



## Supplement of

## How well can inverse analyses of high-resolution satellite data resolve heterogeneous methane fluxes? Observing system simulation experiments with the GEOS-Chem adjoint model (v35)

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**Figure S1.** Monthly TROPOMI sampling density at  $0.25^{\circ} \times 0.3125^{\circ}$  resolution (~25 km) between May 2018 and April 2019, after filtering for data quality and clouds. The total number of observations, and percent over-land grid cell data coverage, is indicated in each panel.



**Figure S2.** Methane column concentration differences for 2018-08-01 resulting from the individual model transport errors employed in the OSSE (see Sect. 2 for details). Shown are differences incurred from: a) using 6 versus 3 buffer grid cells at the domain boundary; b) averaging over 13:00-14:00 LT versus sampling the model instantaneously at the satellite overpass time; c) employing non-local versus full PBL mixing schemes; d) alternate convection and tropopause treatments; and e) all model transport errors. Shown for comparison are f) the column differences that arise from employing the true versus spatially biased prior emissions. The root-mean-square errors (RMSE) relative to the true fluxes are labeled in each panel.



**Figure S3.** Cost function analysis and determination of the regularization parameter  $\gamma$  based on one-week inversions with spatially uniform prior errors. Panel a) shows the L curve comparing the prior and observational deviation terms in the cost function as a function of  $\gamma$ , following the method in Hansen and O'Leary (1993). As shown in Eq. 1 of the main text, the prior term is given by  $(x - x_a)^T S_a^{-1}(x - x_a)$  and the observational term is given by  $(y - F(x))^T S_{obs}^{-1}(y - F(x))$ . Panel b) shows the prior term divided by the total cost function computed at  $\gamma = 1000 (J_{1000}$ , where the solution is mostly determined by the observations; blue line), the observational term divided by the total cost function computed at  $\gamma = 0 (J_0$ , where the solution is solely determined by the prior; red line), and the sum of the blue and red lines (in grey).



**Figure S4.** Initial guess and optimized emissions for each inversion framework. Labels inset indicate the domain-wide total emissions and spatial correlation to the true fluxes.

**Emission increments** 



Figure S5. Emission increments added to the prior fluxes in the V-OBSGuess inversions. See Sect. 4 for details.



Figure S6. Same as Figure 8, but with prior and posterior results degraded to 4°×5° horizontal resolution.



Figure S7. Same as Fig. 4, but showing results with model transport error.

Parameter	Range
Solar Zenith Angle (SZA)	$\leq 70^{\circ}$
Viewing Zenith Angle (VZA)	$\leq 60^{\circ}$
Surface albedo	$\geq 0.02$
Aerosol Optical Thickness (AOT)	< 0.3
XCH <sub>4</sub> precision (noise-related error)	< 10 ppb
Signal-to-noise ratio	$\geq$ 50
$\chi^2$	< 100
Fraction of non-corrupted/unphysical spectral pixel <sup>1</sup>	$\geq 70\%$
Cloudiness level	Confidently clear <sup>1</sup>

## Table S1. Filters applied for TROPOMI data quality assurance

<sup>1</sup>Pixel quality is determined per orbit, for details see the TROPOMI Product User Manual (2019).

## Reference

- Hansen, P. C., and O'Leary, D. P.: The Use of the L-Curve in the Regularization of Discrete Ill-Posed Problems, SIAM Journal on Scientific Computing, 14(6), 1487-1503, 10.1137/0914086, 1993.
- TROPOMI product user manual: <u>http://www.tropomi.eu/data-products/methane</u>/, last access: 15 April 2019.