



Supplement of

Implementing the Water, HEat and Transport model in GEOframe (WHETGEO-1D v.1.0): algorithms, informatics, design patterns, open science features, and 1D deployment

Niccolò Tubini and Riccardo Rigon

Correspondence to: Niccolò Tubini (niccolo.tubini@unitn.it)

The copyright of individual parts of the supplement might differ from the article licence.

1 Analytical solution Srivastava and Yeh (1991)

1.1 Homogeneous soil

We consider a one-dimension homogeneous soil layer of 1 [m] depth (TP1). The saturated hydraulic conductivity value is assumed to be 1.0 [cm h⁻¹], with $\theta_s = 0.45$ [m³ m⁻³], $\theta_r = 0.2$ [m³ m⁻³], and $\alpha = 0.01$ [cm⁻¹]. The initial condition is determined by imposing as lower boundary condition $\psi = 0$ [m] and a constant water flux at the soil surface $q_A = 0.1$ [cm h⁻¹]. For times greater than 0 the water flux at the soil surface is $q_B = 0.9$ [cm h⁻¹]. Here we compare the performance of WHETGEO-1D against the analytical solution considering different spatial and temporal discretizations. The spatial discretization considered are $\Delta z = 0.001$ [m], $\Delta z = 0.005$ [m], $\Delta z = 0.01$ [m], $\Delta z = 0.05$ [m], and the time step are, $\Delta t = 60$ [s], $\Delta t = 300$ [s], $\Delta t = 900$ [s], $\Delta t = 1800$ [s]. In all the configurations the model accuracy is enhanced by allowing two Picard iterations per time step. Figure (1) shows a comparison between the numerical and the analytical solutions, and Fig. (2) shows a local comparison between the analytical solution and the numerical solution by calculating the pressure potential relative error. As can be seen from Fig. (1) and Fig. (2) the error between the numerical and the analytical solution becomes larger as the time step size increases. The numerical method is first order accurate in time therefore the accuracy, in the sense used in Numerics, decrease as the time step increases.

