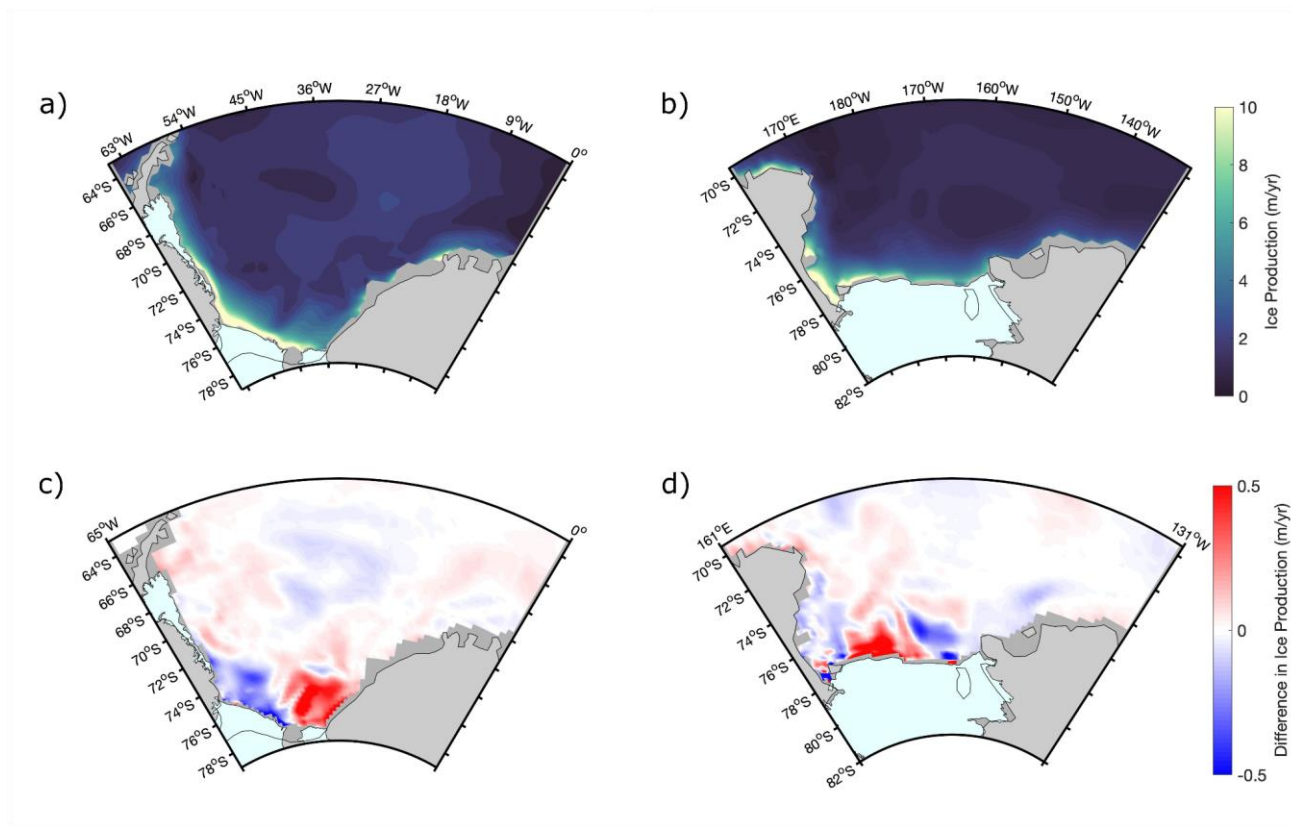
80 S3.2 Impact of explicit sub-ice shelf circulation on polynya activity