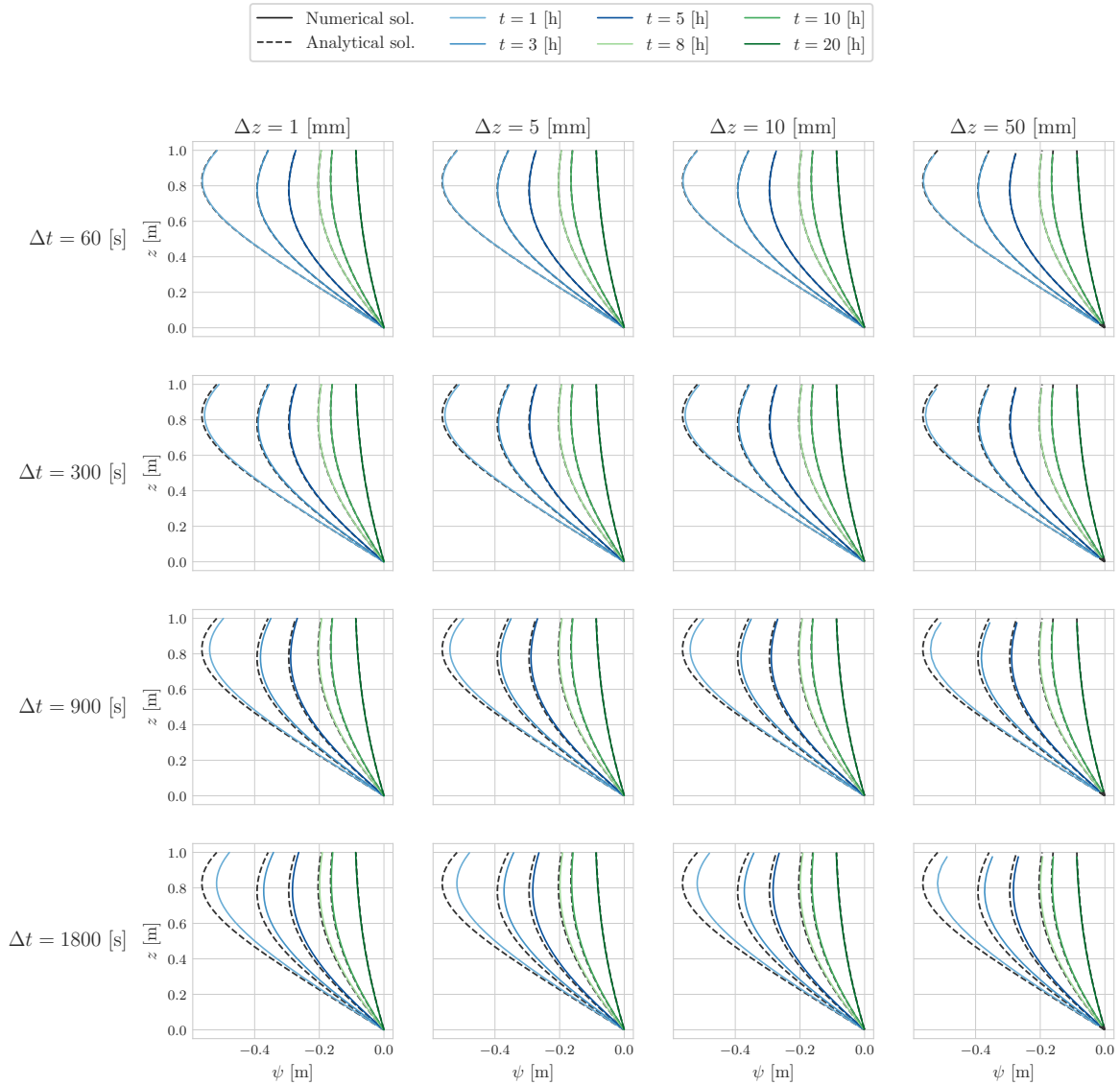


Figure 1: Comparison between the analytical and numerical solutions for the test problem TP1.

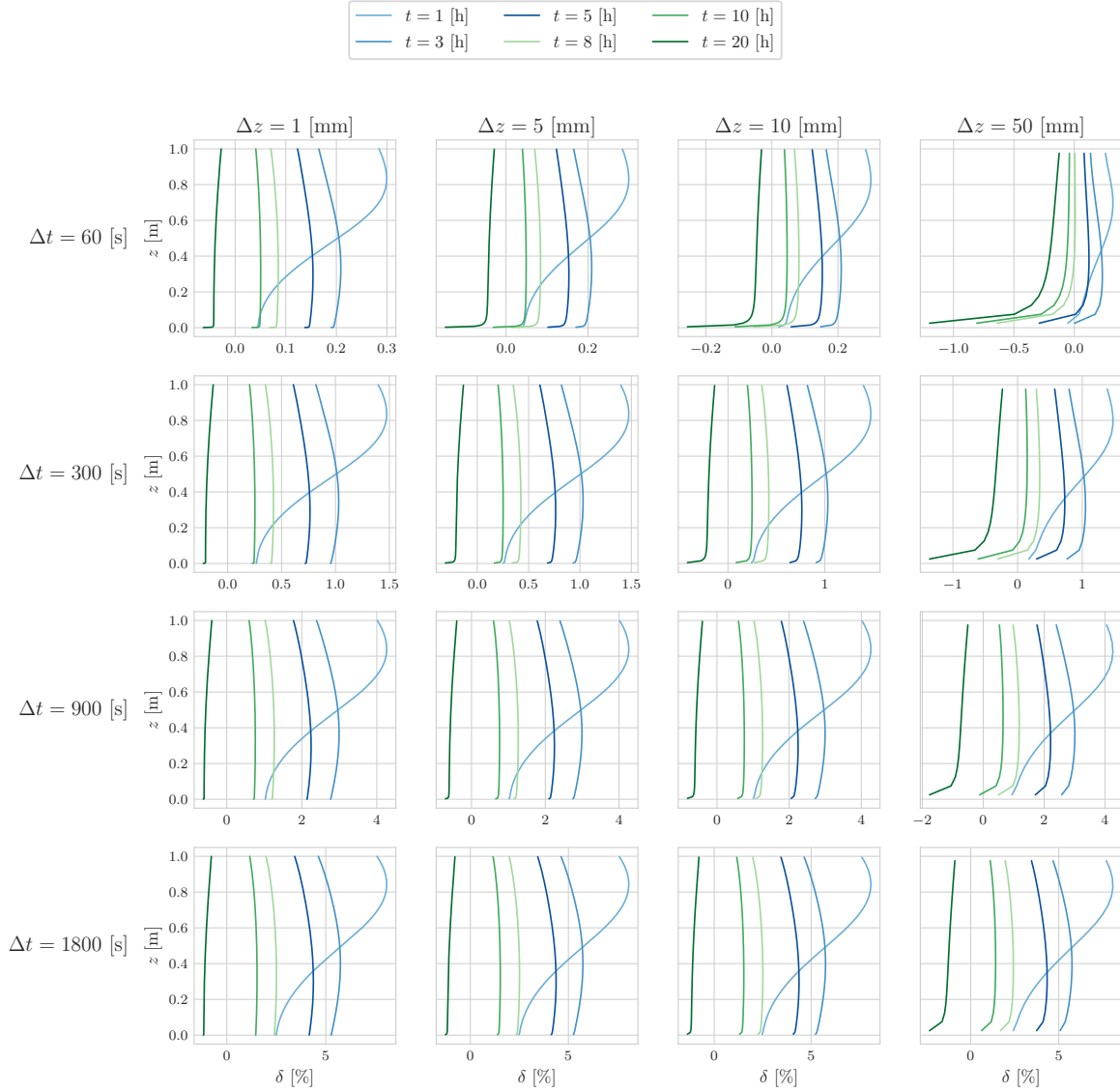


Figure 2: Relative water suction error for the test problem TP1.

2 Effect of temperature on water viscosity

Temperature affects water viscosity, which effectively doubles in passing from 5 to 20 [°C] (Eisenberg et al., 2005), with a positive feedback on the infiltration process. This has been clearly observed in natural systems (Eisenberg et al., 2005; Engeler et al., 2011; Ronan et al., 1998) where infiltration rates follow diurnal and seasonal temperature-cycles. In fact, according to (Muskat and Meres, 1936), the unsaturated hydraulic conductivity can be expressed as

$$K(\theta) = \kappa_r(\theta) \kappa \frac{\rho g}{\nu} \quad (1)$$

where $\kappa_r(\theta)$ [-] is the relative permeability, κ [L²] is the intrinsic permeability, ρ [L³M⁻¹] is the liquid density, g is the acceleration of gravity, and ν [L²T⁻¹] is the kinematic viscosity of the liquid. Thus, for constant θ , variations in $K(\theta)$ due to temperature can be accounted as (Constantz and Murphy, 1991):

$$K(\theta, T_2) = K(\theta, T_1) \frac{\nu(T_1)}{\nu(T_2)} \quad (2)$$

In Eq. (2), T_1 is a reference temperature while T_2 is the soil water temperature. Based on this equation the value of $K(\theta)$ increase by over three-fold between 5 to 60 [°C] (Constantz and Murphy, 1991).

Below we repeat the synthetic experiment presented in Section C3.1 C3.2 regarding respectively the infiltration excess and saturation excess processes, taking into account different soil temperatures. In these simulation temperature is kept constant over time.

2.1 Infiltration excess

In this numerical experiment we consider a homogeneous soil of 3 [m] depth. Soil hydraulic properties are described with the Van Genuchten’s model, Table (1).

Table 1: Hydraulic properties of the silty clay loam soil Bonan, 2019 for the Horton runoff numerical experiment.

θ_r [$\text{m}^3 \text{m}^{-3}$]	θ_s [$\text{m}^3 \text{m}^{-3}$]	α [m^{-1}]	n [-]	K_s [m s^{-1}]
0.089	0.43	1.0	1.23	$1.9447e - 07$

The initial condition is assumed to be hydrostatic with $\psi = 0$ [m] at the bottom. The surface boundary condition is a synthetic rainfall, as in Fig. (3) (a), lasting 15 [min] with constant intensity of 0.028 [mm s^{-1}]. At the bottom we prescribed a Dirichlet boundary condition with constant $\psi = 0$ [m] so the transient is driven only by the surface boundary condition. For the soil temperature we consider the following values: 5 [°C], 10 [°C], 15 [°C], 20 [°C], 25 [°C], 30 [°C], where the reference value is 20 [°C]. As temperature is lower the reference value of $T = 20$ [°C] the infiltration is slower