81 The changes in polynya activity in response to opening FRIS and RIS are minor. We find no change in the location of
82 polynyas. Ice production does, however, slightly increase from $24 \times 10^9 \text{ m}^3$ in the Ronne Polynya region and slightly
83 decrease from $368 \times 10^9 \text{ m}^3$ in the Ross Polynya region. Ice production slightly decreases to the west of the ice shelf
84 fronts and increases eastward in both analyzed regions when cavities are opened, with changes smaller than 0.5 m yr^{-1} .
85 Changes in ice production are consistent with simulated temperature shifts, with warming to the west and cooling to the east
86 of FRIS and RIS (see Figs. 2 and 3) in the “Open” cavity simulation. Due to the very minor changes in volume of ice
87 production and the absence of a location shift in polynyas, the volume of HSSW produced in each simulation is comparable.
88 The majority of the salinity alterations observed in Figures 2i and 3i are thus likely driven by a change in circulation patterns
89 and conversion of HSSW to ISW when the paths under the ice shelves are opened, and not by an alteration in volume of
90 HSSW produced from polynya activity.

91 S3.3 Summary

92 Sea ice production is reasonable for the two large polynya regions we resolve (Ronne & Ross). Changes in polynya activity
93 due to the opening of the sub-ice shelf cavities can be explained as a response to adjustments in temperature patterns. In
94 conclusion, these effects are minor, do not change the overall locations of the polynyas, and the feedback of sea ice changes
95 on water properties is considered weak.

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98 Figure S2: The annual mean sea ice production in NEMO “Closed” configuration for (a) the Weddell Sea and (b) the Ross
99 Sea. The difference in ice production between the “Open” and “Closed” cavity runs (Open-Closed) are shown in plots (c)
100 and (d) for Weddell and Ross seas respectively.

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102 **Data availability:**

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104 CATS2008 is available for download through the U.S. Antarctic Program Data Center: Data DOI: 10.15784/601235