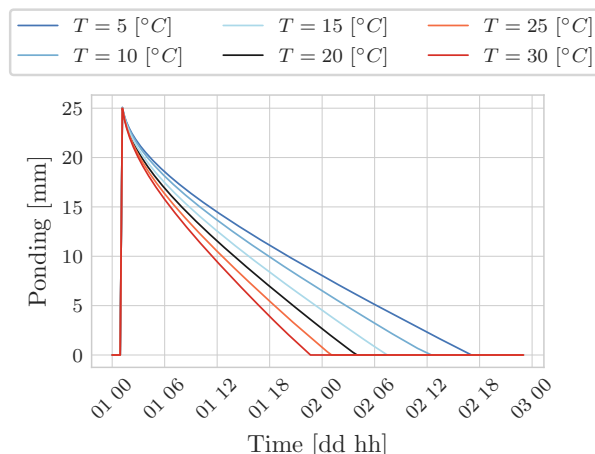


Figure 3: Comparison of the water ponding time evolution considering different soil temperatures. The temperature in these simulations is uniform and kept constant over time. The reference value of for temperature is $T = 20$ [°C]. As soil temperature is lower the reference value the infiltration is lower, vice-versa when temperature is higher.

because the hydraulic conductivity is lower the reference value, whereas when temperature is higher the infiltration is faster. Although the saturated hydraulic conductivity does not change by order of magnitude with temperature, the relative small variations in its values are amplified by the capillary gradient.

2.2 Saturation excess

We consider a layered soil of 3 [m] depth. The thicknesses of the loamy sand layer and the clay layer are, respectively, 0.3 [m], and 2.7 [m]. The soil hydraulic properties are described with the Van Genuchten’s model, Table (2). The initial condition is assumed to be hydrostatic with $\psi = -2$ [m] at the bottom. The surface boundary condition is a synthetic rainfall Fig. (4) (a), at the bottom we prescribed a Dirichlet boundary condition with constant $\psi = -2$ [m] so the transient is driven only by

Table 2: Hydraulic properties of the loamy sand layer and clay layer, respectively, Bonan, 2019 for the numerical experiment on Dunnian runoff due rainfall.

θ_r [$\text{m}^3 \text{m}^{-3}$]	θ_s [$\text{m}^3 \text{m}^{-3}$]	α [m^{-1}]	n [-]	K_s [m s^{-1}]
0.057	0.41	12.4	2.28	$4.0528e - 05$
0.068	0.38	0.8	1.09	$5.5556e - 07$

the surface boundary condition. For the soil temperature we consider the following values: 5 [°C], 10 [°C], 15 [°C], 20 [°C], 25 [°C], 30 [°C], where the reference value is 20 [°C].

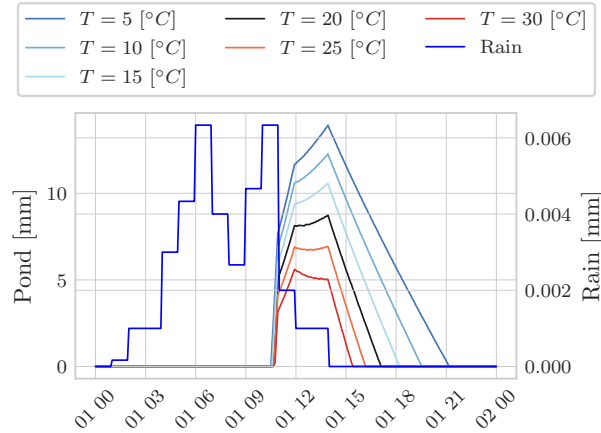


Figure 4: In this numerical experiment the run-off varies with soil temperature. As the soil becomes colder the maximum for water ponding increases and water takes more time to infiltrate.

As temperature increase the ponding water at the soil surface is smaller. The depth of ponding water when $T = 30$ [°C] is almost one third compared to that for $T = 5$ [°C].

3 GEOframe Community Publication Policy (GCPP v1.0)

3.1 Introduction

GEOframe-NewAge is an open-source, semi-distributed, component-based hydrological modeling system. It was developed in Java and based on the environmental modeling framework Object Modeling System V3 (OMS3).