105 **References:**

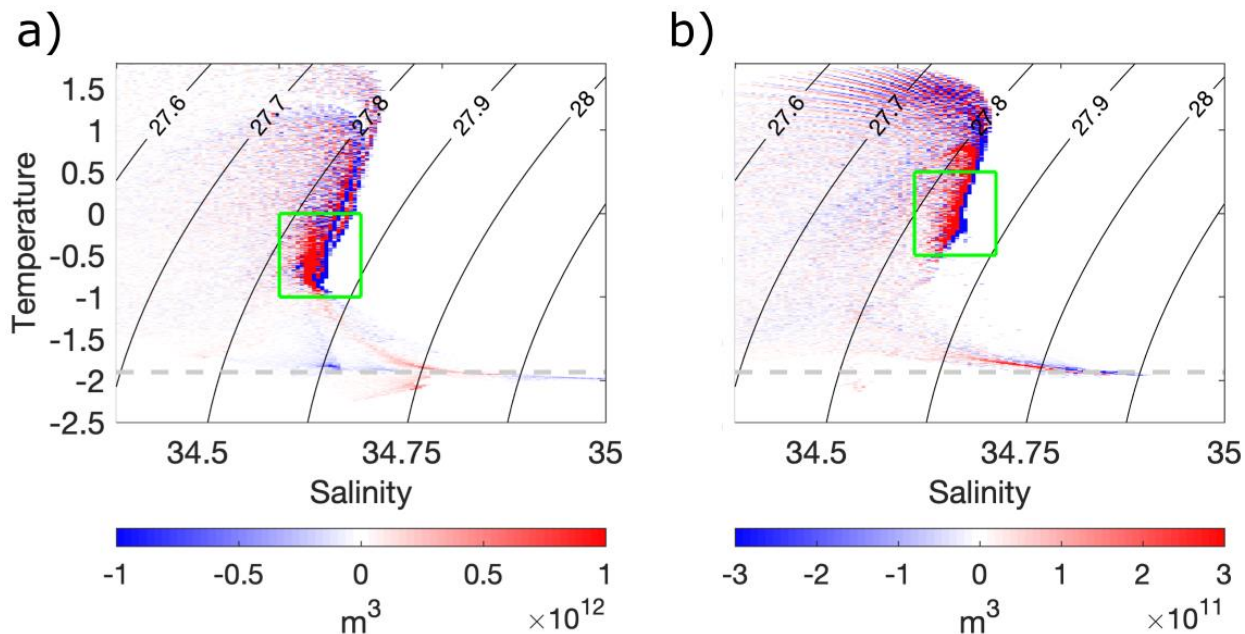
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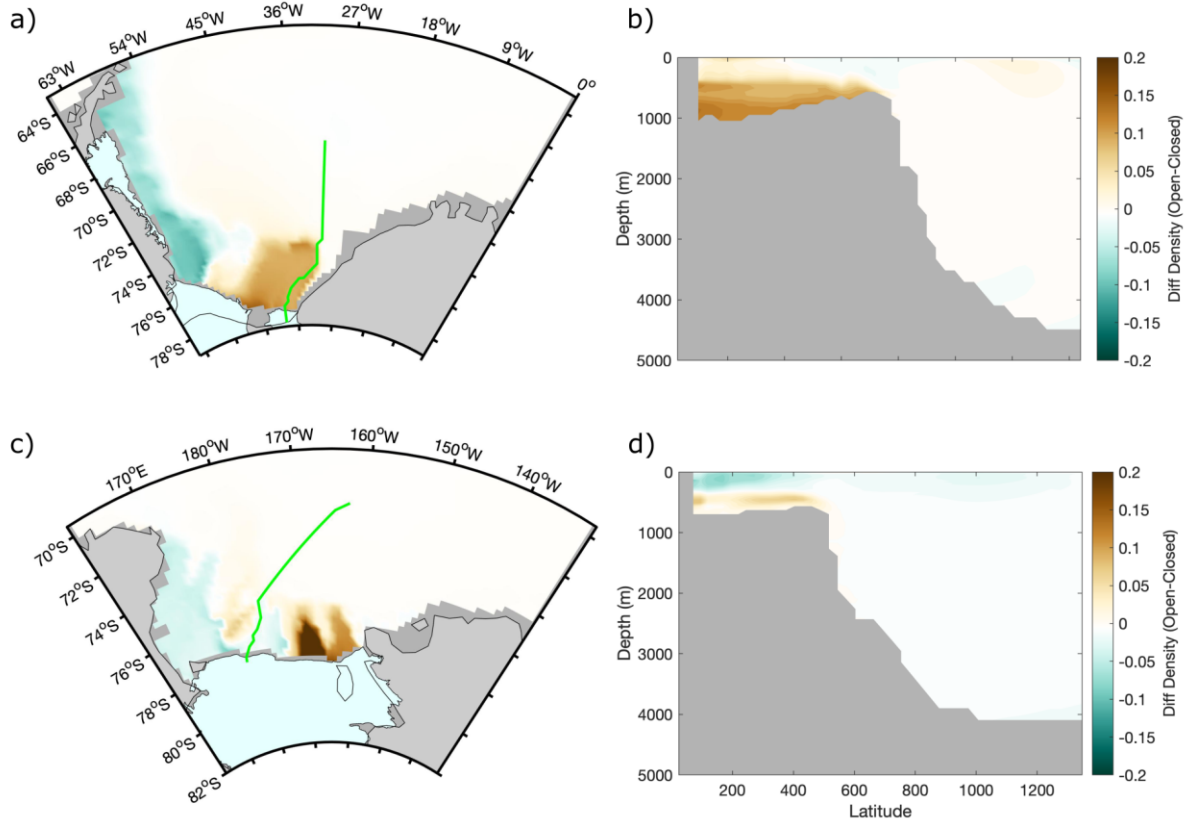
113 **Supplementary Figures**



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Figure S3: Difference (“Open” - “Closed”) in volumetric temperature versus salinity distributions for (a) the Weddell Sea (80 - 60 °S ; 65 °W - 20 °E) and (b) the Ross Sea (85 - 68 °S; 130 °W - 160 °E) for model output excluding data underneath the ice shelves. The scatter dots are placed in T-S space according to their position in the “Closed” cavity simulation and the coloring shows the “Open”-“Closed” volumetric difference. The green boxes delimit the properties corresponding to AABW.



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Figure S4: Density difference (kg m^{-3}) plots for the Weddell (a-b) and Ross (c-d) Seas with bottom values in subplots (a) and (c) and the cross sections of the Filchner and Challenger troughs illustrated by green lines shown in subplots (b) and (d).