The core of the project was born from the idea of Professor Rigon and mainly developed at the Department of Civil, Environmental and Mechanical Engineering of University of Trento, Italy. During the last decade, GEOframe community grew and now is made of many scientists around the world who share their work, codes, knowledge and experiences for the benefit of all GEOframe users, whilst pursuing their individual research interests and careers.

Therefore, giving appropriate credits for the intellectual input through co-authorships or citations should be the proper functioning of the community.

This document sets out how members of the GEOframe community should recognise the intellectual contribution of the GEOframe community's members.

3.2 Principal web references

In the following, the principal GEOframe web references, where you can find the latest achievements, developments, publications, code versions, courses and ideas are reported:

- <http://abouthydrology.blogspot.com/search/label/GEOFRAME>
- <http://geoframe.blogspot.com/>
- <https://github.com/geoframecomponents>
- <https://github.com/GEOframeOMSProjects>
- <https://osf.io/fk8ta/>

3.3 General principles

The formal, legal conditions that govern the use of GEOframe at present are given by the G.P.L. v 3. Each GEOframe component can have its own license though.

This Policy applies to all uses of GEOframe products, including but not limited to data and computer code, for research and teaching. It is not intended to restrict what can be done with them, rather to ensure appropriate acknowledgement and communication between users and developers. This policy will be updated regularly.

A developer is any person whose expertise has either significantly influenced the design of GEOframe code or who has written code, with no distinction between scientific and technical inputs. Developers are encouraged to publish their work in reasonable time, while potential users should approach developers early in their study to avoid duplication or wasted effort on new developments. Developers may reserve the right for the first scientific application of their scheme and will be able to advise if and when co-authorship, citation or acknowledgement is appropriate.

A list of new developments and the scientists responsible for them will be maintained on the GEOframe website. These contributions should be recognised by citations.

When writing the source code of a component, GEOframe developers should consider the following:

- Provide a brief description of what the program does.
- State the authors of the code and the following modifiers.
- Describe the input required to run the component and its output.
- Some notes concerning the limitations, and the algorithms used within the component. A wish-list for the future version and/or information.
- Articles or books which have inspired the codex or justified its necessity. Users are encouraged to cite these papers in their own work

- If you want to contribute code or documentation, create pull requests, we will consider them.

Ideally a committed code should conform to the rules required by Joss.

Acknowledgments should be considered for a wider list of scientists who contributed to the modelling system, but whose contributions may not be documented in publications. A list of such scientists will be maintained on the GEOframe web page.

When writing a paper, GEOframe users and developers should consider the following:

Co-authorship

- Is expected if your published research benefited from a new development, i.e. the development influenced your study to the extent that it was discussed in the paper.
- Is expected if your research required substantial direct input from a developer, e.g. to make substantial modifications to the code that you used, to help design the experiments etc.
- Should be considered for a wider list of scientists who contributed to the modelling system, but whose contributions may not be documented in publications. A list of such scientists will be maintained on the GEOframe web page

Acknowledgements

- Should be considered for scientists involved in GEOframe code developments that have become established.

Citation of a published paper

- Is expected if a citable paper describing a development exists. A narrative description of the model and a list of papers describing developments will be maintained on the GEOframe web page.

Please be generous in offering credit for other people's work, as everyone benefits in the end. Use best judgement and, if in doubt, error the side of inclusiveness.

References

- Bonan, G. (2019). *Climate change and terrestrial ecosystem modeling*. Cambridge University Press.
- Constantz, J., & Murphy, F. (1991). The temperature dependence of ponded infiltration under isothermal conditions. *Journal of Hydrology*, 122(1-4), 119–128.
- Eisenberg, D., Kauzmann, W., & Kauzmann, W. (2005). *The structure and properties of water*. Oxford University Press on Demand.
- Engeler, I., Franssen, H. H., Müller, R., & Stauffer, F. (2011). The importance of coupled modelling of variably saturated groundwater flow-heat transport for assessing river–aquifer interactions. *Journal of Hydrology*, 397(3-4), 295–305.
- Muskat, M., & Meres, M. W. (1936). The flow of heterogeneous fluids through porous media. *Physics*, 7(9), 346–363.
- Ronan, A. D., Prudic, D. E., Thodal, C. E., & Constantz, J. (1998). Field study and simulation of diurnal temperature effects on infiltration and variably saturated flow beneath an ephemeral stream. *Water resources research*, 34(9), 2137–2153.